
#### Abstract

Title of Document: QUANTIFICATION OF PERMEABILITY- POROSITY RELATIONSHIPS IN SEAFLOOR VENT DEPOSITS: DEPENDENCE ON PORE EVOLUTION PROCESSES

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Hydrothermal mineral deposits formed along seafloor spreading centers help regulate the transfer of heat and mass from Earth's interior to the oceans. Aqueous fluids circulate within the seafloor and are emitted through vent deposits, formed from interaction between vent fluids and seawater. These deposits evolve as they react physically and chemically with venting fluids and seawater, therefore changing transport properties, such as permeability and porosity. In this study, measurements of permeability ( $k$ ) and porosity ( $\phi$ ) were used in conjunction with microstructural observations to identify evolution of permeability-porosity relationships (EPPRs) for vent deposits. EPPRs are power-law relationships relating permeability and porosity through an exponent, $\alpha$, which is sensitive to changes in these properties. These relationships are important for understanding pore evolution processes and fluid distribution, in addition to their effects on environmental conditions within vent deposits.


# QUANTIFICATION OF PERMEABILITY-POROSITY 

 RELATIONSHIPS IN SEAFLOOR VENT DEPOSITS: DEPENDENCE ON PORE EVOLUTION PROCESSESBy<br>Jill Leann Gribbin

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## Chapter 1: Introduction

Seafloor hydrothermal vents provide a gateway between the Earth's interior and the oceans, regulating the flow of heat and mass between them. Vents facilitate both conductive and convective heat transport, making them important for maintaining environmental conditions within the oceans [Lowell, 1991; Wilcock and Delaney, 1996]. Heat from magma chambers associated with seafloor spreading centers is easily conducted through seafloor rock [Lowell et al., 1995]. The convective transfer of heat, however, is dependent on the flow of hydrothermal fluids and the circulation of seawater from the seafloor to deep within the crust. This process relies on the percolation of cold seawater down into the crust through broad recharge zones in forms of seafloor faults, dikes, fissures, and natural basaltic pore spaces [Sleep, 1991; Tivey et al., 1995]. As the water travels downward possibly reaching Moho depths [Lowell and Germanovich, 2004], it becomes hotter and loses many of its original dissolved ions, such as $\mathrm{Mg}^{2+}$ and $\mathrm{SO}_{4}{ }^{2-}$, while gaining various sulfides and oxides (i.e. Fe and Mn oxides). This chemically altered fluid becomes increasingly hot and buoyant at which point it begins to ascend to the seafloor [Von Damm et al., 1998; Henderson et al., 2005]. As the fluid rises through the subsurface and then to the seafloor, interactions with cold seawater will lower fluid temperatures. Changes in temperature, in addition to a change in fluid pressure, result in mineral precipitation and the formation of hydrothermal vent deposits.

Vent fields are found on the seafloor along mid-ocean ridges and back-arc basins; however, environmental conditions present at vent fields can vary. Despite
differences between many vent fields, they each generally have regions of both high-temperature and lower-temperature deposits [Elderfield and Schultz, 1996]. High-temperature areas (Figure 1.1a) are generally characterized by large black smoker chimneys, which form from the rapid ascension (1-5 m/s) of $350-400^{\circ} \mathrm{C}$ fluid emission [Haymon, 1983; Delaney et al., 1992; Tivey et al., 1995]. The extreme thermal gradient existing between this fluid and the $2^{\circ} \mathrm{C}$ seawater it enters causes the precipitation of sulfides, anhydrite, and amorphous silica that build upwards creating the chimney structure [Haymon, 1983; Tivey and McDuff, 1990]. Black smokers tend to be situated closer to the ridge axis [German and Parson, 1998], where fluids can be subjected to higher temperatures within the crust and also have greater channelization due to near axis faults, which can accommodate the fluid's high speed.


Figure 1.1: Schematic diagrams of seafloor vent deposits. a) Black smoker. High-temperature fluid emits through central chimney conduit with fluid diffusely transferred from structure sides; b) White smoker. Low-temperature fluid percolates through small branching channels. Vent deposits usually lack a welldefined central conduit.

Lower temperature diffuse vents (Figure 1.1b) are also very common at vent fields and possibly elsewhere throughout the seafloor [Delaney et al., 1992;

Scheirer et al., 2006]. These deposits form in areas where fluid within the seafloor is only weakly channeled. Consequently, the hydrothermal fluid will entrain, and mix with colder seawater, resulting in precipitation of metal sulfides and anhydrite [Tivey et al., 1995; Mills et al., 1996]. Fluid that is diffusively expelled will generally have a temperature less than $150^{\circ} \mathrm{C}$ [Pester et al., 2008]. Although the fluid temperature is lower than that of black smokers, diffuse deposits are believed to transfer several times more heat to the oceans than chimney deposits due to their expansiveness on the seafloor [Rona et al., 1993; Elderfield and Schultz, 1996; Juteau and Maury, 1999]. The amount of flow exiting through these deposits, however, is poorly constrained, making it important to study parameters that impact this flow.

Physical properties that influence the evolution of both high- and lowtemperature vent deposits include permeability and porosity. Permeability $(k)$ is the ability of a material to transmit fluid, while porosity $(\phi)$ is the volume fraction of void space in a material (Figure 1.2) [Norton and Knapp, 1977]. Permeability is dependent upon porosity, as well as pore geometry and connectivity.


Figure 1.2: Permeability and porosity are properties that vary in materials - differences between the two properties for the same medium are illustrated, a) permeability is indicated by green arrows representing available flow pathways through connected pores, b) the medium's porosity, shaded in green, includes all void space regardless of connectivity.

Depending on the location of a vent deposit and its exposure to hydrothermal fluids, permeability and porosity of a deposit will vary. Fluid interactions can lead to a number of processes, such as mineral dissolution, precipitation, and thermal cracking, which can significantly alter pore space and thus affect permeability. Conversely, changes in permeability and porosity will affect the mixing and distribution of vent fluids. The extent to which chemical processes resulting from fluid interactions alter permeability and porosity and thereby impact fluid flow within vent deposits is not fully understood.

In this project, permeability and porosity of vent deposit samples were measured to determine the relationships between permeability and porosity in vent deposits. Various types of vent deposits from many different vent sites were studied. Such data for the same type of vent deposits at different stages provide an evolution of permeability-porosity relationships (EPPRs) [Zhu et al., 2007]. Identified EPPRs can be used to constrain flow properties of vent deposits, which are crucial in modeling how vents evolve over time.

### 1.1 Samples Measured in Study

The seafloor vent deposits were divided into five groups based on sample type. The first sample group was composed of massive anhydrite deposits (Figure 1.3a). Anhydrite $\left(\mathrm{CaSO}_{4}\right)$ forms as a result of increasing seawater temperature. As the seawater is heated above $\sim 150^{\circ} \mathrm{C}$ by hot rock and fluid, dissolved $\mathrm{CaSO}_{4}$ becomes less soluble and begins to precipitate as anhydrite (at depths near 3000 m) [Tivey et al., 1995; Kuhn et al., 2003]. It commonly forms within seafloor cracks and on the seafloor where it forms as massive deposits or develops as the
framework structure necessary for the development of metal-rich chimney structures [Haymon, 1983; Tivey and McDuff, 1990].

The second group of samples consists of portions of flanges, slabs, and crust. Flanges (Figure 1.3b) are deposits that form horizontally like tiers or ledges off the sides of larger vent structures from the precipitation of sulfides, sulfates, and possibly carbonates. Vent fluid that is emitted radially from a primary chimney structure [Goldfarb, 1988; Tivey et al., 1999] will pool under the flange until it gradually flows out from under the flange or is diffused upwards through the flange [Kerr, 1997]. Conductive cooling of fluids flowing through flanges results in the precipitation of amorphous silica along grain edges [Tivey et al., 1999]. Hydrothermal slabs (Figure 1.3c) are found at the Lucky Strike Vent Field along the Mid-Atlantic Ridge (MAR). They are layered silicified volcanic deposits (i.e. hydrothermally cemented breccias) rich in sulfides, barite, silica and basalt fragments [Langmuir et al., 1997; Rouxel et al., 2004]. Fluid trapped beneath slabs will circulate and gradually diffuse through cracks in the slab. A crust deposit from the Trans-Atlantic Geotraverse (TAG) field active mound along the MAR will also be included in this group (Figure 1.3d). TAG crust is notable for its brecciation and incorporation of re-cemented vent debris [Thompson et al., 1985; Humphris et al., 1995].

The remaining three groups are composed of Zn -rich actively diffusing spires, black smoker chimneys and relict spires, respectively. The group of Zn rich actively diffusing spires (Figure 1.3e), which includes white smoker chimneys, consists of spires that form from the emission of roughly $250-300^{\circ} \mathrm{C}$
fluid. This fluid is conductively cooled at depth where it will precipitate much of its sulfide components [Hannington et al., 1995; Humphris et al., 1995]. Fluids from these spires are depleted in sulfur, but enriched in zinc compared to fluids from higher temperature deposits. The zinc abundance is due to both the dissolution of sphalerite and remobilization of zinc at depth, which increases the fluid's zinc concentration. As the fluid is emitted the zinc is re-precipitated as sphalerite in the deposit [Humphris et al., 1995]. These Zn -rich diffusing spires lack a central conduit, but instead contain numerous small channels through which fluid can be passed. In contrast, the black smoker chimneys, discussed earlier in the paper, (Figure 1.3f) rely on the precipitation of anhydrite, sulfides, and silica resulting from the interaction of high-temperature fluid $\left(350-400^{\circ} \mathrm{C}\right)$ with seawater on the seafloor. For these structures to develop, an initial ring of anhydrite must first precipitate on the seafloor. As fluid is channeled through the anhydrite structure, sulfide minerals precipitate along the inner walls, forming a well-defined central conduit [Haymon, 1983; Tivey and McDuff, 1990]. Additionally, pore spaces within the outer anhydrite layers gradually infill due to this sulfide precipitation. Relict spires (Figure 1.3 g ) are spires that were not actively venting when they were recovered from the seafloor. Their inactivity presumably results from pathways becoming blocked by mineral precipitation over time. Because there are many different types of samples being measured in this study, a large range of both permeability and porosity values is expected.


### 1.2 Geologic setting

In this study, permeability and porosity of vent samples recovered from ten different vent fields have been measured (Figure 1.4). Each location has variations in environmental conditions that are important to consider when linking the effects of pore evolution processes on the evolution of permeability and porosity within the deposits.


Figure 1.4: Vent field locations (red circles) from which deposit samples were recovered. From left to right: Fenway, Kilo Moana \& ABE, Juan de Fuca (includes Main Endeavour Field and Cleft), Guaymas Basin, Trans-Atlantic Geotraverse (TAG), Lucky Strike, and Central Indian Ridge (CIR).

A few samples have been collected from the Fenway vent field in the Manus back-arc basin in the Bismarck Sea near Papua New Guinea. The vent field is dominated by a large two-tiered mound upon which several hightemperature black smoker chimneys have developed [Craddock and Bach, 2010]. Outer portions of the mound are covered in sulfide chimney debris, massive anhydrite outcrops, and hydrothermal sediment, which accommodate lower temperature diffuse flow [Craddock and Bach, 2010]. Recovered anhydrite from

Fenway tends to be coarse-grained, suggesting that the anhydrite likely precipitates within cavities accessible to fluid flow that are capped by a less permeable sulfide-rich layer. Over time, these sulfide layers collapse, exposing the anhydrite deposits [Craddock and Bach, 2010].

There is one sample that has been recovered from the Edmond vent field along the Central Indian Ridge (CIR) nearby the Rodriguez Triple Junction. The vent field is partially covered by lava flows, pillow basalts, and sediment and is believed to have experienced robust magmatism for about 2 Myr [Kumagai et al., 2008]. Vent fluids sampled from the Edmond field show low $\mathrm{Na} / \mathrm{Cl}$ ratios (though fluid is enriched in both Na and Cl ) and high $\mathrm{Ca} / \mathrm{Cl}$ ratios, which can be explained by ongoing albitization of subseafloor rocks [Gallant and Von Damm, 2006]. Measured temperatures from spires show typical values, with high temperatures around $380^{\circ} \mathrm{C}$ [Kumagai et al., 2008].

Two carbonate-rich vent samples have been recovered from the Guaymas Basin vent field in the Gulf of California. Guaymas Basin is notable for its rapid rate of sediment deposition and high concentrations of dissolved carbon compounds that result from the interaction of fluids with organic-rich sediments [Pearson et al., 2005]. Previous work has shown that both sulfides and calcite/barite are large constituents of hydrothermal deposits at Guaymas [Koski et al., 1985]. Sampled vent fluids are depleted in $\mathrm{Zn}, \mathrm{Cu}$, and Fe due to sediment and subsurface reactions, which is unusual for sulfide deposits [Koski et al., 1985]. Fluid temperatures and alkalinity are largely consistent with other vent fields.

Four samples will be measured from the Lucky Strike hydrothermal field along the MAR. Lucky Strike is dominated by a large volcanic seamount that has deposited multiple lava flows across the area. Many different vent structures are present at Lucky Strike: high- and low-temperature chimneys, relict spires, flanges, anhydrite deposits, and slabs [Langmuir et al., 1997]. Sampled vent fluids exhibit slightly lower silica concentrations than found in most mid-ocean ridge vent fields and also lower hydrogen sulfide contents [Langmuir et al., 1997]. Fluid temperatures depend on location within the vent field and appear to change over time-scales as short as a few days, suggesting influences from seafloor eruptions.

Moving further south along the MAR, the TAG vent field has a large central mound constructed from the build-up and collapse of hydrothermal material situated atop an area of brecciated seafloor [Humphris et al., 1995]. Strong mineralogical zoning is present on the TAG mound, with hightemperature, chalcopyrite-anhydrite rich chimneys clustered predominantly at the top of the mound. Lower-temperature Zn -rich diffusing spires are common lower down on the sides of the mound [Humphris et al., 1995]. Extensive anhydrite veining is also present at TAG as a result of seafloor crack infill. The anhydrite abundance indicates that the TAG mound environment regularly maintains seawater temperatures $>150^{\circ} \mathrm{C}$ to allow for retrograde anhydrite precipitation [Humphris et al., 1995]. Multiple samples from the TAG field have been measured in this study.

Along the Juan de Fuca Ridge (JFR), several vent fields have been discovered from which vent samples have been included in this study. A few vent samples have been collected from the Main Endeavour Field (MEF) located along the northern portion of the JFR. The MEF is characterized by large sulfide chimneys, areas of diffuse flow, and many inactive spires and flanges [Tivey et al., 1999]. Vent fluids from MEF are unusual in that they have a high pH (4.24.5), exhibit a strong temperature and chlorinity gradient across the field, and have high concentrations of methane and ammonia [Tivey et al., 1999]. Samples from the Cleft segment of the southern JFR have also been measured. The Cleft segment is focused around an untectonized lava plain that houses three major areas of high temperature vents [Embley and Chadwick, 1994]. Both hightemperature and diffuse deposits have been detected along the segment, with most found along seafloor fissures. A large majority of the vents along the Cleft segment are not actively venting; it is believed that many of them vented in accordance with lava producing seafloor eruptions [Embley and Chadwick, 1994]. These older deposits have high silica content and pyritic mineralogy [Embley and Chadwick, 1994].

Samples have also been recovered from both the Kilo Moana and ABE vent fields located along the Eastern Lau Spreading Center, located near the islands of Fiji and Tonga. Seafloor bathymetry of Kilo Moana shows two broad low relief volcanic domes cross-cut by fissures where vent structures have formed [Ferrini et al., 2008]. Both high-temperature and diffuse vent structures have been detected there. The ABE vent field is highly faulted with evidence for several lava
flows and heavy sedimentation [Ferrini et al., 2008]. Most venting spires are surrounded by diffusely venting structures. Flanges forming off the sides of spires and occasionally off of lava flows are common. Between the two vents fields, ABE is slightly shallower and less acidic (4.3-4.9) than Kilo Moana [Ferrini et al., 2008].

## Chapter 2: Pore Evolution Processes and Permeability Change

There exists no single relationship between permeability and porosity that is applicable to all materials. However, correlations can be made for materials under specific conditions. Permeability $(k)$ can be related to porosity $(\phi)$ through a power-law relationship (Equation 2.1) whereby the exponent, $\alpha$, is sensitive to changes in a material's pore structure [e.g., Turcotte and Schubert, 1982; Zhu et al., 1999; 2007; Bernabé et al., 2003]. The value $k_{0}$ is the permeability at a reference porosity $\phi_{0}$ [Zhu et al., 1995].

$$
\begin{equation*}
\left(\frac{k}{k_{0}}\right) \propto\left(\frac{\phi}{\phi_{0}}\right)^{\alpha} \tag{2.1}
\end{equation*}
$$

This relationship can be depicted as a line in $\log (k)$ vs. $\log (\phi)$ plot where $\alpha$ is the slope of this line [Bernabé et al., 2003]. Higher $\alpha$ values (i.e., steeper slopes) represent greater changes in permeability with respect to changes in porosity.

A number of diagenetic processes (e.g. compaction, dissolution) can alter a materials pore structure; these processes define the evolution of permeabilityporosity relationships (EPPRs). EPPRs can be used to provide a convenient description of how transport properties evolve during a specific diagenetic process.

A variety of theoretical [e.g., Paterson, 1983; Walsh and Brace, 1984], numerical [e.g., Steefel and Lasaga, 1994; Quispe et al., 1995; Zhu et al., 1995, 1999], and experimental studies [Bernabé et al., 1982; Bourbie and Zinszner, 1985; Zhang et al., 1994] have examined permeability and porosity under
different conditions and have found a range of EPPRs [Guéguen and Palciauskas, 1994; Bernabé et al., 2003]. Deep sea hydrothermal vent deposits are products of the interaction between aqueous fluids and sea water. The vent structures, in turn, exert important control of fluid flow distribution. To understand this coupled system, it is important to quantitatively characterize EPPRs of vent deposits from various sites. Several physical and chemical processes that can significantly alter pore structures have been described in previous studies on sedimentary rocks. These processes are pertinent to vent formation and thus will be summarized here.

A simple analog of a porous material is a cubic matrix embedded with identical tubes (Figure 2.1). In this analog the tubes constitute all of the material's pore space. Assuming laminar flow conditions, permeability can be related to porosity with an $\alpha$ value of 2 [Turcotte and Schubert, 1982]. If the tubes are replaced by cracks where the apertures of the tubes differ from their diameters, then permeability and porosity can be related with an $\alpha$ value of 3 [Guéguen and Palciauskas, 1994]. While the isotropic and homogeneous tube model is idealized


Figure 2.1: Diagram of a simple cubic matrix of circular tubes. This pore geometry yields a permeability-porosity relationship with an $\alpha$ value of 2 . This relationship is a good reference for considering more complex EPPRs.
compared to those of true materials, it provides a good frame of reference from which permeabilityporosity relationships of more complex pore networks can be estimated. Understanding these basic EPPRs provide a good foundation for exploring the effects of different pore evolution processes. Next, the effects of mechanical compaction, hot isostatic pressing, thermal
cracking, precipitation, and dissolution, all of which are processes pertinent to the formation of vent deposits, on pore structure are discussed, in addition to how they result in permeability-porosity relationships significantly deviating from the simple tube model (Table 2.1).

| Processes | Materials | $\alpha$ |
| :---: | :---: | :---: |
| Plastic compaction | Synthetic aggregates | increasing with decreasing $\phi$ if <br> disconnection occurs |
| Sintering |  | 4.5 for $\phi<0.10$ |
|  | Porous glass | disconnection at $\phi \approx 0.04$ |
| Semi-brittle compaction | Salt aggregates | $5-7$ |
| Elastic compaction | Sandstones | $1-25$ |
|  |  | depending on microstructure |
| Cataclastic compaction | Sandstones | $\approx 20$ |
| (hydrostatic) | $\phi>0.30$ |  |
|  | $0.15>\phi>0.30$ | $10-20$ |
| Cataclastic compaction | $\phi<0.15$ | $\approx 10$ |
| (triaxial) | Sandstones | $5-10$ |
|  | $\phi>0.30$ | $10-20$ |
| Dilatant microcracking | $0.15>\phi>0.30$ | $\approx 20$ |
|  | $\phi<0.15$ | $7-8$ |
| Thermal microcracking | Dense rocks | $\alpha$ decreasing with increasing $\phi$ |
|  | Dense rocks | $5-7$ |
| Dissolution |  | $\alpha \approx 1$ at very low $\phi$ |
| Precipitation | Sedimentary rocks | $>20$ |
| Chemical alteration | Sedimentary rocks | Porous glass |
| (roughening) | Sedimentary rocks | $\alpha$ decreasing with decreasing $\phi$ |
| Diagenesis |  | $\alpha \approx 2$ at $\phi<0.10$ |

Table 2.1: Chart listing compiled $\alpha$ values for multiple pore evolution processes obtained from experimental studies on a variety of natural and synthetic rock types. Chart taken from Bernabé et al. [2003].

### 2.1 Mechanical Compaction

If pressure is applied to a porous material, its granular structure and pore space will be compacted. This compaction will result in tighter grain packing, which will limit the amount of space available between grains (Figure 2.2). This reduction in pore space, in addition to the changes in pore geometry, will result in a permeability decrease within the material. The magnitude of the permeability
reduction will depend on many factors, including the initial pore geometry of the material and the magnitude of the pressure applied.

Previous studies identified two different regimes of mechanical compaction on porous sandstones [e.g., Zhang et al., 1990; Zhu and Wong., 1996]. When the applied pressure is relatively low, resultant changes in pore space are mostly reversible. Many studies analyzed suites of elastically deformed sandstones and found a large range of observed EPPRs with $\alpha$ values spanning roughly 1-21 [Bernabé et al., 1991; Fredrich et al., 1993; David et al., 1994]. However, when the pressure applied exceeds a given threshold, grain crushing and pore collapse occur, where resultant changes in pore space of previously pressure-insensitive pores are non-reversible (Figure 2.2c). The extent to which $k$ will change as a result of changes in $\phi$ depends highly on the initial pore geometry of the sample. Samples having a high initial $\phi$ show a steeper trend as they move into the brittle regime, whereas samples with a low initial $\phi$ experience a decrease in trend during this transition [Bernabé et al., 2003].

These correlations determined from porous sandstones may be applicable to vent deposits. Many vent deposits have structures consisting of interlocking crystals; however, some have a granular structure with pore space distributed somewhat similarly to that of sandstone. In deposits that exhibit this structure it is important to look for evidence of mechanical compaction. Larger isolated pores located at grain junctions may be indicative of elastic compaction, whereas crushed grains may provide evidence for brittle compaction.


Figure 2.2: Diagrams illustrating the effects mechanical compaction on pore structure. a) Initial pore structure prior to compaction. b) Uniform application of pressure (green arrows) on pore structure from a) results in tighter grain packing and a reduction in $k$ and $\phi$. Dashed box represents initial sample size. c) When the applied pressure exceeds the threshold pressure, non-reversible grain crushing and pore collapse occur further decreasing $k$ and $\phi$.

### 2.2 Hot Isostatic Pressing (HIP)

A porous material subjected to high temperatures and pressures will densify during the process of hot isostatic pressing (HIP) (Figure 2.3). As a result of the high temperature conditions, HIP is a plastic pore evolution process. Experiments on synthetic rock aggregates are conducted under both dry and wet conditions [e.g., Bernabé et al., 1982, 2003; Zhang et al., 1994; Wark and Watson, 1998], similar to environmental conditions during diagenesis.

Previous studies show that there is a critical porosity during HIP of calcite aggregates and carbonate rocks. In rocks with porosity greater than the critical porosity, the exponent $\alpha$ value is approximately 3 , whereas below the critical porosity higher $\alpha$ values are observed [Bernabé et al., 1982, 2003; Zhang et al., 1994]. Zhu et al. [1999] show that the interplay between pore space shrinking due to plastic deformation of grains and isolation of pores at grain junctions from the pinching off of connecting tubes is responsible for the change in EPPRs above and below the critical porosity.

Unlike in calcite, HIP of wet quartz does not produce major changes in EPPR, which suggests that pore disconnection through tube pinch off does not regularly occur within these rocks [Wark and Watson, 1998].

HIP experiments are good analogs to natural pore evolution processes and allow us to investigate the mechanisms that operative during plastic compaction and how they affect EPPRs. Although pressures and temperatures exerted on vent deposits are lower than those in most HIP experiments, seafloor hydrothermal deposits may experience deformation similar to that seen during HIP.


Figure 2.3: Diagrams illustrating the effects of hot isostatic pressing (HIP) on pore structure. a) Initial pore structure prior to HIP. b) Uniform application of pressure (green arrows) and heat on pore structure from a). Dashed box represents initial sample size. Pressure compacts the grains, reducing pore space, while the heat causes the grains to expand outward into pore space. Dotted lines show initial grain sizes and darker shaded areas around grains represent thermally expanded portions of the grains. Together the compaction and thermal expansion create a dense material with limited permeability and porosity.

### 2.3 Thermal Cracking

Stresses caused by large changes in temperature can produce isotropic microcracks throughout a porous material (Figure 2.4). Internal stresses are caused by the thermal expansion of grains, which can create grain size mismatches and anisotropy within a material [LeRavalec et al., 1996]. Therefore,
the magnitude of these stresses is controlled by the thermoelastic properties of the mineral grains within the material [deMartin et al., 2004]. The development of thermal cracks can potentially have a significant impact on the permeability of a material.


Figure 2.4: Diagrams illustrating the effects of thermal cracking on pore structure. a) Initial pore structure prior to heat application and cracking. b) Application of heat on pore structure from a) causing the grains to expand outward into pore space and develop isotropic cracks. Dotted lines show initial grain sizes with darker shaded areas around grains representing thermally expanded portions of the grains. Thermal cracks produced during grain expansion can potentially increase permeability despite pore space loss to grain expansion.

Studies investigating the effects of thermal cracking on permeability involve the controlled heating of rocks, in some instances via HIP, for extended periods of time followed by a period of controlled cooling. Depending on the techniques employed, the nature of thermal cracking varies within samples, which leads to some differences in how the cracks influence permeability. In mylonite samples that were exposed to a simple thermal treatment (no HIP), increasing temperatures resulted in increased porosity in the form of well-connected cracks [LeRavalec et al., 1996]. The increased connectivity in these samples yielded higher permeability values; however, these permeabilities are sensitive to changes
in pressure. A small increase in confining pressure on the samples would close the microcracks and decrease the permeability [LeRavalec et al., 1996].

With experiments using HIP, samples are heated and pressurized to create a relatively impermeable sample with few cracks. A study with olivine aggregates used the HIP technique and then evaluated crack growth and permeability development during cooling and depressurization [deMartin et al., 2004]. This study found that longer HIP durations inhibited the development of large crack networks, because pressing reduced the number of potential initiation sites. Also, larger grains were more prone to developing cracks because of increased stresses [deMartin et al., 2004]. As a result porosity would somewhat increase in these samples due to crack formation without significantly enhancing permeability. On average, Bernabé et al. [2003] notes $\alpha$ values ranging from 5-7 for samples having experienced thermal cracking (Table 2.1).

The contrasting temperatures of hot hydrothermal fluids and cold seawater present at seafloor vent fields may make vent deposits susceptible to pore structure changes resulting from thermal cracking. Cracking, if present, may increase deposit permeability and potentially channel fluids moving through the deposits. Most vent deposits are initially quite porous, so the effects of thermal cracking may be expected to be more similar to those obtained without HIP.

### 2.4 Precipitation

Although the processes discussed thus far have been physical processes, chemical processes can also produce large changes in permeability. Precipitation occurs as fluids saturated in various ions pass through a material and react to
produce solid mineral phases. Because the fluids must pass through pore space, precipitates formed during the chemical reactions will be deposited within the pore space. This precipitation can lead to a narrowing of channels and pores, thereby decreasing permeability and restricting the overall flow of fluid.

A good example on how precipitation affects EPPRs is Fontainebleau sandstone from Paris basin. This well-sorted sandstone is monomineralic (99.9\% quartz) with a wide range of porosity due to different degrees of cementation resulted from groundwater oversaturated with quartz. Bourbie and Zinszner [1985] conducted permeability measurements on a suite of Fontainebleau sandstone samples with initial porosity from 3-28\%. They found that for samples with porosity greater than $7 \%$, the exponent $\alpha$ value is approximately 3 , whereas below the 7\%, $\alpha$ value increases to $\sim 7$ (Table 2.1). Zhu et al. [1995] show that the increase in the $\alpha$ value can be explained by pore connectivity loss in low porosity samples.

Experiments transmitting fluids through sandstones [Reis and Acock, 1994; Todd and Yuan, 1992] have shown that mineral precipitation, as expected, produces a loss of both permeability and porosity. These studies found that precipitation within pore space produces roughness along pore walls (Figure 2.5). Roughness creates irregularities along pore walls with small pockets that cannot effectively transmit flow [Bernabé et al., 2003]. Although the precipitated roughness may contain pores, lessening the net loss of porosity, permeability is greatly restricted by the roughness, particularly because the pore wall topography will narrow or even pinch off pore channels limited connectivity. Because of this
large decrease in permeability with respect to minor reductions in porosity, samples experiencing precipitation will have high $\alpha$ values, averaging about 8 (Table 2.1) [Bernabé et al., 2003].

Seafloor vent deposits regularly transmit fluids rich in a range of ions, many of which begin to precipitate upon emission to the seafloor. It has been previously shown [Zhu et al., 2007] that precipitation can dramatically impact the permeability and porosity of vent deposits, such as relict spires. High degrees of precipitation with spires lead to a loss of connectivity and an inability to continue transmitting fluids. Therefore, it is important to identify occurrences of late-stage precipitation within the deposits included in this study.


Figure 2.5: Diagrams illustrating the effects of precipitation on pore structure. a) Initial pore structure prior to precipitation. b) Late-stage mineral precipitation (green areas) within the pore structure of a) creating roughness along pore walls. Roughness narrows pore channels and in some cases pinches off channels. These effects decrease the porosity and restrict the permeability of the material.

### 2.5 Dissolution

Much like precipitation, dissolution is a chemical process that can produce large changes in a material's pore structure. Dissolution occurs when the chemical components of a fluid react with and breakdown the surrounding grain/crystal host structure. The dissolved components of the material become entrained in the
fluid and get carried away or potentially re-precipitated as the fluid passes through the host structure. As a result, mineral dissolution increases the porosity of a material, while typically also enhancing the permeability.

The extent of dissolution a material experiences depends highly on the chemical composition of the fluid, particularly its acidity, and the composition of the host material. For example, studies have been conducted on sandstones and carbonate rocks whereby acidic fluid is passed through the rocks and changes in $k$ and $\phi$ are then observed [McCune et al., 1979; Luquot and Gouze, 2009]. Dissolution was shown to enhance permeability in both cases; however, the increase in permeability in the carbonate rocks was several orders of magnitude higher than the increase seen in the sandstones. The dissolution along grain edges helps create wider and more connected flow channels, particularly in areas of the pore network that were initially well-conducting pores (Figure 2.6) [Bernabé et al., 2003]. The differences in the extent of dissolution between the two rock types is confirmed by their respective $\alpha$ values, with the sandstones have $\alpha$ values from 8-10 and the carbonate rocks having much higher $\alpha$ values $\sim 20$ (Table 2.1) [Bernabé et al., 2003].

The majority of analyzed hydrothermal fluids have been found to be acidic, making dissolution reactions common during fluid migration. Certain vent deposits that contain higher abundances of carbonate minerals may be more susceptible to dissolution. Given the large impact dissolution can have on the pore structure and permeability of any material, it is important to identify whether dissolution has occurred within the vent deposit samples.


Figure 2.6: Diagrams illustrating the effects of dissolution on pore structure. a) Initial pore structure prior to dissolution. b) Fluid reacts with and dissolves grains, removing material (green regions) and decreasing grain size, thereby increasing pore space and significantly enhancing permeability.

## Chapter 3: Experimental Methods

### 3.1 Probe permeability

Vent deposits are heterogeneous. The permeability of vent deposits is generally heterogeneous and anisotropic. To assess the heterogeneity and anisotropy of the various deposits, permeability measurements were taken at multiple sites from multiple facets of each deposit sample using a probe permeameter. Variations in probe permeability values provide a measure of sample heterogeneity [Zhu et al., 2007].

Permeability measurements on vent deposits were conducted by using a portable probe permeameter, the NER TinyPerm $\mathrm{II}^{\mathrm{TM}}$ (Figure 3.1a). The probe permeameter measurements provide a quantitative measurement of permeability heterogeneity within the samples. Comparison of permeability values obtained from different facets of a vent sample can be used to identify existing anisotropy.

The probe permeameter is a syringe-like device that pulls air out of a sample through a compressible rubber tip (Figure 3.1b). The compressible tip is held firmly against the sample surface to reduce the possibility of leaked air (Figure 3.1c). A micro-controller unit within the permeameter monitors the syringe volume as air is pulled from the sample. While air is pumped from the vicinity of the sample surface, transient vacuum pulses are created. Air is pulled from the near-surface region around the measurement site, where it is not directionally restricted. A signal processing algorithm determines the permeability using the vacuum pulses and syringe volume.

To assess the permeability anisotropy of each sample, permeability values along different orientations were compared by conducting measurements on multiple facets of the sample. Along each facet, permeability was measured at several different sites. At each measurement site, five permeability measurements were collected (Figure 3.1c). Permeability values provided for each site are given by the geometrical mean of the 5 measurements.


Figure 3.1: a) NER TinyPerm II ${ }^{\mathrm{TM}}$ probe permeameter. b) Operation of a probe permeameter-the probe tip seal against a sample surface and air being pumped from the sample. c) A probe permeameter in usemultiple measurement sites along the sample surface are shown by the black labels. At each site, 5 measurements were conducted.

### 3.2 Core permeability

Because the flow pattern around the probe tip is unknown (Figure 3.1b), the probe permeability are not unidirectional. Furthermore, there are no porosity data available for the vent samples. To quantify the permeability-porosity relationships in vent deposits, right-cylindrical cores were then taken near the selected probe permeameter measurement sites so that core permeability data can be compared to the probe permeability data obtained. The cores are 2.54 cm (1 inch) in diameter and their lengths vary from $\sim 1-6 \mathrm{~cm}$.

Using these cylindrical cores, permeability was measured along the axial direction with a nitrogen permeameter, the UltraPerm ${ }^{\mathrm{TM}} 400$. The permeameter
uses Darcy's law, whereby fluid flow is proportional to a differential pressure over a given length, to determine sample permeability. This relationship can be rearranged to solve for permeability, assuming a number of flow properties are known (Equation 3.1). This expression requires a well constrained flow of nitrogen through the sample that can then be used in conjunction with recorded upstream and downstream pressures to determine the permeability of the core. The sample length $(L)$ and cross-sectional area $(A)$ are input for each core, while the nitrogen flow rate $(Q)$, upstream pressure $\left(P_{1}\right)$, downstream pressure $\left(P_{2}\right)$, and flow viscosity $(\mu)$ are measured by the permeameter. Since the permeameter uses Darcy's law, it is critical that the flow rate is low ( $\sim 3 \mathrm{~cm}^{3} / \mathrm{min}$ ), because Darcy's law only holds true for linear laminar fluid flow.

$$
\begin{equation*}
k=\frac{Q L \mu}{\left(P_{1}-P_{2}\right) A} \tag{3.1}
\end{equation*}
$$

Unlike with the probe permeameter, fluid flow through the sample is restricted along the axial direction by the application of impermeable jackets against the cylindrical surface. To circumvent problems caused by surface roughness, such as large voids positioned along the sample surface, plastic wrap was also applied around the sample. To prevent by-flow between the sample and the impermeable jacket, higher confining pressure was applied. Additionally, to account for gas slippage within pore space that can occur during measurement, a Klinkenberg correction has been applied to the permeability data.

Higher confinement could cause mechanical compaction of the sample and consequently alter sample porosity and permeability. To address this problem, permeability measurements were taken at multiple confining pressures, increasing
incrementally from $0.5-2.7 \mathrm{MPa}$ ( $\sim 70-400 \mathrm{psi}$ ). After completing measurements taken at each incremental confining pressure (loading cycle), measurements were then taken with confining pressure decreasing from the peak pressure (unloading cycle) to 0.5 MPa . The permeability data obtained during the loading and unloading cycles can be used to gauge the effect of confining pressure on the core's permeability. Five permeability measurements were made at each confining pressure. Geometrical mean and the standard deviation of these values were determined.


Figure 3.2: a) UltraPerm ${ }^{\mathrm{TM}} 400$ nitrogen permeameter (bottom) and UltraPore ${ }^{\mathrm{TM}} 300$ helium porosimeter (top). b) Pressure chamber used with the nitrogen permeameter and helium porosimeter. 2.54 cm diameter cores were measured in the left chamber.

### 3.3 Porosity

Porosity measurements have been made on the sample cores using a helium porosimeter, the UltraPore ${ }^{\mathrm{TM}} 300$. The operation principal of the porosimeter is based on Boyle's law (Equation 3.2), where pressure and volume are inversely proportional at a constant temperature.

To obtain the porosity, it is first necessary to measure the sample pore volume. Measuring the pore volume requires the porosimeter to use a reference volume. In this case, a known volume of helium is held in a reference cell under a known pressure of 1.4 MPa . These values will serve as the initial conditions $V_{l}$ and $P_{1}$, respectively. As measurement begins, helium from the reference tank is directed into the sample core resulting in a pressure drop within the reference tank. Once the helium fills the sample's available pore space, the decreased pressure in the reference tank is recorded as $P_{2}$. The remaining volume of helium within the tank is recorded as the dead volume $\left(V_{\text {dead }}\right)$, with $V_{2}$ equaling the sum of the sample's pore volume $\left(V_{\text {pore }}\right)$ and $V_{\text {dead }}$ (Equation 3.3). Using these values, $V_{\text {pore }}$ can be solved for (Equation 3.4).

$$
\begin{gather*}
\frac{P_{1} V_{1}}{T_{1}}=\frac{P_{2} V_{2}}{T_{2}} \quad \text { where } T_{1}=T_{2}  \tag{3.2}\\
P_{1} V_{1}=P_{2}\left(V_{\text {dead }}+V_{\text {pore }}\right)  \tag{3.3}\\
V_{\text {pore }}=\frac{P_{1}}{P_{2}}\left(V_{1}-V_{\text {dead }}\right) \tag{3.4}
\end{gather*}
$$

To determine the sample porosity, the core's bulk volume $\left(V_{b u l k}\right)$ is needed, which can be found using measured values of the core's length $(l)$ and diameter (d) (Equation 3.5). The porosity is then calculated by dividing the pore volume by the bulk volume (Equation 3.6).

$$
\begin{align*}
& V_{\text {bulk }}=\left(\frac{d}{2}\right)^{2} l \pi  \tag{3.5}\\
& \phi=\frac{V_{\text {pore }}}{V_{\text {bulk }}} \times 100 \tag{3.6}
\end{align*}
$$

The procedure for sample preparation is similar to that described for permeability measurements. Because using helium makes the experimental setup more susceptible to by-flow, higher confinements are applied for porosity measurements. Five pore volume measurements are made at confining pressures of 2.1-3.1 MPa. Cores having a length $\leq 2.54 \mathrm{~cm}$ were measured with a steel core (with a known volume of $14.8 \mathrm{~cm}^{3}$ ) so that the measurement is conducted on a more substantial bulk volume that yielded more accurate pore volume measurements. In these cases, the total calculated pore volume for the combined cores was accepted as the pore volume for the vent core, assuming negligible pore volume for the steel core. The porosities were then determined using this calculated pore volume and the bulk volume of the sample core.


Figure 3.3: Experimental setup for porosity measurements using a helium porosimeter. a) A core sample with plastic wrap along its side is placed within the pressure chamber. b) Core shorter than 2.54 cm together with a steel core within the pressure chamber. The steel core helps to create a better seal around the sample thus increasing the accuracy of the porosity measurements. In both cases a confining pressure is applied along the radial directions, while helium is flowing through the core along the axial direction.

Permeability-porosity relationships of vent deposits can be obtained. Compiling the porosity and permeability values collected on vent samples at different formation stages provides the evolution of permeability and porosity of seafloor hydrothermal vents.

### 3.4 Petrography

After completing the permeability and porosity measurements, thin sections were made from a large subset of the cores. Thin sections were $30 \mu \mathrm{~m}$ thick and were impregnated with epoxy. Sections were cut both radially (Figure 3.4a) and axially (Figure 3.4b) through the cores. Microstructural analyses using reflected and refracted light petrography were then conducted on thin sections to identify pore evolution processes. These analyses are necessary in order to interpret the EPPRs observed for the different samples. Evidence for processes, such as mineral dissolution or precipitation, were recorded in addition to other sample characteristics, including grain size, sorting, and packing. The overall mineralogy of the samples was also noted. While these sections only represent a 2-D cross-section through the sample, in many cases, different orientations of sections from the same sample could be compared and used to make inferences regarding the 3-D structure of the sample (Figure 3.4c).


Figure 3.4: Diagrams showing orientation of thin sections cut from cores and their relation to the 3-D structure of the sample. a) A radial cut through a core. b) An axial cut through a core. c) How radial and axial thin sections for the same sample can be used to make inferences about the 3-D structure.

## Chapter 4: Massive Anhydrite

Measured anhydrite samples were recovered from three different vent fields: the Fenway vent field within the Manus Basin, the Edmond vent field along the Central Indian Ridge (CIR), and the Trans-Atlantic Geotraverse (TAG) along the Mid-Atlantic Ridge. For this study, there were a total of 8 samples measured: 3 from Fenway, 1 from CIR, and 4 from TAG.

### 4.1 Permeability and Porosity

Probe permeability measurements were made along the surface of each of the samples; these data are plotted in Figure 4.1. The permeability data ranged from $\sim 1 \times 10^{-14}$ to $6 \times 10^{-12} \mathrm{~m}^{2}$. These data represent permeability measurements made along the various sides of the samples, and no systematic difference in permeability values was observed. This suggests that there is not significant permeability anisotropy within these samples. In a few samples, such as J2-216-5-R1, the degree of the surface roughness varies considerably (Figure 4.2). For these samples, permeability


Figure 4.1: Histograms showing probe permeability data for massive anhydrite samples. Plots show frequency of measurements at a given permeability. Colors indicate sample locations: Fenway $=$ purple, CIR $=$ red, and TAG $=$ green . a) Data for three Fenway samples. b) Data for one CIR sample. c) Data for four TAG samples.
values obtained along the cut surfaces (Figure 4.2) are generally smaller than the ones obtained along the rough surfaces. Greater permeability values are likely consequences of the imperfect seal between the probe tip and the rough surfaces, thus does not represent actual anisotropy.


Figure 4.2: Anhydrite sample J2-216-5-R1 with one cut side (highlighted in green). A cut surface has significantly reduced surface roughness. Measurements along this surface would have lower permeability values, since the probe permeameter is able to form a tighter seal against the sample.

Cylindrical cores were taken from selected sites where probe permeability measurements were taken. Permeability and porosity along axial direction of a total of 17 anhydrite cores were measured using the nitrogen permeameter and helium porosimeter, respectively.

During permeability measurements, cores were subjected to different confining pressures of $\sim 0.5-2.7 \mathrm{MPa}$. The incremental application and removal of confining pressure allowed us to gauge the effects of pressure on anhydrite permeability. Permeability data of all the cores is compiled in Table 4.1. As confining pressure increases, a slight reduction in permeability is observed (Figure 4.3). This permeability reduction is more pronounced at relatively low confining pressures, which is generally attributed to crack closure [Walsh, 1965]. Data during the removal of confining pressure (unloading) show that this crack




Figure 4.3: Plots showing pressure profiles for select samples during permeability measurement. Colors indicate sample location: Fenway $=$ purple, $\mathrm{CIR}=$ red and $\mathrm{TAG}=$ green. Symbol shapes represent different sample cores. Solid symbols for pressure loading and empty symbols for unloading cycles. Generally minor changes in permeability with pressure. a) Pressure profiles for Fenway sample J2-216-14-R1 cores. b) Pressure profiles for CIR sample J301-3 cores. c) Pressure profiles for TAG sample MIR-1 2/78, Sta 2417 cores.
closure is predominantly reversible and therefore has little effect on the permeability of the cores.

Permeability values for the cores range from $\sim 10^{-16}$ $6 \times 10^{-13} \mathrm{~m}^{2}$, while porosity values range from 2-15\%. Permeability vs. porosity obtained at a confining pressure of 2.1 MPa is plotted in Figure 4.4. The EPPR best fitting the data has a trend with an $\alpha$ value of 4 . From this plot, it is clear that the data fall into roughly two groups: one with high $k$ and $\phi$ values and another with low $k$ and $\phi$. To understand the differences between the two groups and the variation within the data, it is necessary to identify the elements of the pore structure and evidence for pore evolution processes controlling $k$ and $\phi$.

| Sample | Core | $\begin{gathered} \text { Probe } k \\ \left(\times \mathbf{1 0}^{-15} \mathbf{m}^{2}\right) \end{gathered}$ | Core $k\left(\times 10^{-15} \mathrm{~m}^{2}\right)$ |  |  | Core $\phi$ (\%) |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $\phi \quad 2 \sigma$ | $\phi \quad 2 \sigma$ |
| Fenway: |  |  |  |  |  |  |  |
| J2-210-8-R2 | A-1 | $39.0 \pm 0.4$ | $0.3 \pm 0.02$ | $0.3 \pm 0.01$ | $0.3 \pm 0.01$ | $2.3 \pm 1.1$ | $2.7 \pm 1.7$ |
|  | A-2 | -- | $0.1 \pm 0.01$ | $0.1 \pm 0.01$ | $0.1 \pm 0.01$ | $1.9 \pm 2.7$ | $2.0 \pm 2.2$ |
|  | A-3 | -- | $0.2 \pm 0.01$ | $0.1 \pm 0.01$ | $0.1 \pm 0.01$ | $2.8 \pm 4.6$ | $2.1 \pm 4.6$ |
| J2-216-5-R1 | A2 | $402.8 \pm 1.9$ | $237.9 \pm 0.7$ | $230.4 \pm 0.5$ | $225.2 \pm 0.2$ | $11.5 \pm 4.8$ | $11.8 \pm 4.0$ |
| J2-216-14-R1 | A | $2268.4 \pm 2.9$ | $0.3 \pm 0.02$ | $0.3 \pm 0.01$ | $0.3 \pm 0.01$ | $2.1 \pm 2.2$ | $3.1 \pm 1.1$ |
|  | B | $2144.6 \pm 3.2$ | $0.4 \pm 0.04$ | $0.3 \pm 0.01$ | $0.3 \pm 0.01$ | $5.6 \pm 0.1$ | $5.5 \pm 0.1$ |
| CIR: |  |  |  |  |  |  |  |
| J301-3 | A | $1938.5 \pm 1.1$ | $39.8 \pm 0.3$ | $38.5 \pm 0.4$ | $38.0 \pm 0.3$ | $14.8 \pm 0.1$ | $14.7 \pm 0.1$ |
|  | B | $1694.3 \pm 1.3$ | $1.1 \pm 0.04$ | $0.7 \pm 0.04$ | $0.6 \pm 0.05$ | $7.9 \pm 0.1$ | $7.8 \pm 0.1$ |
| TAG: |  |  |  |  |  |  |  |
| ALV 2581-8 | A | $1713.4 \pm 1.6$ | $357.8 \pm 2.0$ | $352.6 \pm 2.1$ | $349.2 \pm 1.8$ | $12.7 \pm 0.1$ | $12.6 \pm 0.1$ |
| MIR 1 1/74 | A-1 | $29.0 \pm 0.4$ | $125.3 \pm 0.7$ | $123.2 \pm 0.9$ | $121.4 \pm 0.4$ | $10.9 \pm 0.2$ | $10.8 \pm 0.1$ |
| Sta 2403 | A-2 | -- | $340.4 \pm 2.2$ | $336.7 \pm 1.9$ | $335.1 \pm 2.1$ | $12.2 \pm 0.2$ | $12.0 \pm 0.2$ |
|  | B | $509.8 \pm 2.2$ | $79.3 \pm 0.6$ | $77.1 \pm 0.7$ | $75.7 \pm 0.3$ | $9.0 \pm 0.1$ | $8.9 \pm 0.1$ |
| MIR 1 2/78 | A | $106.5 \pm 3.0$ | $497.3 \pm 1.2$ | $477.9 \pm 1.0$ | $473.0 \pm 1.1$ | $15.0 \pm 0.1$ | $14.8 \pm 0.1$ |
| Sta 2417 | B | $1043.7 \pm 3.9$ | $607.4 \pm 1.5$ | $595.9 \pm 0.7$ | $590.9 \pm 0.4$ | $13.5 \pm 0.2$ | $13.3 \pm 0.1$ |
| ALV 21837-0 | 2 | -- | $1.0 \pm$--- | $0.4 \pm$--- | $0.1 \pm$--- | $5.0 \pm$--- | $3.6 \pm$--- |
|  | 3 | -- | $0.4 \pm$--- | $0.4 \pm$--- | $0.3 \pm$--- | $6.1 \pm$--- | $2.4 \pm$--- |
|  | B | $674.9 \pm 3.9$ | $0.6 \pm$--- | $0.3 \pm$--- | $0.2 \pm--$ | $6.6 \pm$--- | $4.0 \pm$--- |

Table 4.1: Permeability and porosity data for anhydrite cores.


Figure 4.4: Permeability and porosity data plotted for each of the anhydrite cores. Symbol colors consistent with Figure 4.3. Dashed line indicates trend of EPPR with $\alpha \sim 4$. Data are divided into roughly two groups high $k$ and $\phi$ and low $k$ and $\phi$.

### 4.2 Microstructural Analyses

Twelve thin sections were made from a subset of the measured anhydrite cores. Permeability and porosity data of cores from which thin sections were taken are plotted in Figure 4.5. Thin section descriptions have been organized into two groups, high $k$ and $\phi$ and low $k$ and $\phi$, consistent with the permeability and porosity data.


Figure 4.5: Permeability and porosity data plotted for each of the anhydrite cores from which thin sections were made. Symbol colors consistent with Figure 4.3. Symbol shape represents sample and symbol labels indicate respective core. Data are divided into roughly two groups: high $k$ and $\phi$ and low $k$ and $\phi$.

### 4.2.1 High $k$ and $\phi$ cores

There were six thin sections taken from the high $k$ and $\phi$ cores: MIR 1 2/78, Sta 2417 (Cores A and B), ALV 2581-8 (Core A), MIR 1 1/74, Sta 2403 (Cores A-1 and B), and J301-3 (Core A). Each section predominantly contains anhydrite with minor amounts of sulfides (i.e. chalcopyrite, pyrite, and
sphalerite). Anhydrite crystals are typically tabular with well defined cleavage planes. They initially form in a tightly interlocked grain structure with little space between crystals. All of the thin sections in this group have, however, experienced dissolution, which has created pore space and channels between crystals (Figures 2.6 and 4.6). Dissolution has weathered crystal edges causing many of the crystals to lose their tabular shape and also develop cracks. Pore spaces are irregularly shaped and well-connected likely due to dissolution. This can account for the higher permeability and porosity values measured for these samples.

The thin sections for the cores (A and B) from the MIR 1 2/78, Sta 2417 sample are very similar. They both show blocky anhydrite crystals, ranging in size from about 40-1200 $\mu \mathrm{m}$, with smaller crystals generally appearing to be broken portions of larger crystals (Figure 4.6a). Most crystals adjacent to larger pores and channels have rough edges indicative of crystal breakdown via dissolution. Additionally, dissolution has produced narrow cracks along anhydrite cleavage planes, which in some cases appears to have lead to the division and eventual breakdown of large crystals into several smaller crystals. Very small amounts of chalcopyrite have begun to precipitate within some of the newly created pore space, though these crystals do not sufficiently block any pore space between the anhydrite crystals.

The thin section for ALV 2581-8, Core A is also very similar to those of MIR 1 2/78, Sta 2417. The section contains anhydrite crystals that have been dissolved along crystal edges and internally along cleavages. The dissolution has also created broad channels between crystals, in some cases as wide as $100 \mu \mathrm{~m}$
(Figure 4.6b). This section does exhibit a greater abundance of chalcopyrite that has precipitated within the pore space. The chalcopyrite crystals are typically small $(10-40 \mu \mathrm{~m})$ and do not appear to effectively block much of the pore space.


Figure 4.6: Reflected light (5x) images of thin sections from high $k$ and $\phi$ cores. Width of images is approximately 2.7 mm . a) MIR $12 / 78$, Sta 2417, core A. Large anh grains with dissolved edges, minor amounts of cp. b) ALV 2581-8, core A. Broad channels between anh grains, some cp and py. c) MIR $11 / 74$, Sta 2403, core A-1. Broad pore space around grains, small clusters of cp and py. d) MIR 1 1/74, Sta 2403, core B. Anh grains with dissolved edges. Multiple cp/py clusters in pore space. e) J301-3, core A. Highly fragmented anh grains with a few py grains surrounded by smaller grain fragments.

Thin sections for cores A-1 and B of sample MIR 1 1/74, Sta 2403 are consistent with those discussed above. Both cores have clearly experienced a significant about of dissolution, which has created significant pore space within the initially close-packed anhydrite structure (Figures 4.6c,d). These sections
differ slightly from those previously discussed in that they both show precipitation of chalcopyrite and some pyrite, which appear to cluster within larger pores. While this precipitation does not fully block most of the pore channels between crystals, it is clearly restrictive. The greater proportion of sulfides within these two cores can likely account for the somewhat lower permeability and porosity values measured compared to those of the previous cores discussed.

Lastly, core A from sample J301-3 also plots within the high $k$ and $\phi$ group of samples, though it plots slightly off from the EPPR trend. The thin section for this core is structurally quite different from those of the previously discussed cores (Figure 4.6e). There is evidence for extensive anhydrite dissolution, perhaps multiple dissolution events, that have severely cracked and broken down many of the crystals (Figure 4.7a). Large anhydrite crystals that have remained intact are heavily weathered. The dissolution has created more pore space around and within crystals, though the development of broad channels as seen in the previous


Figure 4.7: Cross-polarized images (5x) of sample J301-3. Sample has experienced severe dissolution, which has cracked and broken anh grains. Pore space has been created within and around grains. Width of images is $\sim 2.7 \mathrm{~mm}$. a) Core A. b) Core B. thin sections appears inhibited by the packing of many of the small, anhydrite
crystal fragments. The tight packing of these anhydrite fragments likely somewhat limits the permeability of the core, which is why this core plots below the high $k /$ high $\phi$ trend. Late-stage sulfide precipitation is less pronounced within this thin section. A few pyrite crystals have developed within the section and are surrounded by anhydrite fragments. It is possible that the pyrite precipitated in pore space created by early dissolution, and that further precipitation was impeded by the infill of anhydrite fragments within the pore space during subsequent dissolution events.

### 4.2.2 Low $k$ and $\phi$ cores

Six thin sections were made from cores that plotted in the low $k /$ low $\phi$ group: J301-3 (Core B), ALV 21837-0 (Core B), J2-216-14-R1 (Cores A and B) and J2-210-8-R2 (Core A-1). Similar to the previous cores, these all show evidence for having experienced anhydrite dissolution. Anhydrite edges are weathered and rough, while the crystals themselves are heavily cracked. These cores, however, appear to have undergone secondary processes that reduced the pore space created by dissolution.

Much like sample J301-3 core A, J301-3 core B exhibits evidence for extensive amounts of anhydrite dissolution. Anhydrite crystals have been thoroughly cracked and fragmented with few large crystals remaining (Figure 4.7b). The majority of the anhydrite is present as small crystal fragments that have been tightly packed. The close packing of these crystals has clearly limited the availability of pore space between the crystals, more so than in core A. Core B also contains more pyrite crystals than core A , which probably helped fill pore
spaces created during an early dissolution event (Figure 4.8a). These pyrite crystals have also been tightly packed with the anhydrite fragments, suggesting that they precipitated prior to some of the more intense dissolution events. The


Figure 4.8: Reflected light (5x) images of thin sections from low $k$ and $\phi$ cores. Width of images is approximately 2.7 mm . a) J301-3, core B. Tightly packed anh grain fragments surrounding py grains. b) ALV 21837-0, core B. Vugs created from dissolution have been infilled by newly precipitated smaller anh grains. c) J2-216-14-R1, core B. Precipitation of sulfides and anh fill in pore space. d) J2-216-14-R1, core A. Precipitation of sulfides and anh fill pore space created during dissolution. Precipitated anh grains are large and add significant surface roughness along pore walls. e) J2-216-14R1, core A. Fragmented anh grain surrounded by cp. Cp precipitated later and will prevent further dissolution of the anh grain. f) J2-210-18-R2, core A-1. Dissolution to a lesser extent produces narrow channels between grains that can be easily restricted or blocked by sulfide precipitation.
tightly packed structure and greater abundance of pyrite within core B likely accounts for the variation observed between the two J301-3 cores.

Core B from sample ALV 21837-0 has also experienced considerable dissolution which has created broad channels between anhydrite crystals and cracks within crystals. Some crystals have been fragmented following severe cracking, though not nearly to the extent as seen in sample J301-3. Small amounts of both chalcopyrite and pyrite have precipitated in the pore space without creating much restriction. These sulfide crystals are dispersed widely and seldom form clusters. This core does, however, contain large patches of more recently precipitated interlocking anhydrite crystals. These crystals are well-formed, but much smaller (5-30 $\mu \mathrm{m}$ ) than the original crystals. They appear to have precipitated within vugs (large pore spaces), where they protrude outwards into remaining pore space, creating significant roughness along the pore channels (Figure 4.8 b). Permeability and porosity enhancements resulting from anhydrite dissolution are reduced due to this late-stage precipitation of smaller anhydrite crystals.

Sample J2-216-14-R1 has three thin sections, two for core B and one for core A. The core B thin sections show that pore space created by dissolution has been restricted by the precipitation of sulfides and small anhydrite crystals (Figure 4.8c). These new crystals block pore space between crystals and created added roughness along channel walls, which decrease the permeability of the core. Core A shows evidence for similar precipitation of sulfides and newer anhydrite. The precipitated anhydrite crystals in core A are well-formed and larger ( $30-400 \mu \mathrm{~m}$ )
than those seen in core B (Figure 4.8d). These crystals extend into pore spaces, often blocking flow channels. Sulfide precipitation is also more pronounced in core A with chalcopyrite and pyrite filling pore space around crystal edges (Figure 4.8e).

Lastly, core A-1 from sample J2-210-8-R2 also plots in the low $k /$ low $\phi$ group. While this core shows signs of dissolution, the extent of dissolution is less than in the other samples. Pore space has been created around crystals, but is narrower than as seen in other sample thin sections (Figure 4.8f). Small amounts of chalcopyrite and pyrite have precipitated within some of the pore space, and in most cases because the pore space is narrow, it blocks flow channels. The lesser amounts of initial anhydrite dissolution in conjunction with the sulfide precipitation restrict the permeability of the core.

### 4.3 Discussion

Seafloor massive anhydrite deposits form by the mixing of hot fluids with seawater, and likely behave similarly to subsurface deposits. Although anhydrite precipitation is recognized as being a key constraint on the flow of hydrothermal fluids, effects of anhydrite precipitation on transport properties of the seafloor and its subsurface structures are not well quantified [Mills and Tivey, 1999; Lowell et al., 2003].

Microstructural analyses show that the pore geometry of anhydrite deposits is controlled by both dissolution and precipitation. As fluids pass through the deposits, environmental conditions such as temperature and pH are altered, which may cause anhydrite deposits to become unstable and dissolution to
occur. In general, dissolution creates pore space and enhances permeability. The anhydrite samples investigated in this study all show evidence for dissolution.

Precipitation is sensitive to environmental conditions and pore size [Aquilano et al., 1992; Pape et al., 2005]. In a study of anhydrite cementation in sandstones, Pape et al. [2005] found that anhydrite preferentially precipitates in larger pore spaces, such as in wide flow channels, as opposed to smaller spaces that could be more easily infilled. It is observed that anhydrite precipitation in smaller pores quickly becomes unstable with drops in the calcium sulfate concentration [Pape et al., 2005]. Anhydrite crystal growth is shown to begin with the nucleation of a single crystal within a large pore, not through growth along the edges of pre-existing crystals as is common with other minerals such as quartz. Thus anhydrite precipitation creates highly efficient blockages that impede fluid transmission, and as such results in large permeability reduction.

This mode of crystal growth may explain the difference between the two observed sample groups. In high $k$ and $\phi$ samples, pore space consists primarily of channels along crystal edges. In low $k$ and $\phi$ samples, abundant vugs are observed. The vugs are ideal for the growth of secondary anhydrite that results in low permeability. In contrast, channels are less conducive to anhydrite growth.

The precipitation of sulfides, such as chalcopyrite and pyrite, is also observed within many of the anhydrite samples. Precipitation of such minerals requires higher temperatures [Fontaine et al., 2001]. Sulfides preferentially precipitate within constrictive spaces making them more likely to restrict pore spaces and decrease permeability [Fontaine et al., 2001].

The effects of both dissolution and precipitation of secondary anhydrite or sulfides within the anhydrite deposits depend highly on their initial crystal structures and environmental conditions. Vugs within the anhydrite pore structure are not necessarily caused by dissolution and can be original to the grain structure. Vug-rich anhydrite deposits are predisposed to secondary anhydrite precipitation. Additionally, late-stage precipitation can occur at any time, regardless of the timing of dissolution events. Precipitation can occur within pore space original to the grain structure, as well as in pores created through dissolution. The extent and timing of these processes is controlled by the chemistry of the vent fluids.

The $\alpha \sim 4$ trend identified for the anhydrite deposits reflects this interplay between dissolution and precipitation. As previously discussed, dissolution, a permeability and porosity enhancing process, has been shown to produce $\alpha$ values as high as 20 [Bernabé et al., 2003]. Precipitation, which results in lower permeability and porosity, generally has $\alpha$ values around 8 [Bernabé et al., 2003]. Because both of these processes occur in anhydrite deposits, the $\alpha$ value will reflect a balance of their effects on pore structure (Figure 4.9). By selecting a starting $k$ and $\phi$ and then imposing either process, the pore structure of the material will change. If this altered composition then experiences the opposing process, the trend from the starting structure to the final structure will be similar to that of $\alpha \sim 4$.


Figure 4.9: $\alpha \sim 4$ trend reflects interplay between dissolution and precipitation.

## Chapter 5: Flanges, Slabs, and Crust

Flanges, slabs and crust are diffusive vent deposits that facilitate fluid transport laterally as well as upwards through deposit cracks, which results in a layered structure (Figure 5.1). This structural characteristic indicates that the processes controlling the formation of their layers, such as mineral precipitation or dissolution, may be similar in these vents [Delaney et al., 1992; Tivey et al., 1995; Cooper et al., 2000]. Ten deposit samples were included in this group: 5 flanges, 4 slabs, and 1 crust. Prior to discussing the results, it is first necessary to understand variations between the deposits, such as how and where they develop.

5.1: Two sets of cores illustrating the prominent structural layering throughout the samples. a) Slab sample ALV 2608-4-1, Pc 1. Left core was taken perpendicular to layering and right core was taken parallel to layering. b) Flange sample ALV 2415-1B. Top core was cut perpendicular to layering, while the bottom core was taken parallel to layering.

### 5.1 Geologic Descriptions

The most widespread of these sample types are the flanges (Figure 5.2a), which have been identified at multiple vent sites. Flanges are deposits that extrude horizontally from the sides of larger chimney structures, like a ledge or tier, where fluid will pool and form layers within the flange parallel to the surface of the fluid [Delaney et al., 1992; Woods and Delaney, 1992]. Fluids pooled beneath a flange
will percolate upwards through the flange, or in the case where the fluid flux from the adjacent chimney structure exceeds the maximum amount able to percolate, the fluid will overflow around the edge of the flange [Kerr, 1997; Tivey et al., 1999]. As fluid overflows, mineral precipitation along the flange edge occurs. This precipitation, particularly that of silica, results in the lateral growth of the flange [Turner, 1995; Kerr, 1997; Tivey et al., 1999]. In general, flanges are compositionally abundant in sulfides, sulfates, and silica $\pm$ carbonates [Lonsdale and Becker, 1985; Delaney et al., 1992; Tivey et al., 1999]. Flange samples for this study are from the Guaymas Basin in the Gulf of California and also the Main Endeavour Field (MEF) located along the northern portion of the Juan de Fuca Ridge. Guaymas Basin is notable for its rapid rate of sediment deposition and hydrothermal


Figure 5.2: Diagrams illustrating the general structure and flow pathways within flanges, slabs, and crust with insets showing samples used in this study. a) Flange extending from the side of a larger chimney structure. Fluid emitted from the side of the chimney will pool under the flange and percolate upwards through the flange. b) Slab with circulating fluids. c) Crust with fingerlike protrusions at the base of the TAG black smoker edifice.
reactions between fluids and organic-rich sediments [Pearson et al., 2005]. As a result, calcite and barite are common constituents of vent deposits from this region [Koski et al., 1985]. The MEF is atypical in that it has unusually high fluid pH and high concentrations of methane and ammonia, which likely influence vent fluid reactions [Tivey et al., 1999].

Hydrothermal slabs (Figure 5.2b) are found at the Lucky Strike vent field along the Mid-Atlantic Ridge (MAR). The Lucky Strike vent field is unique in that it is situated adjacent to a large seamount that periodically produced lava flows that extend across the vent field [Langmuir et al., 1997]. Fluid temperature and chemistry at the vent field are highly controlled by the frequent volcanic eruptions, which have been found to change environmental conditions as often as every few days [Langmuir et al., 1997]. Slabs from Lucky Strike are hydrothermally silicified breccias composed of amorphous silica rimmed volcanic fragments (basaltic glass and plagioclase), barite, and sulfides [Cooper et al., 2000; Rouxel et al., 2004]. Slabs also exhibit layering parallel to the slab surface, likely developed in part to the ongoing volcanic activity. Research on the composition of fluid emitted from slabs suggests that seawater enters and circulates within the slab [Cooper et al., 2000; Rouxel et al., 2004]. This fluid will convect through the slab and eventually exit diffusely from cracks within the slab.

Crust samples (Figure 5.2c) are from the TAG vent field, also located along the MAR. This vent field is focused around a large central mound consisting of collapsed vent fragments [Thompson et al., 1985; Humphris et al., 1995]. Situated atop the center of the mound is a high-temperature black smoker
edifice. Hydrothermal crust forms a platelike, massive-sulfide layered surface that surrounds the TAG active mound upon which the black smoker edifice sits
[Humphris et al., 1995; Tivey et al., 1995]. The crust is formed through the gradual deposition and recementation of older vent debris and because of this, crust is highly cracked [Humphris et al., 1995; Tivey et al., 1995]. Crust also characteristically forms fingerlike protrusions [Tivey et al., 1995]. Fluid that has pulled within the black smoker edifice will seep out from these cracks and protrusions and rise, where it will become entrained within the black smoker plume [Tivey et al., 1995].

### 5.2 Permeability and Porosity

Probe permeability data provide a quantitative measure of permeability heterogeneity within the samples (Figure 5.3). Comparison of


Figure 5.3: Histograms of probe permeability data for flange, slab, and crust samples. Plots show frequency of measurements at a given permeability. Colors indicate sample locations: Guaymas $=$ purple, MEF = blue, Lucky Strike = gold and TAG $=$ green. a) Data for two Guaymas samples. b) Data for three MEF samples. c) Data for four Lucky Strike samples. d) Data for one TAG sample.
permeability values obtained from different sides of each sample show that measurements oriented parallel to the layering within the samples were predominantly on the order of $10^{-12} \mathrm{~m}^{2}$. Measurements taken perpendicular to the layering yielded a much broader range of values.

Cores were taken both parallel and perpendicular to layering from each of the samples in order to better quantify permeability anisotropy using the nitrogen permeameter. In total, 40 cores were obtained with 9 cores oriented parallel and 31 cores perpendicular to layering.

Pressure profiles from select cores showing the effects of confining pressure on permeability are shown in Figure 5.4. In general, the permeability decreased slightly with increases in confining pressure - likely due to crack


Figure 5.4: Plots showing pressure profiles for select samples during permeability measurement. Colors consistent with Fig. 4.3. Symbol shapes represent different sample cores (solid = pressure loading, empty = unloading) and asterisk indicates parallel-to-layering core. Generally negligible changes in permeability with pressure. a) Guaymas flange sample ALV 3517-R1 cores. b) MEF flange sample ALV 2415-1B cores. c) Lucky Strike slab sample ALV 2608-4-1, Pc 1cores. d) TAG crust sample ALV 2179-1-1 cores.
closure [e.g., Walsh, 1965]. Permeability values obtained during unloading agree
well with the values obtained during loading, indicating that the pressure effect is
mostly reversible with no permanent damage introduced during pressurization.

| Sample | Core | $\begin{gathered} \text { Probe } k \\ \left(\times 10^{-15} \mathrm{~m}^{2}\right) \end{gathered}$ | Core $k\left(\times 10^{-15} \mathrm{~m}^{\mathbf{2}}\right)$ |  | 2.7 MPa | Core $\phi$ (\%) |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $\phi \quad 2 \sigma$ | $\phi 2 \sigma$ |
| Guaymas Flange: |  |  |  |  |  |  |  |
| ALV 3517-R1 | D1 | $3230.4 \pm 3.6$ | $59.1 \pm 0.3$ | $55.1 \pm 0.3$ | $52.9 \pm 0.4$ | $2.3 \pm 3.3$ | $21.6 \pm 3.1$ |
|  | D3 | $1884.9 \pm 4.5$ | $32.0 \pm 0.3$ | $29.5 \pm 0.2$ | $28.0 \pm 0.2$ | $2.8 \pm 4.4$ | $19.8 \pm 4.0$ |
| ALV 3521-R2 | A2 | $29981.0 \pm 4.5$ | $1425.8 \pm 3.2$ | $1420.6 \pm 3.3$ | $1414.2 \pm 13.0$ | $11.5 \pm 0.1$ | $40.5 \pm 0.2$ |
|  | Ex 1 | -- | $1311.7 \pm 3.3$ | $1243.4 \pm 3.8$ | $1213.1 \pm 8.5$ | $2.1 \pm 0.4$ | $44.8 \pm 0.2$ |
|  | Ex 2 | -- | $969.7 \pm 4.2$ | $954.7 \pm 5.4$ | $953.8 \pm 2.0$ | $5.6 \pm 0.4$ | $44.7 \pm 0.3$ |
| MEF Flange: |  |  |  |  |  |  |  |
| ALV 2415-1B | A1 | $1694.3 \pm 3.8$ | $56.5 \pm 0.3$ | $55.4 \pm 0.4$ | $54.7 \pm 0.5$ | $20.2 \pm 1.2$ | $20.5 \pm 0.4$ |
|  | B1 | $9704.2 \pm 3.3$ | $1594.6 \pm 4.8$ | $1552.3 \pm 3.4$ | $1514.3 \pm 4.7$ | $31.2 \pm 0.3$ | $30.9 \pm 0.4$ |
|  | 1 | -- | $46.1 \pm$--- | $40.5 \pm$--- | $36.5 \pm$--- | $17.1 \pm$--- | $15.3 \pm$--- |
|  | 2 | -- | $5.3 \pm$--- | $3.5 \pm$--- | $2.6 \pm$--- | $19.2 \pm$--- | $17.2 \pm$--- |
| ALV 2927-3 | B1 | $2193.3 \pm 2.5$ | $1930.6 \pm 10.8$ | $1889.1 \pm 12.7$ | $1863.5 \pm 6.2$ | $38.2 \pm 0.1$ | $37.9 \pm 0.1$ |
|  | 1 | -- | $1950.2 \pm$--- | $1717.7 \pm$--- | $1595.6 \pm$--- | $41.8 \pm$--- | $39.6 \pm$--- |
|  | 2 | -- | $155.6 \pm$--- | $148.7 \pm$--- | $145.0 \pm$--- | $32.7 \pm$--- | $29.5 \pm$--- |
|  | 3 | -- | $994.2 \pm$--- | $967.3 \pm$--- | $948.7 \pm$--- | $40.7 \pm$--- | $38.2 \pm$--- |
| J2-286 | A1 | $2062.0 \pm 3.7$ | $1.3 \pm 0.1$ | $0.5 \pm 0.02$ | $0.2 \pm 0.02$ | $20.7 \pm 0.1$ | $20.5 \pm 0.3$ |
|  | A3 | $4181.9 \pm 4.3$ | $279.4 \pm 1.3$ | $233.1 \pm 1.3$ | $196.0 \pm 1.6$ | $25.3 \pm 1.6$ | $25.3 \pm 0.9$ |
|  | C2-1 | $46446.3 \pm 3.8$ | $0.8 \pm 0.04$ | $0.8 \pm 0.02$ | $0.7 \pm 0.03$ | $27.2 \pm 0.5$ | $27.3 \pm 0.3$ |
|  | C2-2 | -- | $1.3 \pm 0.04$ | $1.2 \pm 0.02$ | $1.2 \pm 0.1$ | $24.1 \pm 0.2$ | $24.2 \pm 0.4$ |
|  | C3-1 | $79602.4 \pm 3.2$ | $162.9 \pm 0.4$ | $146.0 \pm 1.3$ | $138.3 \pm 1.4$ | $27.6 \pm 4.5$ | $27.8 \pm 3.4$ |
|  | C3-2 | -- | $49.9 \pm 0.5$ | $48.7 \pm 0.2$ | $48.3 \pm 0.3$ | $21.2 \pm 1.1$ | $20.7 \pm 1.5$ |
|  | C4 | $94066.3 \pm 2.8$ | $18.7 \pm 0.1$ | $17.7 \pm 0.2$ | $17.0 \pm 0.1$ | $17.7 \pm 2.4$ | $19.9 \pm 2.1$ |
| Lucky Strike Slab: |  |  |  |  |  |  |  |
| ALV 2608-3-3 | B2 | $1960.4 \pm 4.4$ | $2744.7 \pm 13.5$ | $2737.1 \pm 11.9$ | $2746.1 \pm 5.6$ | $42.9 \pm 0.4$ | $42.5 \pm 0.3$ |
|  | C1 | $1051.5 \pm 4.5$ | $1445.3 \pm 5.9$ | $1447.1 \pm 3.1$ | $1436.6 \pm 3.5$ | $29.2 \pm 0.3$ | $28.8 \pm 0.3$ |
| ALV 2608-4-1 | A1 | $816.8 \pm 3.9$ | $787.8 \pm 2.7$ | $740.6 \pm 3.4$ | $706.5 \pm 1.1$ | $34.9 \pm 0.2$ | $34.8 \pm 0.3$ |
| Pc 1 | C1-1 | $428.4 \pm 3.9$ | $530.0 \pm 1.9$ | $486.7 \pm 0.4$ | $461.0 \pm 2.3$ | $37.2 \pm 1.0$ | $36.5 \pm 1.3$ |
|  | C1-2 | -- | $757.9 \pm 8.2$ | $655.1 \pm 5.1$ | $604.9 \pm 4.0$ | $46.4 \pm 0.8$ | $46.3 \pm 0.4$ |
|  | C3-1 | $1196.4 \pm 3.8$ | $28.1 \pm 0.2$ | $24.1 \pm 0.2$ | $22.2 \pm 0.2$ | $35.0 \pm 0.9$ | $34.8 \pm 0.3$ |
|  | C3-2 | -- | $120.3 \pm 1.2$ | $111.7 \pm 1.1$ | $106.9 \pm 0.8$ | $37.2 \pm 0.8$ | $37.1 \pm 0.6$ |
| ALV 2608-4-1 | A3 | $2319.9 \pm 4.6$ | $2056.1 \pm 4.5$ | $2016.9 \pm 6.5$ | $1994.5 \pm 5.5$ | $39.8 \pm 0.3$ | $39.6 \pm 0.1$ |
| Pc 2 | B3 | $1176.4 \pm 4.0$ | $561.6 \pm 2.2$ | $538.3 \pm 2.8$ | $523.0 \pm 0.9$ | $43.1 \pm 0.1$ | $42.8 \pm 0.2$ |
|  | 1 | -- | $5475.2 \pm$--- | $5023.6 \pm$--- | $4885.3 \pm$--- | $45.6 \pm$--- | $44.1 \pm$--- |
|  | 2 | -- | $661.4 \pm$--- | $574.6 \pm$--- | $534.1 \pm$--- | $48.0 \pm$--- | $46.3 \pm$--- |
|  | 4 | -- | $962.2 \pm$--- | $695.1 \pm$--- | $528.8 \pm$--- | $44.0 \pm$--- | $42.4 \pm$--- |
| JAS 177-2-1 | A2 | $3975.9 \pm 4.1$ | $198.5 \pm 1.8$ | $170.4 \pm 2.1$ | $152.1 \pm 1.5$ | $46.1 \pm 0.4$ | $45.7 \pm 0.4$ |
|  | B1 | $2729.9 \pm 3.5$ | $1827.7 \pm 3.8$ | $1774.8 \pm 2.9$ | $1745.9 \pm 4.5$ | $38.8 \pm 1.5$ | $38.9 \pm 1.2$ |
|  | B2 | $9435.7 \pm 3.2$ | $2998.6 \pm 13.7$ | $2954.2 \pm 11.8$ | $2898.5 \pm 4.1$ | $41.2 \pm 1.8$ | $41.0 \pm 1.7$ |
|  | C2 | $69963.2 \pm 3.9$ | $194.1 \pm 0.5$ | $192.1 \pm 0.6$ | $191.9 \pm 0.2$ | $41.2 \pm 3.7$ | $39.5 \pm 3.4$ |
| TAG Crust: |  |  |  |  |  |  |  |
| ALV 2179-1-1 | A1 | $13897.7 \pm 3.0$ | $605.0 \pm 2.8$ | $575.9 \pm 0.7$ | $560.7 \pm 1.2$ | $36.9 \pm 3.1$ | $35.3 \pm 4.3$ |
|  | A2 | $5234.4 \pm 3.6$ | $1106.7 \pm 6.9$ | $1079.1 \pm 5.5$ | $1062.7 \pm 2.9$ | $37.9 \pm 5.0$ | $37.8 \pm 3.2$ |
|  | A3 | $23030.1 \pm 1.4$ | $927.9 \pm 4.4$ | $887.3 \pm 4.7$ | $863.2 \pm 4.7$ | $38.4 \pm 4.5$ | $37.2 \pm 2.5$ |
|  | B2 | $1464.2 \pm 3.4$ | $1487.5 \pm 5.0$ | $1476.5 \pm 9.7$ | $1462.7 \pm 4.6$ | $43.2 \pm 0.2$ | $42.4 \pm 0.4$ |

Table 5.1: Average probe permeability, core permeability, and porosity and $2 \sigma$ values for flange, slab, and crust samples. Additional data for analyses was taken from cores without corresponding probe permeability measurements. Text in blue represents data for measurements taken parallel to layering.

The permeability reduction in all cores became negligible as the applied confining pressure increased from 2.1 MPa to 3.1 MPa . Thus, the measurements taken at 2.1 MPa were used for analyses. Table 5.1 shows the average permeability and porosity values at 2.1 MPa alongside the respective average probe permeability measurements. The majority of the cores measured had corresponding probe measurements, but additional measurements were also conducted on several cores taken where no probe permeameter measurements were made. Comparison between the permeability values obtained by the probe permeameter and those obtained by the nitrogen permeameter indicate that for the same sample, probe permeability is consistently higher. Part of the difference is due to the application of the confining pressure during core permeability measurements. Another reason is that flow measured by the probe permeameter is not directionally restricted, so the values obtained are a representation of a near surface volume average, whereas the nitrogen permeameter gives an axial permeability of a cylindrical core [Zhu et al., 2007]. In addition, probe measurements were sometimes made on surfaces that were trimmed off of cores (because the core measurements need to be made on cylindrical cores with parallel upper and lower surfaces). Notwithstanding these differences, data of the core samples reinforce the observation that within each sample, the parallel-tolayering permeability values do not show a lot of variability and are high, on the order of $10^{-12} \mathrm{~m}^{2}$, compared to the perpendicular-to-layering permeability values, which range over several orders of magnitude.

Permeability and porosity values obtained from these cylindrical cores can be divided into two groups: cores with axes oriented parallel-to-layering and cores with axes oriented perpendicular-to-layering (Figure 5.5). The parallel-to-layering cores all had relatively high permeabilities on the order of $10^{-12} \mathrm{~m}^{2}$ and porosities ranging from $\sim 30-40 \%$. For the perpendicular-to-layering cores, permeability values ranged from $10^{-16}-10^{-12} \mathrm{~m}^{2}$ and porosities ranged from $\sim 20-40 \%$, indicating larger variabilities compared to the parallel-to-layering cores.


Figure 5.5: a) Permeability versus porosity data for all of the cores. In general, permeability values for cores taken parallel-to-layering, symbols outlined in black, are higher $\left(\sim 10^{-12} \mathrm{~m}^{2}\right)$ than for cores taken perpendicular-to-layering $\left(10^{-16}-10^{-12}\right.$ $\mathrm{m}^{2}$ ). Differences in permeability as a function of porosity can be best fit by power-law relationships (black dashed lines), with a power-law exponent of $\alpha \sim 1$ (or 2) for the parallel-to-layering cores and $\alpha \sim 5$ (or 8 ) for the perpendicular-tolayering cores. Symbol color denotes location of sample origin: purple $=$ Guaymas Basin, blue $=$ MEF, yellow $=$ Lucky Strike, and green = TAG. b) Permeability versus porosity for only Lucky Strike cores with power-law relationships identified. c) Permeability versus porosity for MEF cores with power-law relationships labeled.

From Figure 5.5, two distinct trends of EPPRs are evident for the parallel-to-layering and perpendicular-to-layering cores. The exponent $\alpha \sim 1$ is for the
parallel-to-layering cores, whereas for the perpendicular-to-layering cores $\alpha \sim 5$. To correctly interpret the observed EPPRs, it is critical to relate the power-law relationships to the actual pore evolution processes using sample thin sections. Thin sections were made from a large subset of the cores, both parallel and perpendicular to the layers within the samples.

### 5.3. Microstructural Analyses

Microstructural analyses using reflected and transmitted light petrography were conducted on thin sections to identify pore evolution processes. Permeability-porosity data for samples from which thin sections were obtained are plotted in Figure 5.6. Data are grouped according to vent field and sample type: Guaymas flanges, MEF flanges, Lucky Strike slabs, and TAG crust. A summary of observations for each thin section is provided in Table 5.2.

### 5.3.1 Guaymas Flanges

The flanges have been separated into two groups, carbonate-dominated samples from the Guaymas, and sulfide-dominated samples from the MEF. The permeability and porosity values of two flange samples from Guaymas, ALV 3517-R1 and ALV 3521-R2, differ considerably (Figure 5.6a). For each sample, one thin section was cut axially through one core, and a second was cut transversely through another core (Table 5.2). For sample ALV 3517-R1, the axial cut (core D1) exposes a moderately layered structure with large calcite crystals $(\sim 250 \mu \mathrm{~m})$ at the base of the core, adjacent to where hot fluid was pooled. Smaller calcite crystals $(\sim 50 \mu \mathrm{~m})$ are present above these larger crystals, and


Figure 5.6: Permeability and porosity data for cores from which thin sections were obtained for microstructural analyses. Symbol colors are consistent with Figure 2 and are indicative of sample location: purple $=$ Guaymas Basin, blue $=$ MEF, yellow $=$ Lucky Strike, and green $=$ TAG. Different samples from each group are marked by differently shaped symbols, with black-rimmed symbols denoting parallel-to-layering cores. a) Guaymas flange data, b) MEF flange data, c) Lucky Strike slab data, and d) TAG crust data.

| Sample / Section | Mineral Present | Grain <br> Packing | Grain <br> Size ( $\mu \mathrm{m}$ ) | Pore Size ( $\mu \mathrm{m}$ ) | Pore Connectivity | Channel Width ( $\mu \mathrm{m}$ ) | Section Orientation |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{gathered} A L V 3517-R 1 \\ \text { D1 } \\ \text { D3 } \end{gathered}$ | $\begin{aligned} & \mathrm{ca}, \mathrm{st}(\mathrm{cp}, \mathrm{sp}) \\ & \mathrm{ca}, \mathrm{st}(\mathrm{cp}, \mathrm{sp}) \end{aligned}$ | tight <br> tight | $\begin{aligned} & 250 ; 50 \\ & 250 ; 50 \\ & \hline \end{aligned}$ | $\begin{aligned} & 350 ; 60 \\ & 350 ; 60 \end{aligned}$ | $\begin{aligned} & \text { low } \\ & \text { low } \end{aligned}$ | $\begin{aligned} & 10 \\ & 10 \\ & \hline \end{aligned}$ | $\begin{array}{r} (2) \\ (3) \\ \hline \end{array}$ |
| $\begin{gathered} \text { ALV } 3521-R 2 \\ \text { Ex } 1 \\ \text { Ex } 2 \\ \hline \end{gathered}$ | $\begin{aligned} & \text { ca, st (brt, cp, sp) } \\ & \text { ca, st (brt, cp, sp) } \end{aligned}$ | $\begin{aligned} & \text { tight } \\ & \text { tight } \\ & \hline \end{aligned}$ | $\begin{aligned} & 100 ; 50 \\ & 100 ; 50 \\ & \hline \end{aligned}$ | $\begin{aligned} & 300 ; 70 \\ & 300 ; 70 \\ & \hline \end{aligned}$ | moderate moderate | $\begin{aligned} & 10 \\ & 10 \\ & \hline \end{aligned}$ | $\begin{array}{r} (2) \\ (3) \\ \hline \end{array}$ |
| ALV 2927-3 <br> 2 <br> 3 <br> B1 <br> 2 | py, wz (po, cp, am Si) <br> py, wz (po, cp, am Si) <br> py, wz (po, cp, am Si) | tight <br> moderate <br> loose | $\begin{aligned} & 45 \\ & 50 \\ & 50 \end{aligned}$ | $\begin{array}{r} 50 \\ 80 \\ 30 \\ \hline \end{array}$ | moderate <br> high <br> high | $\begin{array}{r} 20 \\ 40 \\ 20 \\ \hline \end{array}$ | (2) <br> (3) <br> (1) |
| J2-286 <br> A1 <br> C2-1 <br> C3-1 <br> C3-2 | $\begin{aligned} & \text { py, po, am Si (wz) } \\ & \text { py, wz (po) } \\ & \text { py, wz, po (am Si) } \\ & \text { py, am Si (wz, cp) } \\ & \hline \end{aligned}$ | moderate <br> moderate <br> tight <br> tight | $\begin{gathered} 50 \\ 35 \\ 30 ; 150 \\ 50 \\ \hline \end{gathered}$ | $\begin{aligned} & 50 \\ & 75 \\ & 20 \\ & 20 \\ & \hline \end{aligned}$ | low moderate low low | $\begin{aligned} & 20 \\ & 20 \\ & 20 \\ & 20 \\ & \hline \end{aligned}$ | $\begin{array}{r} (2) \\ (3) \\ (2) \\ (2) \\ \hline \end{array}$ |
| $\begin{gathered} A L V 2415-1 B \\ \text { A1 } \\ \text { B1 } \\ \hline \end{gathered}$ | $\begin{aligned} & \text { py, po, am Si (cp, wz) } \\ & \text { py, po, am Si (cp, wz) } \end{aligned}$ | tight-mod moderate | $\begin{aligned} & 75 \\ & 75 \end{aligned}$ | $\begin{aligned} & 75 \\ & 50 \end{aligned}$ | $\begin{aligned} & \text { low } \\ & \text { moderate } \end{aligned}$ | $\begin{aligned} & 20 \\ & 20 \end{aligned}$ | $\begin{aligned} & (2) \\ & (1) \end{aligned}$ |
| $\begin{gathered} \text { ALV 2608-3-3 } \\ \text { B2 } \\ \text { C1 } \\ \hline \end{gathered}$ | $\begin{aligned} & \mathrm{gl}(\mathrm{pl}, \mathrm{am} \mathrm{Si}, \mathrm{py}) \\ & \mathrm{gl}, \mathrm{am} \mathrm{Si}(\mathrm{pl}, \mathrm{py}) \end{aligned}$ | loose <br> moderate | $\begin{aligned} & 300 \\ & 200 \\ & \hline \end{aligned}$ | $\begin{aligned} & 175 \\ & 150 \\ & \hline \end{aligned}$ | $\begin{gathered} \text { high } \\ \text { moderate } \end{gathered}$ | $\begin{aligned} & 40 \\ & 10 \\ & \hline \end{aligned}$ | $\begin{array}{r} (1) \\ (1) \\ \hline \end{array}$ |
| $\begin{gathered} \text { ALV 2608-4-1, Pc } 1 \\ \text { A1 } \\ \text { C3 (top) } \\ \text { C3 (bottom) } \\ \hline \end{gathered}$ | $\begin{aligned} & \mathrm{gl}, \mathrm{am} \mathrm{Si}, \mathrm{cl}(\mathrm{pl}) \\ & \mathrm{gl}, \mathrm{am} \mathrm{Si}, \mathrm{cl}(\mathrm{pl}) \\ & \mathrm{gl}, \mathrm{am} \mathrm{Si}, \mathrm{cl}(\mathrm{pl}) \end{aligned}$ | tight moderate tight | $\begin{gathered} 80 \\ 100 \\ 80 \\ \hline \end{gathered}$ | $\begin{gathered} 30 \\ 100 \\ 60 \\ \hline \end{gathered}$ | low moderate moderate | $\begin{aligned} & 15 \\ & 50 \\ & 20 \\ & \hline \end{aligned}$ | $\begin{array}{r} (1) \\ (2) \\ (2) \\ \hline \end{array}$ |
| ALV 2608-4-1, Pc 2 A3 B3 2 | $\begin{aligned} & \mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl}) \\ & \mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl}) \\ & \mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl}) \end{aligned}$ | moderate <br> tight <br> tight | $\begin{aligned} & 200 \\ & 100 \\ & 100 \\ & \hline \end{aligned}$ | $\begin{aligned} & 175 \\ & 175 \\ & 175 \\ & \hline \end{aligned}$ | high <br> high <br> high | $\begin{aligned} & 40 \\ & 40 \\ & 40 \\ & \hline \end{aligned}$ | $\begin{aligned} & (1) \\ & (2) \\ & (2) \\ & \hline \end{aligned}$ |
| $\begin{array}{r} \text { JAS } 177-2-1 \\ \text { A2 } \\ \text { B1 } \\ \text { B2 } \\ \hline \end{array}$ | $\begin{aligned} & \mathrm{gl}, \mathrm{pl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}) \\ & \mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl}) \\ & \mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl}) \end{aligned}$ | loose loose loose | $\begin{aligned} & 140 \\ & 200 \\ & 200 \end{aligned}$ | $\begin{aligned} & 150 \\ & 175 \\ & 175 \end{aligned}$ | high <br> high <br> high | $\begin{aligned} & 40 \\ & 50 \\ & 50 \end{aligned}$ | (2) <br> (1) <br> (1) |
| $\begin{gathered} \hline \text { ALV 2179-1-1 } \\ \text { A1 } \\ \text { B2 } \\ \hline \end{gathered}$ | $\begin{aligned} & \text { cp, py, sp } \\ & \text { cp, py, sp } \end{aligned}$ | moderate moderate | $\begin{array}{r} 30 \\ 50 \\ \hline \end{array}$ | $\begin{aligned} & 60 \\ & 60 \\ & \hline \end{aligned}$ | $\begin{aligned} & \text { high } \\ & \text { high } \\ & \hline \end{aligned}$ | $\begin{array}{r} 10 \\ 60 \\ \hline \end{array}$ | $\begin{array}{r} (3) \\ (1) \\ \hline \end{array}$ |
| ${ }^{(1)}$ parallel-to-layering core <br> thin section cut radially transects different layers |  | ${ }^{(2)}$ perpendicular-to-layering core <br> ${ }^{(3)}$ perpendicular-to-layering core <br> thin section cut axially <br> thin section cut radially transects different layers consists of a single layer |  |  |  |  |  |
| $\begin{aligned} \mathrm{am} \mathrm{Si} & =\text { amorphous } \\ \mathrm{pl} & =\text { plagioclase } \end{aligned}$ | $\text { ica } \quad \begin{array}{ll} \text { brt } & =\text { barite } \\ \text { po } & =\text { pyrrhotite } \end{array}$ | = calcite <br> $=$ pyrite | $\begin{aligned} & \mathrm{cp}=\mathrm{ch} \\ & \text { sp }=\text { sp } \end{aligned}$ | yrite <br> ite | $\begin{aligned} & =\text { clay } \\ & =\text { stevensite } \end{aligned}$ | $\begin{aligned} \mathrm{gl} & =\mathrm{palag} \\ \mathrm{wu} & =\text { wurtz } \end{aligned}$ | zed glass |

Table 5.2: Summary of microstructural observations for each thin section.
filling pore space between the larger crystals. Pore space is limited, occurring between clusters of calcite crystals, and there is little to no pore connectivity. The transverse section (core D3), taken through a region of the core near the transition from larger to smaller crystals, also reveals limited pore space and a lack of pore
connectivity (Figure 5.7a). Tight packing of calcite crystals limits pore space, pore connectivity, and likely inhibits flow and accounts for the low permeability values measured.


Figure 5.7: Cross polarized images of Guaymas flange cores. Width of each image $\sim 2.7 \mathrm{~mm}$. Larger blocky grains are calcite crystals; dark, elongated and grungy looking crystals are stevensite; white space is pore space. a) Section D3 is from sample ALV 3517-R1. Calcite crystals are very tightly packed. b) Section Ex 2 from sample ALV 3521-R2. Crystals structure contains larger pores.

For sample ALV 3521-R2, which has much higher permeability and porosity values, the axial cut (core Ex 1) is composed of large calcite crystals at the base of the core, adjacent to where hot fluid was pooled, and there are also patches of large crystals throughout the core. Calcite crystal size and packing is more variable in this sample, and there is more pore space (Figure 5.7b), than in sample ALV 3517-R1. Greater pore connectivity is observed in addition to greater porosity.

The distribution of calcite crystals observed in the Guaymas samples suggests that as $\mathrm{CO}_{2}$-rich fluids pool under the flange large calcite crystals precipitate along this surface under relatively constant conditions and high temperatures. Fluid that percolates upwards across a steep thermal gradient cools, and more variable conditions result in precipitation of smaller calcite crystals, stevensite, and sulfide minerals within pore space of upper parts of the flange.

### 5.3.2 MEF Flanges

The MEF flange samples are composed dominantly of sulfide minerals, not carbonate. Nine thin sections were made from three samples: ALV 2927-3 (3 thin sections), ALV 2415-1B (2 thin sections), and J2-286 (4 thin sections) (Figure 5.6b). Seven of the nine thin sections are from cores oriented perpendicular to layering, with five of these sections cut vertically to cross the different layers, and two cut horizontally, parallel to layering; the other two thin sections are from cores oriented parallel to layering, and were cut horizontally to cross the different layers. The thin sections oriented perpendicular-to-layering exhibit the full range of textures in each sample.

The thin section from the highest permeability core, ALV 2927-3 core B1, reveals layers composed of small to moderate sized crystals of wurtzite and pyrite (20 to $125 \mu \mathrm{~m}$, average $40 \mu \mathrm{~m}$ ), with minor amounts of similarly sized chalcopyrite and pyrrhotite and trace clay. Differences in crystal size and packing distinguish one layer from another, with different crystal packing resulting in a range of pore connectivity throughout the sample. For instance, one layer is particularly porous and well connected with channels averaging $\sim 80-100 \mu \mathrm{~m}$ in width; pore connectivity in other layers is not as high, though crystal packing is still loose. The thin section from core 3 of the same sample, a core with lower permeability but higher porosity, reveals only one layer (because it was cut parallel to layering), and thus likely is not representative of the entire core. The layer is similar in mineral content, texture, crystal size, packing, and pore connectivity to the layer of core B1 that exhibits the highest pore connectivity,
though in patches a very thin layer of amorphous silica is present (Figure 5.8a). The thin section from ALV 2927-3 core 2 reveals layers and textures very similar to those in the thin section from core $3, B 1$ of the same sample. Pore connectivity in some layers of core 2 is very low, with sulfide crystals more tightly packed, and some pore connectivity decreased and channels blocked by small amounts of late amorphous silica, particularly at the top of the core, nearest the upper flange surface.


Figure 5.8: Reflected light images. a) Section 3 is from sample ALV 2927-3. The grains are loosely packed. Large well-connected, equant void space is observed, which is usually an indication of high-connectivity and high permeability. b) Section A1 from sample ALV 2415-1B. Precipitation of amorphous silica causes void space reduction and likely causes a loss of pore connectivity, thus reducing permeability.

Thin sections from sample ALV 2415-1B cores B1 and A1 similarly reveal textures of several layers. The thin section from core A1 is composed of pyrite, with minor wurtzite, marcasite, trace pyrrhotite and chalcopyrite, and variable amounts of amorphous silica (Figure 5.8b). Fossil worm tube casts and clasts of outermost marcasite-rich upper flange layers that collected on the upper flange surface as debris are present near the top of the core, coated with a layer of late amorphous silica. Large isolated pores are present in this uppermost layer. Amounts of amorphous silica are least at the base of the core, closest to where hot fluid was pooled, and greatest near the top of the flange. Pore connectivity is only
present in layers where the sulfide crystals were initially widely spaced (loosely packed). The thin section from ALV 2415-1B core B1 also shows variable amounts of late stage amorphous silica, with greater amounts nearer the top of the flange. Due to the broad spacing of the initial sulfide crystals, pore connectivity and channel width have been retained in some of the layers. As with ALV 2927-3 core B1, the highest permeability layer of ALV 2415-1B core B1 should exert the greatest control on the overall permeability of the core.

For sample J2-286, four thin sections were made from cores oriented perpendicular to layering. The thin section from core C3-1 reveals layers of different mineral composition, crystal shape and size, porosity, and packing. The layer at the base of the core is composed of large pyrrhotite (blades $\sim 550 \mu \mathrm{~m}$ long) and cubanite, with large pore spaces adjacent to large crystals. Mid-layers of the core are composed of finer-grained resorbed pyrrhotite, pyrite and wurtzite, with the uppermost layer composed of mixed pyrite and wurtzite (crystals 10-50 $\mu \mathrm{m}$ in size). Cracks are present, with textures consistent with the cracks having been conduits for fluid in the past: one crack is lined with pyrite, and another filled with amorphous silica, and amorphous silica fills pore spaces adjacent to the crack. The thin section from core also reveals a large range in crystal sizes through the layers. The layers are composed primarily of pyrite crystals with minor amounts of wurtzite and marcasite, and nearly all crystals are heavily coated with amorphous silica. The precipitation of amorphous silica in many of the layers appears to have severely restricted and in some cases blocked flow channels between crystals. Although, in one of the layers the crystals were
initially widely spaced, so that even with the later precipitation of the amorphous silica coating there was still ample space between crystals to accommodate flow. Fossil tube worm casts are present in one layer, with the tubes $50 \%$ filled by later precipitated pyrite and amorphous silica spherules. The thin section from core A1 is similar compositionally to that from core C3-2, with the exception that abundant amorphous silica is present coating crystals and filling what was pore space in the majority of the layers in core A 1 , resulting in little to no pore connectivity. There are layers present that do still contain pore space, but these pores are isolated. The thin section from core $\mathrm{C} 2-1$ is different from the other thin sections in that it was taken through one of the core's layers rather than through the whole core. The layer exhibits narrow channels that are moderately well connected. Amorphous silica is absent, some Fe-oxide is present at crystal boundaries, and the fossil worm tube casts present are much smaller in size than in other parts of this sample. It is likely that the low permeability measured for this core may be indicative of the presence of substantial amorphous silica elsewhere in the core, in a layer that was not transected by the thin section.

Flange growth can be closely linked to the microstructures observed in the thin sections, as was true for the Guaymas flanges. Larger sulfide crystals develop along the base of the flange overlying the pooled fluid, under relatively constant conditions. Pore space between these large sulfide crystals accommodates the upward migration of hot fluids. As the fluids cool, smaller sulfide crystals precipitate throughout the flange under more variable conditions. Conductive cooling of vent fluid or vent fluid/seawater mixtures result in
saturation and deposition of amorphous silica on existing surfaces as a thin layer. Over time the precipitation of angular sulfide crystals and coatings of silica will block flow pathways and limit the permeability through the flange.

### 5.3.3 Lucky Strike Slabs

The slab samples are composed of shards of palagonitized glass $\pm$ plagioclase shards and later stage amorphous silica and clay. Eleven thin sections were made from four slab samples: ALV 2608-3-3 (2 sections), ALV 2608-4-1, Pc 1 (3 sections), ALV 2608-4-1, Pc 2 (3 sections), and JAS-177-2-1 (3 sections) (Figure 5.6c), all cut through the cores such that the samples' layers were present in each section. Layers were delineated by changes in grain packing and often grain size. Thin sections from five perpendicular-to-layering cores were examined to identify microstructural features that might explain the relatively steep ( $\alpha \sim 8$ ) permeability-porosity trend for these cores. The thin section from the most permeable core (ALV 2608-4-1, Pc 2 core 2 ) reveals moderately sorted, highly fragmented shards of palagonitized glass $(\sim 100 \mu \mathrm{~m})$, in each of its layers, and amorphous silica and clay coat many of the glass grains; however, pore connectivity remains intact through each of the layers. The thin section from core B3 of the same sample is very similar and contains layers of highly fragmented glass shards, although within this section greater variability in the structure of the layers is apparent. There are two layers visible within this section that can account for the somewhat lower measured permeability and porosity values: a layer of slightly smaller and more tightly packed grains, and a layer where pore space between grains has been almost entirely filled with amorphous silica and clay.

These restrictive features were not observed in the thin section from core 2 . The thin section from sample JAS 177-2-1 core A2 is also composed of palagonitized glass shards, though shard size varies considerably through the different layers. There is a greater abundance of plagioclase shards than in the previously described thin sections, although they too vary in size and shape. The shards are widely spaced, but in several of the layers there is a thick amorphous silica and clay coating that fills previous pore space and blocks flow channels between shards. The lower permeability values can be attributed to this high degree of channel restriction. Thin sections from cores C3-1 and C3-2 of sample ALV 2608-4-1, Pc 1 exhibit similar textures - a mixture of large and small palagonitized glass shards coated in amorphous silica. Much of the space between shards has been infilled with clay that blocks channels and isolates many of the pores. Microstructural observations are consistent with core C3-1 being least permeable because it includes a layer that has been severely infilled by clay and amorphous silica precipitation.

Thin sections from parallel-to-layering slab cores reinforce the observations from the previously described slab thin sections. The most and least permeable of these cores are from sample ALV 2608-3-3. Thin sections from both cores B2 and C1 reveal poorly sorted palagonitized glass shards (and a few plagioclase shards) coated with amorphous silica and minor clay. In the thin section from core B2, grains are loosely packed, and the coating of amorphous silica narrows, but seldom blocks channels (Figure 5.9a, b). In the thin section from core C 1 , glass shards are more densely packed than in core B 1 , and coatings


Figure 5.9: Cross polarized slab images showing variation in grain packing and precipitation between sample layers. Width of images is 2.7 mm . a and b) Section B2 from sample ALV 2608-3-3. A layer with large, broadly spaced glass shards is shown in a), while a more tightly packed layer with highly fragmented grains is shown in b). Close packing of angular grains can cause pinch-offs in void space around grains that can typically lead to limited pore connectivity and a lower permeability. c and d) Section B2 from sample JAS 177-2-1. A well connected, high permeability layer is shown in c), whereas a layer with pore space that had been densely infilled by the precipitation of amorphous silica and clay is shown in d). The precipitation of amorphous silica and clay results in void space reduction, which generally leads to a loss of pore connectivity, thereby reducing permeability.
of amorphous silica and clay are thicker, resulting in less pore connectivity. However grain packing is less dense, and pore connectivity greater, in the layer of core C 1 nearest the top of the slab; this layer likely explains why the permeability of this core (made parallel to layering) has a high permeability despite a low porosity. Textures observed in the thin section from sample JAS 177-2-1 cores B2 and B1 are very similar to those from core A2 from that sample. As with core A2, there is considerable variability between the structure and packing, with both highly porous and well connected layers alternating with a layer that has been heavily coated with amorphous silica and clay (Figure 5.9c, d). Permeability of the cores oriented parallel to layering is likely controlled by the layer with the
highest pore connectivity. The thin section from sample ALV 2608-4-1, Pc 2 core A3 reveals both amorphous silica and clay precipitated throughout the various layers, but pore space and pore connectivity are still both high, as in this sample's core 2 . The core with the lowest permeability of the parallel-to-layering cores is ALV 2608-4-1, Pc 1 core A1. The thin section from this core exhibits abundant amorphous silica and clay, and pore connectivity appears lower than in the other parallel-to-layering cores.

From a mineralogical and textural perspective, seafloor hydrothermal slabs can be classified as hyaloclastites that form from the interaction of hot magma with seawater. Hyaloclastites contain glass shards, as seen in the slab thin sections, which form as thermal stresses break apart large pieces of volcanic glass. The orientation of the slab layers and the grain size within each layer suggests that smaller shards settle to the bottom of the slab while larger fresher grains are found predominantly along the slab surface. Layers are delineated by differences in shard size, initial packing density, and late stage amorphous silica and clay deposition. The presence of amorphous silica indicates that silica-rich vent fluids, or mixtures of vent fluids and seawater, have percolated through the slabs with cooling of the fluids resulting in amorphous silica and clay saturations [e.g., Tivey et al., 1999) and deposition along grain edges, resulting in cementation of some slab layers.

The slab samples are notable in that the perpendicular-to-layering cores plot at a slightly steeper trend than the other vent deposit sample groups (Figure $5.5 b$ ), which may be due to the high angularity of the glass grains. Angular grains
create roughness along flow pathways that can significantly impact the deposit's permeability. Any changes in grain shape over time can have a large effect on permeability while not producing major changes in the overall porosity, resulting in a steeper trend.

### 5.3.4 TAG Crust

Measurements were made on only one crust sample, ALV 2179-1-1. Thin sections were made form core A1 (from a perpendicular-to-layering core) and core B2 (from a parallel-to-layering core) (Figure 5.6d). The thin section from core A1 was cut horizontally through just one of the sample's layers. The layer is composed dominantly of fine-grained chalcopyrite and pyrite $(\sim 30 \mu \mathrm{~m})$ with pore space $(\sim 60 \mu \mathrm{~m})$ along crystal edges. Pore connectivity in this layer is high, but through narrow ( $\sim 10 \mu \mathrm{~m}$ ) channels. Unfortunately no information on pore connectivity in other layers of this sample is available. The thin section from core B2 was cut


Figure 5.10: Reflected light images of layering through crust sample ALV 2179-1-1, core B2. Sulfide-rich layers grade from loosely packed crystals (high connectivity) to more closely packed crystals (lower connectivity). Images are each $\sim 2.7 \mathrm{~mm}$ wide.
across the layering (Figure 5.10). Textures and mineral contents are similar to
those in core A1. Layering was apparent only through changes in crystal size. High pore connectivity is consistent through most of the layers and can explain the high permeability value.

Because crust deposits on the TAG active mound are situated adjacent to high-temperature black smokers, fluids beneath the crust are likely hot (around $300^{\circ} \mathrm{C}$ ). The crust sample appears to have developed much like sulfide flange deposits, with hot fluid percolating upwards from the base of the sample [Tivey et al., 1995]. Textures are consistent with this, with large sulfide crystals at the bottom of the crust where conditions are likely relatively constant. Packing of the crystals is loose, providing space at crystal boundaries to accommodate fluids moving up through the deposit. As the fluid travels upwards it cools, resulting in precipitation of smaller more closely packed crystals under more variable conditions.

### 5.4. Discussion

In this study, permeability and porosity measurements were conducted on flange, slab, and crust samples, each of which exhibit layering that parallels upper and lower surfaces of the deposits. The data document that permeability values in the direction parallel to layering are considerably less variable and higher than permeability values in the direction perpendicular to layering. These differences in permeability suggest different flow behaviors caused by layering, and have broad implications for the overall fluid flux accommodated by flange, slab, and crust seafloor deposits. Given observed permeabilities, fluid will travel horizontally through these deposits, within prominent highest permeability
deposit layers. This layer will be able to continually facilitate the lateral flow of fluid through the deposits, because permeability decreases little as porosity decreases (trend of $\alpha \sim 1$ ).

The difference in magnitude as well as in variation of permeability values of cores taken parallel-to-layering versus perpendicular-to-layering in all samples can be quantified using an effective permeability model. Because the total volume flux parallel to layering is equal to the sum of the volume flux through each layer, the effective permeability $\left(k_{p a l}\right)$ in this case is the sum of each layer's permeability $\left(k_{i}\right)$ multiplied by the fraction of the total thickness that layer constitutes $\left(h_{i} / H\right)$ (5.1), where $h_{i}$ and $H$ are the thickness of each layer and the total thickness of the sample, respectively (e.g., Freeze and Cherry, 1978):

$$
\begin{equation*}
k_{p a l}=\sum_{i=1}^{n} k_{i} \frac{h_{i}}{H} \tag{5.1}
\end{equation*}
$$

In comparison, the fluid flux perpendicular to layering must obey mass conservation while crossing several layers of varying permeabilities. The effective permeability $\left(k_{p e p}\right)$ perpendicular-to-layering equals the total deposit thickness $(H)$ divided by the sum of the ratios between layer thicknesses $\left(h_{i}\right)$ to their respective permeabilities $\left(k_{i}\right)$ :

$$
\begin{equation*}
k_{p e p}=H / \sum_{i=1}^{n} \frac{h_{i}}{k_{i}} \tag{5.2}
\end{equation*}
$$

From eqn. (5.1) and (5.2), it is easy to see that permeability of a layered vent deposit is generally anisotropic, with the parallel-to-layering effective permeability $k_{p a l}$ greater than the perpendicular-to-layering effective permeability $k_{p e p}$.

The parallel versus serial flow patterns within the deposit cores are supported by microstructural observations from sample thin sections. Petrographic examination shows that the flange, slab, and crust deposits generally consist of layers with large contrasts in crystal packing, void space, and pore connectivity (Figures 5.8-5.10). Low-permeability layers, resulting from initial differences in crystal packing, and/or subsequent deposition of crystals in pore space, and/or late stage precipitation of amorphous silica $\pm$ clay, are observed in both the parallel- and perpendicular-to-layering cores. These low permeability layers restrict the overall flux perpendicular to layering thus exerting primary control on $k_{\text {pep }}$, whereas the flux parallel to layering and thus $k_{p a l}$ is affected primarily by the highest permeability layer. This explains why the permeability values in the parallel-to-layering cores are consistently higher than those in the perpendicular-to-layering cores and why the perpendicular-to-layering cores exhibit much greater permeability variations.

This difference in parallel flow and serial flow is best seen by comparing textures observed in thin sections from cores B1 and 2 of sample ALV 2927-3. Both of these thin sections clearly show the same sequence of sample layering with some layers of relatively high pore connectivity, and thus likely permeability, and some with lower pore connectivity and thus likely permeability. Although they are similar, core B1, oriented parallel-to-layering, has a measured permeability over an order of magnitude greater than that of core 2 , which was oriented perpendicular-to-layering.

The microstructural observations provide explanations for the two EPPRs identified for these deposits. For the parallel-to-layering cores, with an EPPR with an exponent of $\alpha \sim 1$, the change in the effective permeability of the samples are relatively small, even for large changes in porosity. The thin sections taken from all of the cores show that the layers within these cores have undergone pore evolution processes, such as late-stage precipitation of amorphous silica or clays or thermal cracking of crystals. Textures resulting from these processes have been observed in the sample thin sections to significantly change the porosity of the layers. With late-stage amorphous silica precipitation, precipitation can initially result in smoothing of crystal edges, but after larger amounts of precipitation these mineral coatings can pinch-off, block, or completely infill pore space and channels along crystal edges. The parallel-to-layering cores have experienced a loss of porosity as a result of these processes; however, layers of high permeability which were retained during these processes will still allow the deposit to accommodate a high flux of fluid.

Data from the perpendicular-to-layering cores reveal an EPPR with an exponent of $\alpha \sim 5$, and pore evolution processes that are significantly more effective in changing the overall permeability. Flow through these samples, perpendicular to layering, is serial. Changes in pore space of least permeable layers will restrict this serial flow and lower the effective permeability of the sample. These effects of pore evolution processes on the sample layers and the effective permeability of the deposits are important for modeling their fluid fluxes.

## Chapter 6: Spire Deposits

Hydrothermal spires are tower-like deposits that grow vertically upwards from the seafloor. Zn -rich actively diffusing spires, black smoker chimneys, and relict spires are all spires deposits. These deposits were grouped together following data collection and analyses, because they have similar structural features and likely experience similar evolution processes. In total the spire sample set includes 9 actively diffusing spires, 6 black smoker chimneys, and 8 relict spires. These spires come from several different vent fields: ABE, MEF, TAG, Cleft, and Kilo Moana.

### 6.1 Permeability and Porosity

Probe permeability measurements for the Zn -rich actively diffusing spires, black smoker chimneys and relict spires are plotted in Figures 6.1, 6.2, and 6.3 respectively. The Zn -rich actively diffusing spires were the most permeable with values ranging from $\sim 6 \times 10^{-13}-8 \times 10^{-10} \mathrm{~m}^{2}$. Permeability anisotropy is observed in most of


Figure 6.1: Histograms showing probe permeability data for Zn -rich diffusing spire samples. Plots show frequency of measurements at a given permeability. Colors represent individual samples. b) Data for four ALV samples. b) Data for three J 2 samples. these samples. In general, permeability measurements made near the center of the spire are lower than those made along the outer rim of the sample. The black


Figure 6.2 (above): Histograms showing probe permeability data for black smoker chimney samples. Plots show frequency of measurements at a given permeability. Colors represent individual samples. a) Data for three ALV samples. b) Data for three J2 samples.

Figure 6.3 (right): Histograms showing probe permeability data for relict spire samples. Plots show frequency of measurements at a given permeability. Colors represent sample locations: $\mathrm{ABE}=$ purple, $\mathrm{Cleft}=$ gold, Kilo Moana $=$ blue and TAG $=$ green. a) Data for two ABE samples. b) Data for two Cleft samples. c) Data for one Kilo Moana sample. d) Data for one TAG sample.

smoker chimneys, which had permeability values ranging from $\sim 2 \times 10^{-14}-8 \times 10^{-11}$ $\mathrm{m}^{2}$, also showed permeability variations within samples. Permeability measurements made along the inner chalcopyrite lining were lower than those made along the outer, more anhydrite-rich layers. For the relict spires, probe permeability values ranged between $\sim 4 \times 10^{-14}-2 \times 10^{-10} \mathrm{~m}^{2}$. The relict spires had
more pronounced variability in measurements made along the same direction, suggesting that these samples may have a higher degree of surface roughness than the other spire samples. The few measurements that were obtained axially along spires have slightly lower permeability values than those made radially for the same sample.

Similar to the other discussed sample types, varying degrees of surface roughness can account for some of the observed variability in the probe permeability data. High degrees of surface roughness can lead to an insufficient permeameter seal against the sample surface, which can increase the potential for air to be leaked during measurement. An improper seal can lead to variations in the data that are not representative of the samples' true permeability anisotropy.

After completing probe permeability measurements, cylindrical cores were taken from select probe measurement sites for axial permeability and porosity measurement. In total, 22 Zn -rich diffusing spire cores, 11 black smoker chimney cores, and 34 relict spire cores were made taken from the deposit samples. Permeability and porosity data for the Zn -rich diffusing spires, black smoker chimneys and relict spires are listed in Tables 6.1, 6.2, and 6.3, respectively. Confining pressure from approximately $0.5-2.7 \mathrm{MPa}$ was incrementally applied to the cores during measurement. Pressure profiles for the Zn -rich diffusing spires, black smoker chimneys and relict spires are shown in Figures 6.4, 6.5, and 6.6, respectively.

| Sample | Core | $\begin{gathered} \text { Probe } k \\ \left(\times \mathbf{1 0}^{-15} \mathbf{m}^{2}\right) \end{gathered}$ | Core $k \quad\left(\times 10^{-15} \mathrm{~m}^{2}\right)$ |  |  | Core $\phi$ (\%) |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $\phi \quad 2 \sigma$ | $\phi \quad 2 \sigma$ |
| $\begin{aligned} & \text { ALV 2187-1-1 } \\ & \text { (top) } \end{aligned}$ | A2 | $8339.8 \pm 3.6$ | $6111.0 \pm 42.5$ | $5462.3 \pm 35.7$ | $4847.3 \pm 30.4$ | $45.5 \pm 0.3$ | $44.1 \pm 0.2$ |
|  | A4 | $8386.7 \pm 4.1$ | $4013.4 \pm 17.4$ | $3433.7 \pm 24.3$ | $3038.1 \pm 14.8$ | $36.5 \pm 0.2$ | $35.8 \pm 0.2$ |
|  | B1 | $27560.5 \pm 4.1$ | $613.1 \pm 1.1$ | $596.3 \pm 2.5$ | $578.6 \pm 1.1$ | $45.5 \pm 0.6$ | $44.2 \pm 0.7$ |
| ALV 2187-1-1 <br> (bottom) | A2 | $25910.6 \pm 2.0$ | $340.2 \pm 1.2$ | $332.9 \pm 1.1$ | $322.6 \pm 0.9$ | $38.9 \pm 0.5$ | $38.4 \pm 0.3$ |
|  | B1 | $31357.7 \pm 3.3$ | $590.4 \pm 1.3$ | $581.7 \pm 4.5$ | $557.6 \pm 2.1$ | $42.1 \pm 0.2$ | $41.8 \pm 0.3$ |
|  | C2 | $1137.5 \pm 3.9$ | $618.9 \pm 2.8$ | $551.0 \pm 2.3$ | $523.8 \pm 1.3$ | $41.0 \pm 0.2$ | $40.5 \pm 0.3$ |
|  | C3 | $641.7 \pm 4.1$ | $1709.6 \pm 10.0$ | $1467.9 \pm 6.7$ | $1306.2 \pm 5.4$ | $40.6 \pm 0.4$ | $40.0 \pm 0.4$ |
| ALV 2187-1-2 | B2-2 | -- | $1817.8 \pm 5.7$ | $1748.7 \pm 2.8$ | $1704.0 \pm 5.3$ | $41.8 \pm 0.2$ | $41.3 \pm 0.2$ |
|  | C1/C2 | $216149.9 \pm 1.4$ | $1813.7 \pm 17.6$ | $1619.4 \pm 14.4$ | $1476.5 \pm 9.0$ | $39.4 \pm 0.1$ | $38.7 \pm 0.2$ |
|  | C3-1 | $2372.5 \pm 3.3$ | $2657.3 \pm 11.0$ | $2602.0 \pm 11.2$ | $2534.5 \pm 3.6$ | $41.3 \pm 0.3$ | $40.9 \pm 0.2$ |
|  | C3-2 | -- | $2809.0 \pm 15.5$ | $2670.0 \pm 7.1$ | $2501.7 \pm 9.2$ | $42.4 \pm 0.3$ | $41.5 \pm 0.5$ |
|  | C5 | $7580.9 \pm 4.0$ | $1460.9 \pm 5.5$ | $1426.3 \pm 3.3$ | $1400.6 \pm 5.4$ | $42.9 \pm 0.2$ | $42.2 \pm 0.1$ |
| ALV 2190-14-1 | 1 A1 | $4066.2 \pm 4.5$ | $2802.2 \pm 25.4$ | $2726.3 \pm 9.8$ | $2661.7 \pm 7.4$ | $47.5 \pm 0.2$ | $46.7 \pm 0.2$ |
|  | B1 | $12847.6 \pm 3.1$ | $1195.5 \pm 3.1$ | $1140.0 \pm 2.6$ | $1073.3 \pm 6.1$ | $47.6 \pm 0.6$ | $46.1 \pm 0.6$ |
|  | B2 | $8541.1 \pm 3.7$ | $797.7 \pm 3.2$ | $775.1 \pm 3.0$ | $761.1 \pm 1.9$ | $41.3 \pm 0.5$ | $41.0 \pm 0.5$ |
| J2-128-8-R1 | Ex | -- | $29.0 \pm 0.5$ | $22.2 \pm 0.3$ | $18.2 \pm 0.1$ | $28.0 \pm 0.2$ | $27.4 \pm 0.5$ |
| J2-137-7-R1 | B1 | $239125.2 \pm 2.5$ | $1181.7 \pm 6.0$ | $1068.4 \pm 5.8$ | $994.2 \pm 3.2$ | $47.8 \pm 1.0$ | $44.3 \pm 0.7$ |
| J2-222-1-R1 | A1 | -- | $367.8 \pm 2.6$ | $327.3 \pm 1.0$ | $305.8 \pm 1.2$ | $35.0 \pm 0.3$ | $34.1 \pm 0.3$ |
|  | A2 | -- | $231.3 \pm 0.3$ | $227.7 \pm 0.5$ | $225.9 \pm 0.6$ | $36.0 \pm 0.5$ | $35.6 \pm 0.7$ |
|  | Ex | -- | $220.8 \pm 1.6$ | $202.9 \pm 1.0$ | $190.8 \pm 0.8$ | $29.8 \pm 0.8$ | $29.4 \pm 0.5$ |
| J2-127-1-R2 | B2 | -- | $379.2 \pm 0.4$ | $371.1 \pm 0.6$ | $366.1 \pm 1.1$ | $37.4 \pm 0.5$ | $37.1 \pm 0.5$ |
|  | B3 | -- | $444.1 \pm 1.2$ | $438.2 \pm 1.0$ | $432.3 \pm 1.2$ | $36.2 \pm 0.2$ | $35.7 \pm 0.3$ |

Table 6.1: Average probe permeability, core permeability, and porosity and $2 \sigma$ values for Zn -rich actively diffusing spires.

| Sample | Core | $\begin{gathered} \text { Probe } k \\ \left(\times 10^{-15} \mathbf{m}^{2}\right) \end{gathered}$ | Core $k \quad\left(\times 10^{-15} \mathrm{~m}^{2}\right)$ |  |  | Core $\phi$ (\%) |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | 1.4 MPa | 2.1 MPa | 2.7 MPa | 2.1 MPa | 2.7 MPa |
|  |  | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $\phi \quad 2 \sigma$ | $\phi \quad 2 \sigma$ |
| ALV 1445-3 | C1 | $129.6 \pm 1.7$ | $21.5 \pm 0.3$ | $16.1 \pm 0.2$ | $13.5 \pm 0.2$ | $24.9 \pm 0.3$ | $24.0 \pm 0.8$ |
| ALV 2179-4-1 | A1 | $1230.4 \pm 2.9$ | $167.6 \pm 0.6$ | $135.6 \pm 0.4$ | $113.7 \pm 0.5$ | $42.7 \pm 1.0$ | $41.4 \pm 0.7$ |
|  | B1 | $309.4 \pm 2.4$ | $27.0 \pm 0.7$ | $17.7 \pm 0.4$ | $12.9 \pm 0.2$ | $36.6 \pm 0.4$ | $35.0 \pm 0.6$ |
|  | C1 | $3675.5 \pm 2.4$ | $132.1 \pm 0.8$ | $124.1 \pm 0.3$ | $122.1 \pm 0.2$ | $42.2 \pm 0.4$ | $41.4 \pm 0.5$ |
| J2-137-1-R1 | D1 | $111.6 \pm 0.1$ | $0.7 \pm 0.02$ | $0.6 \pm 0.03$ | $0.6 \pm 0.02$ | $16.9 \pm 0.5$ | $16.7 \pm 0.5$ |
|  | D2 | $61.3 \pm 0.2$ | $0.4 \pm 0.03$ | $0.4 \pm 0.02$ | $0.4 \pm 0.02$ | $13.9 \pm 0.4$ | $13.8 \pm 0.5$ |
|  | D3 | $223.0 \pm 0.1$ | $0.8 \pm 0.1$ | $0.7 \pm 0.02$ | $0.7 \pm 0.02$ | $18.0 \pm 0.4$ | $17.8 \pm 0.3$ |
|  | D4 | $90.0 \pm 0.3$ | $0.6 \pm 0.02$ | $0.6 \pm 0.01$ | $0.6 \pm 0.02$ | $17.7 \pm 0.3$ | $17.4 \pm 0.5$ |
| J2-213-3-R1 | A1 | $12146.5 \pm 2.9$ | $20.3 \pm 0.1$ | $18.4 \pm 0.1$ | $17.5 \pm 0.2$ | $28.9 \pm 0.5$ | $28.6 \pm 0.8$ |
|  | B1 | $309.4 \pm 3.2$ | $40.5 \pm 0.4$ | $37.7 \pm 0.3$ | $34.9 \pm 0.3$ | $30.6 \pm 0.1$ | $30.1 \pm 0.4$ |
|  | D1 | $3675.5 \pm 3.7$ | $115.2 \pm 0.6$ | $99.5 \pm 1.4$ | $86.6 \pm 0.6$ | $39.0 \pm 0.6$ | $38.6 \pm 0.8$ |

Table 6.2: Average probe permeability, core permeability, and porosity and $2 \sigma$ values for black smoker chimneys.

| Sample Core | $\begin{gathered} \text { Probe } k \\ \left(\times \mathbf{1 0}^{-15} \mathbf{m}^{2}\right) \end{gathered}$ | Core $k\left(\times 10^{-15} \mathrm{~m}^{2}\right)$ |  |  | Core $\phi$ (\%) |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | 1.4 MPa | 2.1 MPa | $\underline{\text { 2.7 MPa }}$ | $\underline{2.1 ~ M P a}$ | 2.7 MPa |
|  | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $k \quad 2 \sigma$ | $\phi \quad 2 \sigma$ | $\phi \quad 2 \sigma$ |
| ABE: |  |  |  |  |  |  |
| J2-129-1-R3 B2 | $88.0 \pm 2.9$ | $751.8 \pm 5.9$ | $740.4 \pm 3.0$ | $731.8 \pm 2.7$ | $41.0 \pm 0.2$ | $40.2 \pm 0.1$ |
| J2-136-6-R1 A1-1 | $152.5 \pm 3.7$ | $155.0 \pm 1.3$ | $140.1 \pm 1.0$ | $133.5 \pm 1.0$ | $31.3 \pm 7.8$ | $30.0 \pm 4.7$ |
| A1-2 | -- | $194.5 \pm 1.2$ | $190.8 \pm 1.4$ | $188.8 \pm 1.6$ | $30.2 \pm 1.1$ | $29.8 \pm 1.3$ |
| C2 | $448.1 \pm 2.3$ | $22.6 \pm 0.4$ | $19.4 \pm 0.3$ | $17.5 \pm 0.2$ | $25.4 \pm 2.4$ | $24.8 \pm 2.0$ |
| Cleft: |  |  |  |  |  |  |
| ALV 2944-3-S1 A1 | $51095.7 \pm 3.8$ | $633.1 \pm 3.4$ | $626.4 \pm 4.8$ | $624.2 \pm 3.1$ | $37.6 \pm 10.0$ | $36.3 \pm 6.2$ |
| Pc 1 A3 | $41748.7 \pm 3.6$ | $43.0 \pm 0.3$ | $40.9 \pm 0.3$ | $39.7 \pm 0.2$ | $30.9 \pm 5.7$ | $31.2 \pm 4.5$ |
| B1 | $27100.4 \pm 4.2$ | $859.4 \pm 5.9$ | $825.7 \pm 8.8$ | $712.9 \pm 7.1$ | $36.0 \pm 8.5$ | $34.6 \pm 5.0$ |
| ALV 2944-3-S1 A1 | -- | $331.4 \pm 0.7$ | $327.2 \pm 1.1$ | $325.6 \pm 0.4$ | $33.5 \pm 5.8$ | $31.1 \pm 10.1$ |
| Pc 2 A2 | -- | $651.9 \pm 1.3$ | $645.8 \pm 2.5$ | $639.6 \pm 1.4$ | $42.6 \pm 4.5$ | $42.1 \pm 3.3$ |
| A3 | -- | $566.8 \pm 3.2$ | $562.9 \pm 1.1$ | $560.7 \pm 1.0$ | $40.6 \pm 6.5$ | $41.9 \pm 6.2$ |
| A4 | -- | $426.5 \pm 1.5$ | $404.5 \pm 1.0$ | $392.7 \pm 0.8$ | $31.6 \pm 4.7$ | $29.8 \pm 5.9$ |
| ALV 2941-6-S1 A1 | $126826.6 \pm 2.5$ | $194.5 \pm 0.4$ | $140.6 \pm 0.3$ | $117.4 \pm 0.1$ | $32.5 \pm 4.4$ | $32.6 \pm 5.1$ |
| Kilo Moana: |  |  |  |  |  |  |
| J2-125-3-B1 B2-1 | $3953.7 \pm 3.3$ | $440.7 \pm 1.5$ | $356.3 \pm 1.6$ | $312.2 \pm 1.4$ | $34.7 \pm 3.5$ | $34.9 \pm 2.3$ |
| B2-2 | -- | $156.6 \pm 0.7$ | $149.4 \pm 0.3$ | $146.6 \pm 0.4$ | $35.4 \pm 5.1$ | $32.7 \pm 5.7$ |
| B3-1 | $1189.7 \pm 4.1$ | $98.6 \pm 0.7$ | $97.0 \pm 0.6$ | $96.3 \pm 0.9$ | $32.6 \pm 4.0$ | $32.5 \pm 3.3$ |
| B3-2 | -- | $235.4 \pm 0.6$ | $227.5 \pm 0.6$ | $221.9 \pm 0.8$ | $34.6 \pm 2.7$ | $32.1 \pm 4.7$ |
| Ex B2 | -- | $509.6 \pm 2.5$ | $502.5 \pm 2.6$ | $499.5 \pm 1.8$ | $37.8 \pm 2.7$ | $37.1 \pm 3.4$ |
| TAG: |  |  |  |  |  |  |
| ALV 2178-4-1 4 | -- | $2306.0 \pm 4.6$ | $2224.3 \pm 8.5$ | $2127.5 \pm 10.0$ | $49.3 \pm 0.5$ | $47.5 \pm 0.2$ |
| 5 | -- | $1595.5 \pm 10.9$ | $1556.6 \pm 4.7$ | $1525.6 \pm 7.9$ | $45.6 \pm 0.1$ | $44.8 \pm 0.2$ |
| A4 | $7207.5 \pm 4.0$ | $628.8 \pm 1.6$ | $592.8 \pm 4.5$ | $581.5 \pm 2.2$ | $39.6 \pm 0.4$ | $38.9 \pm 0.4$ |
| A5 | $30661.6 \pm 3.1$ | $784.3 \pm 3.1$ | $766.9 \pm 2.5$ | $754.7 \pm 4.3$ | $44.9 \pm 0.5$ | $44.2 \pm 0.3$ |
| MEF: |  |  |  |  |  |  |
| ALV 2461-R13 1-1 | -- | $837.4 \pm 2.9$ | $822.4 \pm 2.4$ | $813.7 \pm 1.7$ | $26.7 \pm 0.1$ | $26.6 \pm 0.1$ |
| 1-2 | -- | $1168.1 \pm 3.1$ | $1142.5 \pm 0.7$ | $1118.2 \pm 3.0$ | $22.2 \pm 0.1$ | $22.1 \pm 0.1$ |
| 2 | -- | $740.3 \pm 2.3$ | $733.8 \pm 2.0$ | $731.1 \pm 2.3$ | $28.9 \pm 0.2$ | $28.7 \pm 0.3$ |
| 3-1 | -- | $0.8 \pm 0.01$ | $0.8 \pm 0.03$ | $0.8 \pm 0.02$ | $17.3 \pm 0.4$ | $16.9 \pm 0.3$ |
| 3-2(1-1) | -- | $0.2 \pm 0.04$ | $0.2 \pm 0.01$ | $0.2 \pm 0.01$ | $5.0 \pm 2.5$ | $8.1 \pm 5.7$ |
| 3-2(1-2) | -- | $0.3 \pm 0.01$ | $0.3 \pm 0.02$ | $0.3 \pm 0.03$ | $10.4 \pm 3.4$ | $10.4 \pm 3.3$ |
| 3-2(2) | -- | $7.7 \pm 0.01$ | $6.4 \pm 0.1$ | $5.9 \pm 0.03$ | $12.2 \pm 0.6$ | $11.0 \pm 2.5$ |
| 4-1 | -- | $9.7 \pm 0.1$ | $6.6 \pm 0.1$ | $4.1 \pm 0.2$ | $18.7 \pm 1.6$ | $18.5 \pm 1.8$ |
| 4-2 | -- | $5.9 \pm 0.1$ | $4.4 \pm 0.1$ | $3.6 \pm 1.0$ | $22.0 \pm 1.8$ | $21.4 \pm 2.3$ |
| 6 | -- | $1.8 \pm 0.01$ | $1.4 \pm 0.02$ | $1.1 \pm 0.04$ | $17.5 \pm 0.2$ | $17.3 \pm 0.2$ |
| 7 | -- | $49.7 \pm 0.3$ | $48.3 \pm 0.4$ | $46.8 \pm 0.4$ | $24.4 \pm 1.2$ | $24.7 \pm 0.1$ |
| 8-1 | -- | $118.8 \pm 0.7$ | $115.2 \pm 0.4$ | $113.3 \pm 0.3$ | $20.7 \pm 0.2$ | $20.6 \pm 0.1$ |
| 8-2 | -- | $119.9 \pm 0.5$ | $112.6 \pm 0.6$ | $107.6 \pm 0.3$ | $20.7 \pm 0.1$ | $20.5 \pm 0.04$ |

Table 6.3: Average probe permeability, core permeability, and porosity and $2 \sigma$ values for the relict spires.


Figure 6.4: Plots showing pressure profiles for select Zn -rich diffusing spire samples during permeability measurement. Colors represent sample: ALV 2187-1-1 (top) = purple, ALV 2190-14-1 = orange, J2-122-$1-\mathrm{R} 1=$ green and J2-127-1-R2 = blue. Symbol shapes represent different sample cores (solid = pressure loading, empty $=$ unloading). Changes in permeability with pressure are minor. a) ALV 2187-1-1 (top) cores. b) ALV 2190-14-1 cores. c) J2-122-1-R1 cores d) J2-127-1-R2 cores.


Figure 6.5: Plots showing pressure profiles for select black smoker chimneys during permeability measurement. Colors represent sample: J2-137-1-R1 = red, J2-213-3-R1 = navy, ALV 2179-4-1 = green and ALV 1445-3 = purple. Symbol shapes represent different sample cores (solid = pressure loading, empty $=$ unloading). Changes in permeability with pressure are minor. a) J2-137-1-R1 cores. b) J2-213-3R1 cores. c) ALV 2179-4-1 cores d) ALV 1445-3 core.


Figure 6.6: Plots showing pressure profiles for select relict spires during permeability measurement. Colors are consistent with Figure 6.3. Symbol shapes represent different sample cores (solid = pressure loading, empty $=$ unloading). Changes in permeability with pressure are minor. a) J2-136-6-R1 cores. b) J2-125-3-B1 cores. c) ALV 2944-3-S1, Pc 1 cores d) ALV 2461-R13 core.

During the initial application of pressure, many of the samples experienced a reversible decrease in permeability that can be attributed to the closure of microcracks within the cores [Walsh, 1965]. This permeability decrease may be reversible, because lowering the confining pressure would allow the cracks to re-open, therefore increasing permeability. Overall, permeability values measured during pressure unloading are consistent with the values measured during pressure loading. A few cores had permeability values from the unloading cycle that were lower than those of the loading cycle. In these cases the confining pressure may have slightly compacted the structure of the cores, thereby resulting in slightly lower permeability values during the unloading cycle.

Permeability and porosity for the Zn -rich diffusing spires, black smoker chimneys and relict spires are plotted in Figures 6.7, 6.8, and 6.9. Permeability values for the Zn -rich diffusing spires range from $\sim 2 \times 10^{-14}-5 \times 10^{-12} \mathrm{~m}^{2}$ and
porosity values are between approximately $30-45 \%$. An EPPR trend of $\alpha \sim 6$ best fits the data for these spires. The black smoker chimneys have permeability values between $\sim 3 \times 10^{-16}-3 \times 10^{-13} \mathrm{~m}^{2}$ with porosity values around $15-40 \%$. These data plot along an EPPR trending $\alpha \sim 5$. Lastly, the relict spires have permeability values from $\sim 7 \times 10^{-16}-2 \times 10^{-12} \mathrm{~m}^{2}$ and porosity values from $5-45 \%$. The EPPR trend for the relict spires is $\alpha \sim 5$. Characteristics of the cores' pore structure identified through microstructural analyses helps explain the determined EPPRs, therefore providing a better understanding of how these spires evolve.


Figure 6.7: Permeability versus porosity data for the Zn -rich actively diffusing spires. Differences in permeability as a function of porosity can be best fit by power-law relationships (black dashed line), with a power-law exponent of $\alpha \sim 6$. Symbol color denotes sample. Circles around symbols indicate that a thin section was taken from that core. All circled cores are labeled with the core number.

Figure 6.8: Permeability versus porosity data for the black smoker chimneys. Trend of EPPR (black dashed line), is $\alpha$ ~ 5. Symbol color denotes sample. Circles around symbols indicate that a thin section was taken from that core. All circled cores are labeled with the core number.

Figure 6.9: Permeability versus porosity data for the relict spires. Trend of EPPR (black dashed line), is $\alpha \sim 5$. Symbol color/ shape denotes sample. Circles around symbols indicate that a thin section was taken from that core. All circled cores are labeled with the core number.



### 6.2 Microstructural Analyses

Thin sections were taken from select spire deposits: 10 sections from the Zn -rich actively diffusing spires, 4 from the black smoker chimneys and 20 from the relict spires. Figures 6.7-6-9 show the cores from which these sections were taken.

### 6.2.1 Actively Diffusing Spires

The Zn-rich actively diffusing spires, which included white smoker chimneys, were the most permeable of the spires. The group of white smoker chimneys consisted of thin sections from sample ALV 2187-1-1 (top) cores A2, A4, and B1 and sample ALV 2187-1-2 cores C1/2 and C5. The remaining thin sections were taken from Zn -rich diffusing spire samples ALV 2190-14-1 core A1, J2-137-7-R1 core B3, J2-222-1-R1 cores A2 and Ex, and J2-128-8-R1 core Ex.

Thin sections from the white smoker sample ALV 2187-1-1 (top) are compositionally similar, being composed almost entirely of sphalerite with only trace amounts of chalcopyrite. Core A2, taken axially from the deposit sample, has the highest permeability and has a thin section characterized by broadly spaced sphalerite crystals (Figure 10a). Crystals are small to moderate sized (30$100 \mu \mathrm{~m})$ and are surrounded by wide, well-connected pore channels $(70-150 \mu \mathrm{~m})$. Core A4, also taken axially from the sample, has a similar permeability to that of A2, but with a lower porosity. Sphalerite crystals are similar in size to that of core A2 with crystals ranging from 30-100 $\mu \mathrm{m}$. Pore channels are well-connected, but
narrower $(30-60 \mu \mathrm{~m})$ than in A2 due to a tighter packing of crystals. Core B1 has a similar porosity to A2, but a lower permeability and was made radially from the deposit sample. The section for core B1 reveals layering through a prominent change in crystal size and packing. The layer oriented towards the original center of the spire deposit contains small sphalerite crystals $(10-40 \mu \mathrm{~m})$ that are widely spaced. This layer grades into an outer layer consisting of larger $(50-100 \mu \mathrm{~m})$, more closely packed crystals. Pore space within this layer appears predominantly connected through narrow channels and, in some of the more tightly packed areas, pore spaces are isolated by larger sphalerite crystals.

The thin sections from sample ALV 2187-1-2 are similar to those of ALV 2187-1-1 (top). Both ALV 2187-1-2 sections are primarily composed of sphalerite crystals with minor amounts of chalcopyrite and pyrite. The two cores have comparable permeabilities, though core C 5 has a higher porosity. Core C 5 has small to moderately sized $(30-100 \mu \mathrm{~m})$ sphalerite crystals that are loosely packed (Figure 10b). Pore space between crystals appears well-connected throughout the section. For core $\mathrm{C} 1 / 2$ the crystal size is similar to that of C 5 ; however, $\mathrm{C} 1 / 2$ has some patches of tightly packed crystals. There are no areas of very broadly spaced crystals, though there are a few prominent flow channels through the section. The tighter packing in core $\mathrm{C} 1 / 2$ is likely why it has a lower porosity than core C 5 .

The remaining Zn -rich diffusing spire cores have structures similar to those of the white smokers. Core A1 of sample ALV 2190-14-1 has the highest permeability of the remaining cores. It comprises small to moderate sized (30-100 $\mu \mathrm{m})$ sphalerite crystals with trace amounts of both chalcopyrite and


Figure 6.10: Reflected light (5x) images of Zn rich actively diffusing spires. Width of images is ~ 2.7 mm . a) ALV 2187-1-1 (top), core A2. Broadly space sphalerite crystals. b) ALV 2187-$1-2$, core C5. Loosely packed, moderate sized crystals. c) J2-127-1-R2, core B3. Small crystals with well-connected pore space. d) J2-127-1-R2, core B3. Crystals with weathered edges, likely from dissolution. e) J2-222-1-R2, core A2. Moderately packed sulfides + anhydrite. f) J2-222-1-R2, core Ex. Tightly packed sulfides + anhydrite. g) J2-128-8-R2, core Ex. Clusters of sulfide crystals have been coated in am Si.
pyrite. Sphalerite crystals are moderately packed and surrounded by wellconnected pore space. Channel width varies through different portions of the section but on average is $\sim 70-100 \mu \mathrm{~m}$. This section is similar to ALV 2187-1-1 (top) core A2, but its slightly tighter crystal packing likely makes it a little less permeable.

Core B3 from sample J2-127-1-R2 has a lower permeability and porosity than that of ALV 2190-14-1, core A1. Core B3 is dominated by wurtzite crystals ranging in size from 20-200 $\mu \mathrm{m}$ and lesser amounts of pyrite. The wurtzite crystals are well-formed and equant (Figure 10c). There are clear well-sorted areas of small crystals and of large crystals likely indicating somewhat variable environmental conditions. Wurtzite crystals are moderately to tightly packed, though pore connectivity is maintained through many narrow channels (10-40 $\mu \mathrm{m})$. A few areas within the thin section appear to have experienced some dissolution. Crystals in these areas are more irregularly shaped with rough, weathered edges (Figure 10d). Pore connectivity appears enhanced in these regions by the presence of slightly broader pore channels. The tighter crystal packing observed may account for the decrease in porosity observed between this core and ALV 2190-14-1 core A1.

While core A2 of sample J2-222-1-R2 plots very close to J2-127-1-R2 core B3, the two cores are compositionally and structurally quite different. Core A2 contains an assortment of sulfide minerals including wurtzite, pyrite, chalcopyrite, pyrrhotite, and chalcocite. The core also contains many small, tabular anhydrite crystals $(50-150 \mu \mathrm{~m})$. The section reveals layering within the
core whereby a layer containing broadly packed crystals transitions into a more tightly packed layer. Sulfide crystal sizes vary from $40-180 \mu \mathrm{~m}$ and are moderately sorted throughout the core (Figure 10e). In both the layers the anhydrite crystals create considerable roughness along the pore space edges. In the more closely packed layer these crystals restrict pore channel width. The layer of tight crystals will limit the core's permeability and porosity. Core Ex of the same sample is compositionally and texturally consistent with core A2, but has significantly tighter crystal packing within the core (Figure 10f).

Sample J2-128-8-R1 core Ex is composed primarily of wurtzite with lesser amounts of pyrite. This core shows a large range in wurtzite crystal sizes (10-300 $\mu \mathrm{m}$, with similarly sized crystals clustered together. Both large and small crystals are closely packed throughout the core, limiting pore connectivity. Additionally, most clusters of small crystals have been coated in amorphous silica, further restricting connectivity (Figure 10 g ). As with J2-127-1-R2 there are a few areas of moderate sized crystals that appear to have experienced dissolution. These crystals have weathered edges with no amorphous silica and are surrounded by highly connected pore space. This core has the lowest permeability and porosity of the Zn -rich diffusing spires and likely because of the close packed grain structure and abundance of amorphous silica clogging pore space.

### 6.2.2 Black Smoker Chimneys

Four thin sections were made from the black smoker chimney cores: J2-213-3-R1, cores A1 and D1; ALV 1445-3 core C1; and J2-137-1-R1 core D1. Both thin sections for sample J2-213-3-R1 reveal a layered structure that consists


Figure 6.11: Reflected light (5x) images of black smoker chimneys. Image width $\sim 2.7 \mathrm{~mm}$. Images show prominent layering within the samples. Densely packed chalcopyrite layers (lower images) grade into an anhydrite-rich layer (upper images) that has experienced some pore space infill by assorted sulfide crystals. a) J2-213-3-R1, core D1. b) J2-137-1-R1, core D1.
of a chalcopyrite-rich layer that gradually transitions to a more anhydrite dominated layer (Figure 11a). Core D1, the more permeable of the two, contains a
layer of moderately packed chalcopyrite crystals that range in size between 40-
$100 \mu \mathrm{~m}$. This layer is followed by a layer consisting on an assortment of sulfide
minerals including wurtzite, pyrrhotite, marcasite, and covellite, which are sized similarly to the chalcopyrite crystals. The crystals in this layer are slightly more broadly spaced than in the chalcopyrite layer. This sulfide layer grades into the anhydrite-rich layer, where anhydrite crystals are large (200-350 $\mu \mathrm{m}$ ) and interspersed with smaller sulfide crystals. Pore space is connected through narrow channels $(5-20 \mu \mathrm{~m})$ within the different layers; however, in areas where the sulfide crystal packing is somewhat dense, pores tend to be more isolated. These observations are consistent with core A1. Core A1 has a lower permeability and porosity - a difference that can be attributed to A1 having a tighter grain structure.

Core C1 for sample ALV 1445-3 also exhibits a layered structure. Densely packed chalcopyrite crystals form a chalcopyrite layer similar to that observed in J2-213-3-R1. This layer contains small ( $10-20 \mu \mathrm{~m}$ ), isolated pores and also a few thin cracks that likely formed as a result of pressure application during the measurements. The chalcopyrite layer is bordered by a slightly more porous layer of both pyrite and chalcopyrite, in addition to some anhydrite with crystals between $60-120 \mu \mathrm{~m}$. This layer then transitions into an anhydrite-rich layer characterized by $100-200 \mu \mathrm{~m}$ anhydrite crystals with traces of sulfides. The anhydrite layer is also densely packed with crystals, though narrow channels between crystals are present. Like J2-213-3-R1 core A1, the tight packing of the grain structure likely controls the permeability and porosity of ALV 1445-3 core C1.

As with the previously discussed black smoker chimney cores, J2-137-1R1 core D1 also has a layered structure consisting of a chalcopyrite-rich layer, an intermediate sulfide and anhydrite layer, and a layer that is mostly anhydrite crystals. Most of the pore space within these layers is isolated due to close crystal packing or is simply poorly connected through narrow channels between crystals. Unlike with the other cores the anhydrite layer here is relatively thin, such that the bulk of the core contains sulfides (Figure 11b). Because sulfides tend to form in clusters, crystals become close-packed, therefore restricting pore connectivity. It is the prominence of tightly packed sulfide crystals within this core which causes it to have a significantly lower permeability and porosity.

### 6.2.3 Relict Spires

The relict spire samples represent a large range of permeability and porosity values. Because of this, variations in pore structure across the different samples are to be expected. Twenty thin sections have been made from the relict spire cores in order to provide a better understanding of the processes influencing their EPPR.

The most permeable and porous of the relict spire samples is ALV 2178-41 from which two thin sections were made. The first thin section, core 4 , is the more permeable of the two and is composed primarily of pyrite and chalcopyrite. Pyrite crystals range in size between $50-120 \mu \mathrm{~m}$, while the chalcopyrite crystals are a bit smaller 30-70 $\mu \mathrm{m}$. The crystals are fairly equally distributed throughout the core; however, the crystals are very widely spaced with pore channel width averaging around $100 \mu \mathrm{~m}$ (Figure 12a). The loose crystal packing within this core
allows for high pore connectivity and thus high permeability. The second section, core A5, is composed of almost all chalcopyrite with small amounts of pyrite present. Chalcopyrite crystal size varies considerably from about 20-200 $\mu \mathrm{m}$. These crystals are poorly sorted with spacing between crystals highly variable. Pore connectivity appears high in most areas with the exception of a few tightly packed crystal clusters. The permeability is clearly high due to this pore connectivity, although not quite as high as seen in core 4 .

Sample ALV 2944-3-S1, Pc 2 also has two thin sections: core A1 and A2. Both cores are compositionally similar and are constituted largely of wurtzite and chalcopyrite with minor amounts of pyrite. Small to moderate $(20-100 \mu \mathrm{~m})$ sized chalcopyrite and wurtzite crystals are present and range from loosely packed to more moderately packed grain structure in both cores (Figure 12b). Pore connectivity appears high in both cores with abundant narrow channels ( $\sim 20 \mu \mathrm{~m}$ ) and several wider channels ( $80-100 \mu \mathrm{~m}$ ) surrounding the sulfide crystals. Core A1 has a few large patches of small, tightly packed crystals that lack well-connected channels and are therefore likely restrictive to flow. This may explain why core A1 has a lower permeability and porosity than core A2.

The one thin section from sample J2-129-1-R3 comes from core B2. This core is compositionally comparable to the two ALV 2944-3-S1, Pc 2 cores. It is composed of mostly wurtzite and chalcopyrite crystals with similar size and crystal packing as ALV 2944-3-S1, Pc 2 core A2. The availability of pore space is not consistent throughout the section. There are areas that have broadly spaced crystals, some of which may have been enhanced by dissolution, since crystals
appear irregularly shaped and with tattered edges. The core also has areas where the widely spaced crystals have been heavily coated in amorphous silica, blocking pore space around the crystals. It is possible that both dissolution and precipitation of amorphous silica have been occurring concurrently within the core. There are also a few patches of crystals that appear dissolved and then coated in amorphous silica, suggesting that in some cases dissolution may have preceded the silica precipitation (Figure 12c). Despite the crystal patches that are heavily coated in amorphous silica, the well-connected areas of crystals, including some areas of apparent dissolution, give the core a high permeability.

Sample ALV 2944-3-S1, Pc 1 has thin sections for core A1 and core A3. Core A1 is very similar to ALV 2944-3-S1, Pc 2 core A2. It is abundant in wurtzite and chalcopyrite with both small and moderate sized crystals (30-120 $\mu \mathrm{m})$ throughout the core. There are areas of very loosely packed crystals and also of moderately packed crystals. The more tightly packed crystals have narrower pore channels $(\sim 10-20 \mu \mathrm{~m})$, though the pore space still appears well connected. Core A3 has a lower permeability and porosity than core A1 and is structurally quite different. This core was the only core to have its thin section made axially, which may explain some of the variation between it and core A1. Core A3 appears to have a layered structure with a sulfide-rich layer of mainly wurtzite and chalcopyrite adjacent to a layer of anhydrite, similar to what was observed in the black smoker chimney cores (Figure 12d). The sulfide-rich layer is structurally analogous to core A1, but with more tightly packed crystals. Interestingly, the transition between the two layers has many small sulfide crystals
that have been coated with amorphous silica. This coating appears to greatly restrict pore space between the sulfide crystals and limit pore connectivity. Within the anhydrite layer, anhydrite crystals $(50-100 \mu \mathrm{~m})$ are closely packed with pore space seemingly filled by late stage sulfide crystals. This core clearly has a tightly packed grain structure due to sulfide infill and amorphous silica precipitation that are consistent with it having a lower permeability and porosity than core A1.

Three sections were made from the J2-125-3-B1 sample: core Ex, core B2-1, and core B3-1. Cores Ex and B3-1 are principally composed of wurtzite and chalcopyrite, whereas core B2-1 is nearly all chalcopyrite. Core Ex has the higher permeability of the three cores and has both large and small crystals $(40-250 \mu \mathrm{~m})$. It has patches of very tightly packed crystals and also areas with very broad spaces between crystals. The high connectivity areas of this core are likely enhancing its permeability compared with the other J2-125-3-B1 cores. Core B2-1 mainly has small to moderate sized chalcopyrite crystals that are very rounded. The crystals are closely packed, but many narrow channels are present between crystals to accommodate flow. Core B3-1 is very similar to core A1 from ALV 2944-3-S1, Pc 2 with permeability and porosity only being limited by the tightly packed wurtzite and chalcopyrite grain structure. Sample ALV 2941-6-S1 core A1, which plots adjacent to the J2-125-3-B1 cores, is compositionally and structurally consistent with core B3-1.

Core A1-2 and C2 of sample J2-136-6-R1 are both largely made of wurtzite and chalcopyrite, similar to many of the other relict spire cores. Having a higher permeability, core A1-2 is characterized by small to moderately sized


Figure 6.12: Reflected light (5x) images of thin sections from relict spire cores. Width of images is $\sim 2.7 \mathrm{~mm}$. a) ALV 2178-4-1, core 4. Loose crystal packing. b) ALV 2944-3-S1, Pc 2, core A2. Moderate crystal packing. c) J2-129-1-R3, core B2. Dissolved crystals with am Si coating. d) ALV 2944-3-S1, Pc 1, core A3. Anhydriterich layer adjacent to sulfide-rich layer. e) J2-136-6-R1, core C2. Patchy distribution of am Si. f) ALV 2461-R13, core C1-2. Sulfide grain structure with minimal am Si. g) ALV 2461-R13, core C4-2. Thick am Si coating along sulfide crystals.
crystals $(30-100 \mu \mathrm{~m})$ that range in crystal packing. Pore space around these crystals appears generally well connected, except in more densely packed areas. Core C 2 is very similar to $\mathrm{A} 1-2$, but it has a more tightly packed grain structure and also patches of amorphous silica that block pore space (Figure 12e). These features can explain why core C 2 has a lower permeability and porosity than core A1-2.

The seven thin sections made from the ALV 2461-R13 illustrate a clear progression from a high permeability and porosity structure to one that is heavily restricted by the precipitation of amorphous silica. Core C1-1, C1-2, and C8-2 have high permeability values for this sample. They have moderately packed crystal structures abundant in wurtzite and chalcopyrite with lesser amounts of pyrrhotite (Figure 12f). Pore space exists as narrow channels along crystal edges and as larger isolated pores; however, in some small patches channel connectivity is lost from amorphous silica precipitation (more so in C1-2). Cores C4-1 and C42 have initial crystal structures much more broadly spaced than the previous three cores. Small to moderate sized crystals of chalcopyrite, wurtzite, pyrite and pyrrhotite are present, although a high degree of amorphous silica precipitation has occurred. Amorphous silica forms a thick coating around most of the crystals, yet because the initial grain structure was loosely packed, pore channels and still intact (Figure 12g). Core C3-2(2) is very similar to these two cores, but has a slightly denser initial grain structure and a thicker amorphous silica coating. Lastly, C3-2(1-2), which has a very low permeability and porosity, appears to have had a more tightly packed grain structure than the previous cores that has
since been densely coated with amorphous silica. Pore space between crystals has been blocked by amorphous silica resulting in minimal pore connectivity. The combined abundance of both sulfide crystals and amorphous silica can account for the low permeability and porosity of this core.

### 6.3 Discussion

Each of the spire samples is dominated by a sulfide crystals. As observed through deposit microstructures, the tightness or packing of this grain structure strongly influences permeability and porosity. The initial mineral assemblage and grain structure of the spires will depend on the chemistry and temperature of emitted vent fluids. As fluid flow progresses, pore evolution processes will occur and change the initial grain structure. From the microstructural observations, spire deposits appear to commonly experience late stage precipitation of additional sulfides or amorphous silica, both of which can decrease the permeability and porosity of the spires. In some of the spire cores, weathered crystals with rough edges indicate that dissolution has also occurred.

The pronounced effects of both the precipitation of sulfides and also of amorphous silica on hydrothermal spires were recognized [Zhu et al., 2007]. Using techniques synonymous to those employed in this study, Zhu et al. [2007] identified two EPPRs for spire deposits having experienced precipitation. For deposits that primarily exhibited the precipitation of late stage sulfide crystals they found an $\alpha \sim 9$. Deposits having crystals that had been largely coated by amorphous silica were found to have an EPPR trend of $\alpha \sim 3$. The $\alpha$ value associated with sulfide precipitation is higher, because sulfide crystals are
naturally quite angular making them effective at blocking pore space. Conversely, amorphous silica has a lower $\alpha$ value because it precipitates as a thin, rounded coating along crystal edges that builds up over time gradually decreasing permeability.

Correlating the trends determined by Zhu et al. [2007] to the Zn -rich diffusing spire data is simply not feasible, because there is not clear evidence for either precipitation process having occurred in the cores. Only one of the cores showed precipitation of amorphous silica. Comparing data for this core with measurements from Zhu et al. [2007] places it at the high $k$ and $\phi$ end of the $\alpha \sim 3$ trend, suggesting that it may evolve along this trend. In general, cores measured in the Zhu et al. [2007] study showed a greater range in both permeability and porosity values from which interpretations of pore evolution could be made.

The black smoker chimneys did not show precipitation of amorphous silica (consistent with black smoker observations from Zhu et al. [2007]), but had a close-packed sulfide grain structure. From this grain structure and observations of sulfide infill within anhydrite dominated layers, it can be inferred that late stage precipitation of sulfides had occurred. The data clearly plot along a trend of $\alpha \sim 5$, whereas the cores from Zhu et al. [2007], which are similar compositionally, plot along the $\alpha \sim 9$ trend. The data for the lower $k$ and $\phi$ cores are comparable to those of the lower $k$ and $\phi$ cores from Zhu et al. [2007]; however, the same is not true for the high $k$ and $\phi$ cores. Zhu et al. [2007] do not detail the structure of their black smoker samples, but do state that the cores, of which there were only four, cracked during the coring process. It is possible that these cracks may have
enhanced the permeability values without significantly affecting their porosity, which is consistent with the samples.

The $\alpha$ value of 5 to some extent likely reflects the precipitation of sulfides. Black smoker chimneys should not, however, necessarily follow the $\alpha \sim 9$ trend observed by Zhu et al. [2007] in actively diffusing spires and relict spires. Black smoker chimneys have a distinct layered structure with a high abundance of anhydrite that the other spire types generally lack. The presence of significant anhydrite within black smokers should influence how the pore structure evolves. Anhydrite naturally has a close-packed grain structure, which a sulfide-rich layer does not necessarily have initially. Sulfides will infill pore spaces within the anhydrite, but the effects of this precipitation will be less pronounced compared to sulfide precipitation within an initially widely-spaced, sulfide-rich grain structure. Because of this, anhydrite should dilute or lessen the effects of sulfide precipitation on black smoker chimneys, potentially explaining why the data have an EPPR of $\alpha \sim 5$, as opposed to $\alpha \sim 9$.

For the relict spire samples, both late stage precipitation of sulfides and amorphous silica precipitation were observed in several of the cores - consistent with relict spires observations from Zhu et al. [2007]. As with the Zn-rich diffusing spires correlating the set of relict spires data to theirs is challenging, because many of the cores had permeability and porosity values higher than those of Zhu et al. [2007]. For the cores with porosities higher than $\sim 30 \%$, it is difficult to evaluate the effects of precipitation. Many of the high porosity samples did not have amorphous silica, but some samples did have tightly packed crystal
structures, which can be an indicator of later sulfide precipitation. However, given the generally high permeability values for these samples, it seems unlikely that their initial crystal structures have experienced sulfide precipitation, seeing as they exhibit well-connected pore networks.

The relict spire samples with amorphous silica present compare well with the $\alpha \sim 3$ data of Zhu et al. [2007]. It seems likely that with continued precipitation of amorphous silica these samples would evolve along the $\alpha \sim 3$ trend. Sample ALV 2461-R13 is interesting because it has a large range of permeability and porosity values. Microstructural observations from this sample suggest that precipitation of both sulfides and amorphous silica contributed to the reduction in permeability. At some point sulfide crystal packing became denser and amorphous silica coatings became thicker. This may explain why these cores plot more steeply than the others. The overall $\alpha \sim 5$ trend determined for the relict spires is probably best applicable to samples which have not experienced large amounts of change to their initial pore structures.

The group of spire deposits as a whole follows a trend whereby as the deposits change over time, they lose pore space and connectivity. Spire samples initially have high $k$ and $\phi$, but gradually evolve to lower $k$ and $\phi$ primarily as a result of precipitation. Spires will continue to evolve in this manner until they 'pinch-off'; a point at which precipitation has sufficiently clogged pore channels causing significant permeability reduction. This pinch-off is exemplified by the relict spire sample ALV 2461-R13 (Figure 13).


Figure 6.13 Reflected light (5x) images of relict spire ALV 2461-R13 core C3-2(1-2). Width of images is $\sim 2.7 \mathrm{~mm}$. Core has experience heavy precipitation of amorphous silica (dark gray, rounded coating) along sulfide crystal edges. Most of the pore space between crystals has been blocked by the amorphous silica, thus the precipitation is causing the 'pinch-off' of pore space and a large reduction in permeability.

## Chapter 7: Conclusions

The interaction of hot hydrothermal fluids with seawater controls both chemical and physical processes that can change the structure of various seafloor vents. Much work has been done to identify the composition, structure, and evolution of hydrothermal vents from a range of vent fields. Early studies conducted on deposits from the East Pacific Rise, such as those by Haymon [1983] and Goldfarb et al. [1983], analyzed the most prominent of these vent structures, the black smoker chimneys. The rapid, high-temperature focused fluid emission of these chimneys is in stark contrast to the majority of vent structures through which fluid diffuses out to the seafloor. Diffuse hydrothermal vents accommodate much of the ongoing transfer of fluids between the subsurface and seafloor, yet little information is known about the feedback between the fluid, the vent structures, and the surrounding environmental conditions [Delaney et al., 1992; Lowell et al., 1995].

The formation and evolution of all vent structures, both focused and diffuse, are closely dependent upon their physical and chemical environment. Changes in environmental conditions, such as temperature, flow rate, and the degree of mixing, can significantly impact how vents evolve over time [Tivey and McDuff, 1990; Tivey, 1995]. Conversely, changes in vent structures affect the ability of vents to transfer fluids. Many studies have collected data from vent fluids and deposit samples, and through models have been able to better understand how the fluid feedback within vents works and contributes to vent growth [Lowell et al., 1995]. However, well-constrained transport property data,
such as vent deposit permeability and porosity, are needed to improve these models. Despite its importance, systematic characterization of evolution permeability and porosity relationships (EPPRs) of vent deposits is scarce, with the exception of Zhu et al. [2007].

This study has provided the first set of systematic permeability and porosity data for many different vent deposit types from a large range of vent field locations. From these data and microstructural observations vent deposit permeability-porosity relationships and anisotropy have been identified. Deposit evolution is reflected by the different $\alpha$ values determined for each permeabilityporosity power law relationship. Evidence for pore alteration processes was observed through quantitative microstructural analyses. Some deposits show interplay between these processes, which is reflected in the $\alpha$ values determined. These results can be used to accurately evaluate vent fluid distribution in order to gain a better understanding of mineralogical and biological processes.

The results of this study show that different deposits undergo similar evolutionary processes. Microstructural observations indicate that precipitation is dominant within each of the deposit types, and that other processes, such as dissolution or compaction, are common as well. Precipitation of minerals within the initial grain structure of a deposit has large implications on the pore structure of deposits. Precipitation limits pore space and connectivity, therefore lowering a deposit's permeability. The rate and degree to which precipitation reduces permeability will depend highly on the mineral precipitating, as different minerals have varying effects on permeability. It is important to understand that the pore
structure observed in the deposits is not simply the product of one pore altering process, but rather a result of the interplay between multiple processes, whose timing or sequence is difficult to determine.

Measurements of vent deposit permeability and porosity conducted in this study can be used to identify permeability-porosity relationships and deposit characteristics, such as anisotropy. Microstructural observations provide evidence for pore evolution processes that can be used to explain the permeability-porosity relationships. Pore evolution processes are fundamental to the development of vent deposits, and recognizing key processes in this evolution reinforces the valuable role these deposits play in mass transport between the Earth's subsurface and oceans. Identified anisotropy or a susceptibility of a deposit to a given pore evolution process are characteristics that have important implications for future efforts to model and constrain fluid fluxes from these vent structures. Ongoing modeling efforts emphasize the need for detailing factors that control the evolution of vent deposits. It is clear that the growth of these deposits is dependent on the interplay between fluids and the vent deposits themselves. These interactions influence local flow rates, temperatures, and chemistry, and affect heat and mass flux through different parts of the vent deposits, and the availability of nutrients to organisms living on deposit interiors and exteriors. As such, having flow property data for vent samples is imperative for refining our understanding of hydrothermal processes.

## Appendix 1: Probe Permeability Data

## A1.1 Massive Anhydrite Data

| Sample | Site | Permeability (mD) |  |  |  |  |  | Mean (mD) | Mean ( $\mathrm{m}^{2}$ ) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | 1 | 2 | 3 |  | 4 | 5 |  |  |
| J301-3 | A1 | 2332.9 | 4325.1 | $1 \quad 1331.0$ |  | 678.7 | 3003.2 | 1938.5 | $1.94 \mathrm{E}-12$ |
|  | A2 | 624.0 | 589.9 | - 512.7 |  | 336.6 | 717.9 | 539.2 | $5.39 \mathrm{E}-13$ |
|  | B1 | 2610.0 | 542.3 | - 2205.6 |  | 2839.3 | 1575.1 | 1694.3 | $1.69 \mathrm{E}-12$ |
| J2-210-8-R2 | A1 | 78.2 | 10.1 | 181.5 |  | 47.2 | 13.4 | 39.0 | $3.90 \mathrm{E}-14$ |
|  | A2 | 11.9 | 7.6 | 54.3 |  | 29.3 | 15.8 | 18.7 | $1.87 \mathrm{E}-14$ |
|  | B1 | 34.7 | 26.9 | 23.4 |  | 3.3 | 20.3 | 17.1 | $1.71 \mathrm{E}-14$ |
|  | C1 | 35.7 | 13.0 | 30.1 |  | 45.9 | 26.2 | 27.9 | $2.79 \mathrm{E}-14$ |
| J2-216-5-R1 | A1 | 542.3 | 1189.7 | $7 \quad 606.7$ |  | 433.2 | 557.7 | 624.0 | $6.24 \mathrm{E}-13$ |
|  | A2 | 387.2 | 717.9 | - 247.2 |  | 589.9 | 261.4 | 402.8 | $4.03 \mathrm{E}-13$ |
|  | B1 | 4205.4 | 4838.9 | $9 \quad 4089.1$ |  | 3759.0 | 3088.6 | 3953.7 | $3.95 \mathrm{E}-12$ |
|  | B2 | 6776.1 | 5413.6 | 64838.9 |  | 6406.3 | 4705.0 | 5567.7 | $5.57 \mathrm{E}-12$ |
|  | C1 | 5263.8 | 4574.8 | $8 \quad 3654.9$ |  | 3553.8 | 3865.9 | 4135.2 | $4.14 \mathrm{E}-12$ |
| J2-216-14-R1 | A1 | 624.0 | 826.1 | 573.6 |  | 826.1 | 678.7 | 698.1 | $6.98 \mathrm{E}-13$ |
|  | A2 | 3003.2 | 1489.1 | 12332.9 |  | 2610.0 | 2205.6 | 2268.4 | $2.27 \mathrm{E}-12$ |
|  | B1 | 2332.9 | 2205.6 | $6 \quad 1971.4$ |  | 2839.3 | 1575.1 | 2144.6 | $2.14 \mathrm{E}-12$ |
|  | B2 | 3003.2 | 2399.3 | 32399.3 |  | 2839.3 | 3865.9 | 2855.2 | $2.86 \mathrm{E}-12$ |
|  | C1 | 6776.1 | 6776.1 | $1 \quad 5567.7$ |  | 6776.1 | 6056.7 | 6370.4 | $6.37 \mathrm{E}-12$ |
| ALV 2581-8 | A1 | 3088.6 | 2760.7 | $7 \quad 1156.8$ |  | 849.6 | 1762.1 | 1713.4 | $1.71 \mathrm{E}-12$ |
|  | A2 | 3759.0 | 5567.7 | $7 \quad 5413.6$ |  | 6229.0 | 4574.8 | 5032.7 | $5.03 \mathrm{E}-12$ |
|  | B1 | 1407.8 | 2144.6 | $6 \quad 2610.0$ |  | 1713.4 | 924.2 | 1656.6 | $1.66 \mathrm{E}-12$ |
|  | C1 | 2920.1 | 2684.3 | $3 \quad 3455.5$ |  | 3865.9 | 3759.0 | 3303.8 | $3.30 \mathrm{E}-12$ |
| $\begin{gathered} \text { MIR } 1,1 / 74 \\ \text { Sta } 2403 \end{gathered}$ | A1 | 85.1 | 80.5 | 57.5 |  | 71.9 | 74.0 | 73.1 | $7.31 \mathrm{E}-14$ |
|  | A2 | 31.9 | 44.6 | 39.9 |  | 44.6 | 48.6 | 41.5 | $4.15 \mathrm{E}-14$ |
|  | A3 | 115.9 | 13.7 | 6.6 |  | 106.5 | 18.2 | 29.0 | $2.90 \mathrm{E}-14$ |
|  | A4 | 14.9 | 19.8 | 26.2 |  | 7.4 | 9.8 | 14.1 | $1.41 \mathrm{E}-14$ |
|  | A5 | 33.7 | 112.7 | - 16.7 |  | 17.7 | 48.6 | 35.3 | $3.53 \mathrm{E}-14$ |
|  | A6 | 137.1 | 5.8 | 55.9 |  | 10.1 | 8.8 | 20.8 | $2.08 \mathrm{E}-14$ |
|  | B1 | 240.3 | 309.4 | 4181.5 |  | 186.7 | 227.2 | 224.7 | $2.25 \mathrm{E}-13$ |
|  | B2 | 849.6 | 356.0 | 498.5 |  | 678.7 | 336.6 | 509.8 | 5.10E-13 |
|  | B3 | 678.7 | 781.0 | - 641.7 |  | 781.0 | 950.5 | 759.4 | $7.59 \mathrm{E}-13$ |
|  | B4 | 387.2 | 300.8 | - 292.5 |  | 433.2 | 276.5 | 332.8 | $3.33 \mathrm{E}-13$ |
|  | B5 | 557.7 | 624.0 | - 557.7 |  | 484.7 | 433.2 | 527.3 | $5.27 \mathrm{E}-13$ |
|  | B6 | 527.3 | 458.2 | 2398.3 |  | 318.2 | 398.3 | 414.2 | $4.14 \mathrm{E}-13$ |
|  | C1 | 376.5 | 318.2 | - 346.1 |  | 589.9 | 214.8 | 350.0 | $3.50 \mathrm{E}-13$ |
|  | C2 | 898.6 | 1331.0 | 0 924.2 |  | 1005.3 | 1124.8 | 1045.6 | $1.05 \mathrm{E}-12$ |
| $\begin{gathered} \text { MIR } 1,2 / 78 \\ \text { Sta } 2417 \end{gathered}$ | A1 | 145.0 | 122.6 | 233.7 | 103.6 | 60.8 | 68.0 | 145.0 | $1.10 \mathrm{E}-13$ |
|  | A2 | 106.5 | 62.5 | 51.4 | 42.2 | 34.7 | 32.8 | 106.5 | $5.04 \mathrm{E}-14$ |
|  | A3 | 126.0 | 95.2 | 71.9 | 119.2 | 141.0 | 100.7 | 126.0 | $1.07 \mathrm{E}-13$ |
|  | A4 | 133.3 | 112.7 | 95.2 | 82.7 | 122.6 | 141.0 | 133.3 | $1.13 \mathrm{E}-13$ |
|  | A5 | 85.1 | 119.2 | 100.7 | 122.6 | 126.0 | 47.2 | 85.1 | $9.52 \mathrm{E}-14$ |
|  | A6 | 103.6 | 97.9 | 166.9 | 112.7 | 133.3 | 106.5 | 103.6 | $1.18 \mathrm{E}-13$ |
|  | B1 | 1223.5 | 1575.1 | 346.1 | 409.6 | 717.9 | 849.6 | 741.8 | $7.42 \mathrm{E}-13$ |
|  | B2 | 1189.7 | 1093.6 | 1093.68 | 873.7 | 977.5 | 1063.4 | 1043.7 | $1.04 \mathrm{E}-12$ |
| ALV 2183-7-0 | A1 | 2467.6 | 1916.9 | 9924.2 |  | 950.5 | 1034.0 | 1338.5 | $1.34 \mathrm{E}-12$ |
|  | A2 | 1666.0 | 2332.9 | 91666.0 |  | 2399.3 | 2839.3 | 2132.6 | $2.13 \mathrm{E}-12$ |
|  | B1 | 606.7 | 803.2 | - 606.7 |  | 698.1 | 678.7 | 674.9 | $6.75 \mathrm{E}-13$ |
|  | B2 | 1971.4 | 1916.9 | $9 \quad 950.5$ |  | 1368.9 | 1531.5 | 1497.5 | $1.50 \mathrm{E}-12$ |

## A1.2 Flange, Slab and Crust Data

| Sample | Site | Permeability (mD) |  |  |  |  | Mean (mD) | Mean ( $\mathrm{m}^{2}$ ) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | 1 | 2 | 3 | 4 | 5 |  |  |
| ALV 3517-R1 | A1 | 15288.9 | 12562.4 | 15724.0 | 19136.7 | 7580.9 | 13437.6 | 1.34E-11 |
|  | D1 | 4705.0 | 2537.8 | 3865.9 | 3455.5 | 2205.6 | 3230.4 | $3.23 \mathrm{E}-12$ |
|  | D2 | 20241.3 | 17105.0 | 12562.4 | 20241.3 | 18607.2 | 17493.3 | $1.75 \mathrm{E}-11$ |
|  | D3 | 1863.9 | 1762.1 | 1762.1 | 1916.9 | 2144.6 | 1884.9 | $1.88 \mathrm{E}-12$ |
|  | D4 | 2027.5 | 1223.5 | 1531.5 | 1331.0 | 1971.4 | 1583.9 | $1.58 \mathrm{E}-12$ |
| ALV 3521-R2 | A1 | 65774.9 | 46970.5 | 87080.7 | 92107.4 | 108996.3 | 76966.7 | $7.70 \mathrm{E}-11$ |
|  | A2 | 29981.0 | 26797.9 | 29981.0 | 32614.1 | 30834.2 | 29981.0 | $3.00 \mathrm{E}-11$ |
|  | B1 | 10615.9 | 9226.3 | 12214.8 | 19136.7 | 18607.2 | 13362.4 | $1.34 \mathrm{E}-11$ |
|  | B2 | 28344.8 | 41983.7 | 34496.7 | 22019.0 | 23290.0 | 29151.4 | $2.92 \mathrm{E}-11$ |
| ALV 2415-1B | A1 | 1863.9 | 1619.9 | 2205.6 | 1156.8 | 1812.3 | 1694.3 | $1.69 \mathrm{E}-12$ |
|  | B1 | 8018.5 | 14865.9 | 10322.2 | 5413.6 | 12919.9 | 9704.2 | $9.70 \mathrm{E}-12$ |
|  | C1 | 3266.9 | 10322.2 | 5263.8 | 8018.5 | 10918.0 | 6891.1 | $6.89 \mathrm{E}-12$ |
| ALV 2927-3 | A1 | 1294.2 | 1156.8 | 1368.9 | 2332.9 | 606.7 | 1237.4 | $1.24 \mathrm{E}-12$ |
|  | A2 | 37.7 | 42.2 | 109.5 | 82.7 | 71.9 | 63.6 | $6.36 \mathrm{E}-14$ |
|  | B1 | 1005.3 | 1258.4 | 2684.3 | 5567.7 | 2684.3 | 2193.3 | $2.19 \mathrm{E}-12$ |
| J2-286 | A1 | 1294.2 | 2205.6 | 2760.7 | 2268.4 | 2085.2 | 2062.0 | $2.06 \mathrm{E}-12$ |
|  | A2 | 2684.3 | 2610.0 | 4976.6 | 2610.0 | 2027.5 | 2839.3 | $2.84 \mathrm{E}-12$ |
|  | A3 | 3975.9 | 5118.2 | 3654.9 | 4325.1 | 3975.9 | 4181.9 | $4.18 \mathrm{E}-12$ |
|  | A4 | 7796.7 | 7371.2 | 8722.7 | 4574.8 | 6406.3 | 6814.2 | $6.81 \mathrm{E}-12$ |
|  | A5 | 8722.7 | 8018.5 | 7167.2 | 8246.7 | 7796.7 | 7973.7 | $7.97 \mathrm{E}-12$ |
|  | A6 | 5413.6 | 4838.9 | 3759.0 | 5118.2 | 4448.2 | 4678.6 | $4.68 \mathrm{E}-12$ |
|  | B1 | 1368.9 | 119.2 | 421.3 | 366.1 | 445.6 | 407.3 | $4.07 \mathrm{E}-13$ |
|  | B2 | 458.2 | 327.2 | 37.7 | 356.0 | 433.2 | 244.4 | $2.44 \mathrm{E}-13$ |
|  | B3 | 2144.6 | 738.4 | 1368.9 | 1619.9 | 1575.1 | 1407.8 | $1.41 \mathrm{E}-12$ |
|  | C1 | 6588.6 | 7167.2 | 6776.1 | 8246.7 | 6968.9 | 7127.1 | $7.13 \mathrm{E}-12$ |
|  | C2 | 43178.4 | 33542.2 | 40822.0 | 62185.3 | 58791.6 | 46446.3 | $4.64 \mathrm{E}-11$ |
|  | C3 | 128981.9 | 80050.4 | 115288.0 | 52549.7 | 51095.7 | 79602.4 | $7.96 \mathrm{E}-11$ |
|  | C4 | 87080.7 | 103047.9 | 67646.6 | 128981.9 | 152632.2 | 103627.8 | $1.04 \mathrm{E}-10$ |
|  | C5 | 82328.4 | 75681.7 | 144302.4 | 54045.1 | 57164.8 | 77399.8 | $7.74 \mathrm{E}-11$ |
|  | C6 | 202072.9 | 275141.7 | 180619.0 | 148408.9 | 213737.4 | 199817.6 | $2.00 \mathrm{E}-10$ |
| J2-286 | A1 | 3088.6 | 2760.7 | 1156.8 | 849.6 | 1762.1 | 1713.4 | $1.71 \mathrm{E}-12$ |
|  | A2 | 3759.0 | 5567.7 | 5413.6 | 6229.0 | 4574.8 | 5032.7 | $5.03 \mathrm{E}-12$ |
|  | B1 | 1407.8 | 2144.6 | 2610.0 | 1713.4 | 924.2 | 1656.6 | $1.66 \mathrm{E}-12$ |
|  | C1 | 2920.1 | 2684.3 | 3455.5 | 3865.9 | 3759.0 | 3303.8 | $3.30 \mathrm{E}-12$ |
| ALV 2608-3-3 | A1 | 18092.3 | 22019.0 | 23290.0 | 19136.7 | 15288.9 | 19352.7 | $1.94 \mathrm{E}-11$ |
|  | A2 | 1916.9 | 2205.6 | 2027.5 | 1713.4 | 1971.4 | 1960.4 | $1.96 \mathrm{E}-12$ |
|  | A3 | 4205.4 | 4705.0 | 4976.6 | 4205.4 | 4574.8 | 4523.7 | $4.52 \mathrm{E}-12$ |
|  | B1 | 8018.5 | 11876.9 | 13287.6 | 18607.2 | 17105.0 | 13213.2 | $1.32 \mathrm{E}-11$ |
|  | B2 | 4705.0 | 10322.2 | 7796.7 | 6776.1 | 8722.7 | 7412.6 | $7.41 \mathrm{E}-12$ |
|  | B3 | 3088.6 | 2537.8 | 1762.1 | 1575.1 | 2027.5 | 2132.6 | $2.13 \mathrm{E}-12$ |
|  | C1 | 1124.8 | 1124.8 | 1034.0 | 950.5 | 1034.0 | 1051.5 | $1.05 \mathrm{E}-12$ |
|  | C2 | 12919.9 | 9758.8 | 6056.7 | 8971.0 | 4705.0 | 7973.7 | $7.97 \mathrm{E}-12$ |
|  | C3 | 4205.4 | 4574.8 | 4574.8 | 4448.2 | 3003.2 | 4112.1 | $4.11 \mathrm{E}-12$ |
|  | C4 | 409.6 | 336.6 | 409.6 | 318.2 | 376.5 | 368.2 | $3.68 \mathrm{E}-13$ |
|  | D1 | 1916.9 | 2144.6 | 2027.5 | 1575.1 | 1489.1 | 1812.3 | $1.81 \mathrm{E}-12$ |
|  | D2 | 1713.4 | 2467.6 | 924.2 | 803.2 | 873.7 | 1223.5 | $1.22 \mathrm{E}-12$ |
|  | D3 | 717.9 | 589.9 | 458.2 | 346.1 | 606.7 | 527.3 | $5.27 \mathrm{E}-13$ |
|  | D4 | 2144.6 | 2144.6 | -- | -- | -- | 2144.6 | $2.14 \mathrm{E}-12$ |


| $\begin{aligned} & \text { ALV 2608-4-1 } \\ & \quad \text { Pc } 1 \end{aligned}$ | A1 | 1063.4 | 698.1 | 803.2 | 624.0 | 977.5 | 816.8 | $8.17 \mathrm{E}-13$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | A2 | 10615.9 | 10322.2 | 9758.8 | 11876.9 | 12562.4 | 10979.4 | $1.10 \mathrm{E}-11$ |
|  | A3 | 17591.7 | 14865.9 | 19136.7 | 18092.3 | 15288.9 | 16914.1 | $1.69 \mathrm{E}-11$ |
|  | A4 | 3553.8 | 2085.2 | 3455.5 | 3359.9 | 3553.8 | 3141.1 | $3.14 \mathrm{E}-12$ |
|  | B1 | 33542.2 | 52549.7 | 44407.2 | 31711.6 | 25335.5 | 36283.8 | $3.63 \mathrm{E}-11$ |
|  | C1 | 512.7 | 458.2 | 484.7 | 433.2 | 292.5 | 428.4 | $4.28 \mathrm{E}-13$ |
|  | C2 | 2332.9 | 2144.6 | 1971.4 | 2205.6 | 1812.3 | 2085.2 | $2.09 \mathrm{E}-12$ |
|  | C3 | 1034.0 | 1093.6 | 924.2 | 1531.5 | 1531.5 | 1196.4 | $1.20 \mathrm{E}-12$ |
| $\begin{gathered} \text { ALV 2608-4-1 } \\ \text { Pc } 2 \end{gathered}$ | A1 | 8722.7 | 26056.4 | 32614.1 | 5726.1 | 12919.9 | 14054.6 | $1.41 \mathrm{E}-11$ |
|  | A2 | 292.5 | 641.7 | 498.5 | 421.3 | 717.9 | 490.2 | $4.90 \mathrm{E}-13$ |
|  | A3 | 2205.6 | 2332.9 | 2399.3 | 2467.6 | 2205.6 | 2319.9 | $2.32 \mathrm{E}-12$ |
|  | B1 | 717.9 | 873.7 | 803.2 | 803.2 | 781.0 | 794.2 | 7.94E-13 |
|  | B2 | 3003.2 | 1093.6 | 1407.8 | 1368.9 | 977.5 | 1439.8 | $1.44 \mathrm{E}-12$ |
|  | B3 | 1156.8 | 1489.1 | 1407.8 | 924.2 | 1005.3 | 1176.4 | $1.18 \mathrm{E}-12$ |
| JAS 177-2-1 | A1 | 512.7 | 471.3 | 458.2 | 398.3 | 573.6 | 479.3 | $4.79 \mathrm{E}-13$ |
|  | A2 | 3176.5 | 3759.0 | 3865.9 | 4448.2 | 4838.9 | 3975.9 | $3.98 \mathrm{E}-12$ |
|  | A3 | 1575.1 | 1863.9 | 1331.0 | 950.5 | 1407.8 | 1392.1 | $1.39 \mathrm{E}-12$ |
|  | B1 | 2144.6 | 3865.9 | 1916.9 | 2537.8 | 3759.0 | 2729.9 | $2.73 \mathrm{E}-12$ |
|  | B2 | 14454.6 | 5118.2 | 7796.7 | 12919.9 | 10036.5 | 9435.7 | $9.44 \mathrm{E}-12$ |
|  | B3 | 6229.0 | 22645.6 | 11228.7 | 5726.1 | 8481.4 | 9488.8 | $9.49 \mathrm{E}-12$ |
|  | C1 | 4574.8 | 6968.9 | 4325.1 | 6588.6 | 8246.7 | 5955.5 | $5.96 \mathrm{E}-12$ |
|  | C2 | 54045.1 | 89558.8 | 60464.6 | 87080.7 | 65774.9 | 69963.2 | $7.00 \mathrm{E}-11$ |
|  | C3 | 4574.8 | 4574.8 | 6406.3 | 5263.8 | 6056.7 | 5323.2 | $5.32 \mathrm{E}-12$ |
| ALV 2179-1-1 | A1 | 15288.9 | 23952.8 | 16631.7 | 6588.6 | 12919.9 | 13897.7 | $1.39 \mathrm{E}-11$ |
|  | A2 | 4838.9 | 5118.2 | 8481.4 | 3975.9 | 4705.0 | 5234.4 | $5.23 \mathrm{E}-12$ |
|  | A3 | 18607.2 | 11228.7 | 12214.8 | 11876.9 | 213737.4 | 23030.1 | $2.30 \mathrm{E}-11$ |
|  | B1 | 738.4 | 458.2 | 924.2 | 376.5 | 542.3 | 576.8 | $5.77 \mathrm{E}-13$ |
|  | B2 | 1156.8 | 1575.1 | 1005.3 | 2467.6 | 1489.1 | 1464.2 | $1.46 \mathrm{E}-12$ |

## A1.3 Zn-Rich Actively Diffusing Spire Data

| Sample | Site | Permeability (mD) |  |  |  |  | Mean (mD) | Mean (m2) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | 1 | 2 | 3 | 4 | 5 |  |  |
| $\begin{aligned} & \text { ALV 2187-1-1 } \\ & \text { top } \end{aligned}$ | A1 | 2920.1 | 2839.3 | 3003.2 | 2268.4 | 2537.8 | 2699.4 | $2.70 \mathrm{E}-12$ |
|  | A2 | 8971.0 | 8246.7 | 10322.2 | 7167.2 | 7371.2 | 8339.8 | $8.34 \mathrm{E}-12$ |
|  | A3 | 6968.9 | 8246.7 | 10036.5 | 6776.1 | 5118.2 | 7248.1 | $7.25 \mathrm{E}-12$ |
|  | A4 | 7796.7 | 8971.0 | 8722.7 | 8971.0 | 7580.9 | 8386.7 | $8.39 \mathrm{E}-12$ |
|  | B1 | 24634.4 | 28344.8 | 28344.8 | 26056.4 | 30834.2 | 27560.5 | $2.76 \mathrm{E}-11$ |
|  | B2 | 22019.0 | 22019.0 | 19136.7 | 29151.4 | 23952.8 | 23030.1 | $2.30 \mathrm{E}-11$ |
|  | B3 | 12562.4 | 22019.0 | 5118.2 | 23952.8 | 18607.2 | 14454.6 | $1.45 \mathrm{E}-11$ |
| $\begin{aligned} & \text { ALV 2187-1-1 } \\ & \text { bottom } \end{aligned}$ | A1 | 180619.0 | 202072.9 | 121942.9 | 69571.7 | 180619.0 | 141099.3 | $1.41 \mathrm{E}-10$ |
|  | A2 | 28344.8 | 28344.8 | 23952.8 | 26797.9 | 22645.6 | 25910.6 | $2.59 \mathrm{E}-11$ |
|  | A3 | 10322.2 | 12919.9 | 8018.5 | 8018.5 | 6968.9 | 9021.5 | $9.02 \mathrm{E}-12$ |
|  | A4 | 7580.9 | 3759.0 | 2920.1 | 3359.9 | 4574.8 | 4181.9 | $4.18 \mathrm{E}-12$ |
|  | B1 | 25335.5 | 39692.5 | 26056.4 | 32614.1 | 35478.4 | 31357.7 | $3.14 \mathrm{E}-11$ |
|  | B2 | 14454.6 | 16631.7 | 11876.9 | 8481.4 | 17591.7 | 13362.4 | $1.34 \mathrm{E}-11$ |
|  | B3 | 3176.5 | 6968.9 | 4976.6 | 7167.2 | 7580.9 | 5694.1 | $5.69 \mathrm{E}-12$ |
|  | C1 | 3553.8 | 2760.7 | 2684.3 | 2839.3 | 2399.3 | 2823.4 | $2.82 \mathrm{E}-12$ |
|  | C2 | 1063.4 | 1093.6 | 1223.5 | 1005.3 | 1331.0 | 1137.5 | $1.14 \mathrm{E}-12$ |
|  | C3 | 738.4 | 641.7 | 606.7 | 624.0 | 606.7 | 641.7 | $6.42 \mathrm{E}-13$ |
|  | C4 | 18092.3 | 23952.8 | 25335.5 | 20817.4 | 16631.7 | 20700.9 | $2.07 \mathrm{E}-11$ |


| ALV 2187-1-2 | A1 | 77835.4 | 48307.2 | 80050.4 | 28344.8 | 20817.4 | 44657.1 | $4.47 \mathrm{E}-11$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | A2 | 19136.7 | 21409.8 | 57164.8 | 60464.6 | 35478.4 | 34690.8 | $3.47 \mathrm{E}-11$ |
|  | B1 | 4838.9 | 4574.8 | 5118.2 | 4448.2 | 15288.9 | 5989.1 | $5.99 \mathrm{E}-12$ |
|  | B2 | 4838.9 | 7167.2 | 7580.9 | 5263.8 | 2332.9 | 5032.7 | $5.03 \mathrm{E}-12$ |
|  | B3 | 2268.4 | 1971.4 | 2537.8 | 2399.3 | 2399.3 | 2306.9 | $2.31 \mathrm{E}-12$ |
|  | C1 | 291024.0 | 291024.0 | 232508.7 | 245930.1 | 307823.1 | 272070.8 | $2.72 \mathrm{E}-10$ |
|  | C2 | 213737.4 | 84671.2 | 495982.9 | 219819.8 | 239125.2 | 216149.9 | $2.16 \mathrm{E}-10$ |
|  | C3 | 2537.8 | 2205.6 | 1916.9 | 3176.5 | 2205.6 | 2372.5 | $2.37 \mathrm{E}-12$ |
|  | C4 | 2610.0 | 2537.8 | 2920.1 | 3088.6 | 3455.5 | 2903.7 | $2.90 \mathrm{E}-12$ |
|  | C5 | 8018.5 | 6776.1 | 8246.7 | 8246.7 | 6776.1 | 7580.9 | 7.58E-12 |
| ALV 2190-14-1 | A1 | 3975.9 | 4205.4 | 3865.9 | 4205.4 | 4089.1 | 4066.2 | $4.07 \mathrm{E}-12$ |
|  | A2 | 10322.2 | 10615.9 | 13287.6 | 8246.7 | 7580.9 | 9813.8 | $9.81 \mathrm{E}-12$ |
|  | A3 | 4089.1 | 3865.9 | 4574.8 | 3359.9 | 3654.9 | 3887.7 | $3.89 \mathrm{E}-12$ |
|  | A4 | 9758.8 | 14054.6 | 14454.6 | 15288.9 | 14454.6 | 13437.6 | $1.34 \mathrm{E}-11$ |
|  | A5 | 39692.5 | 49681.9 | 100196.6 | 73587.6 | 316582.9 | 85626.9 | $8.56 \mathrm{E}-11$ |
|  | A6 | 919524.4 | 1704746.3 | 539542.1 | 799157.2 | 364265.9 | 755544.1 | $7.56 \mathrm{E}-10$ |
|  | B1 | 11876.9 | 11548.2 | 10918.0 | 12214.8 | 19136.7 | 12847.6 | $1.28 \mathrm{E}-11$ |
|  | B2 | 8722.7 | 7796.7 | 7580.9 | 10322.2 | -- | 8541.1 | $8.54 \mathrm{E}-12$ |
| $\begin{gathered} \mathrm{J} 2-137-7-\mathrm{R} 1 \\ \mathrm{Pc} 1 \end{gathered}$ | A1 | 275141.7 | 495982.9 | 156975.7 | 385292.8 | 232508.7 | 286165.4 | $2.86 \mathrm{E}-10$ |
|  | A2 | 374631.8 | 694546.3 | 354186.7 | 603629.1 | 570686.7 | 501581.2 | $5.02 \mathrm{E}-10$ |
|  | B1 | 170761.9 | 325591.9 | 354186.7 | 219819.8 | 180619.0 | 239125.2 | $2.39 \mathrm{E}-10$ |
|  | B2 | 374631.8 | 226075.2 | 316582.9 | 275141.7 | 191045.0 | 269034.2 | $2.69 \mathrm{E}-10$ |
|  | C1 | 65774.9 | 144302.4 | 92107.4 | 161442.7 | 191045.0 | 121942.9 | $1.22 \mathrm{E}-10$ |
|  | C2 | 919524.4 | 972603.1 | 468915.2 | 734638.4 | 419130.7 | 664053.8 | $6.64 \mathrm{E}-10$ |
|  | D1 | 226075.2 | 291024.0 | 175621.3 | 455940.4 | 307823.1 | 276690.1 | $2.77 \mathrm{E}-10$ |
|  | E1 | 156975.7 | 275141.7 | 374631.8 | 334857.4 | 115288.0 | 228627.0 | $2.29 \mathrm{E}-10$ |
| J2-139-2-R1 | A1 | 140309.6 | 63954.9 | 175621.3 | 87080.7 | 58791.6 | 95797.7 | $9.58 \mathrm{E}-11$ |
|  | B1 | 115288.0 | 115288.0 | 132652.4 | 191045.0 | 121942.9 | 132652.4 | $1.33 \mathrm{E}-10$ |
| J2-128-8-R1 | A1 | 11228.7 | 11876.9 | 29151.4 | 46970.5 | 67646.6 | 26203.1 | $2.62 \mathrm{E}-11$ |
|  | A2 | 20241.3 | 12562.4 | 11228.7 | 9758.8 | 16631.7 | 13589.2 | $1.36 \mathrm{E}-11$ |
|  | B1 | 125413.0 | 175621.3 | 396257.1 | 144302.4 | 170761.9 | 184719.3 | $1.85 \mathrm{E}-10$ |

## A1.4 Black Smoker Chimney Data

| Sample | Site | Permeability (mD) |  |  |  |  | Mean (mD) | Mean ( $\mathrm{m}^{2}$ ) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | 1 | 2 | 3 | 4 | 5 |  |  |
| ALV 2462-R2 | A1 | 122.6 | 82.7 | 85.1 | 71.9 | 27.7 | 70.3 | $7.03 \mathrm{E}-14$ |
|  | A2 | 69.9 | 47.2 | 14.1 | 19.8 | 20.3 | 28.5 | $2.85 \mathrm{E}-14$ |
|  | A3 | 141.0 | 33.7 | 25.5 | 82.7 | 55.9 | 56.2 | $5.62 \mathrm{E}-14$ |
|  | A4 | 103.6 | 103.6 | 71.9 | 100.7 | 64.3 | 87.0 | $8.70 \mathrm{E}-14$ |
|  | B1 | 24.1 | 16.3 | 16.3 | 15.8 | 15.4 | 17.3 | $1.73 \mathrm{E}-14$ |
|  | B2 | 25.5 | 17.7 | 38.8 | 15.8 | 27.7 | 23.8 | $2.38 \mathrm{E}-14$ |
| ALV 1445-3 | A1 | 109.5 | 186.7 | 254.2 | 292.5 | 292.5 | 213.6 | $2.14 \mathrm{E}-13$ |
|  | B1 | 346.1 | 318.2 | 327.2 | 214.8 | 153.4 | 260.0 | $2.60 \mathrm{E}-13$ |
|  | C1 | 192.0 | 119.2 | 60.8 | 109.5 | 240.3 | 129.6 | $1.30 \mathrm{E}-13$ |
| ALV 2179-4-1 | A1 | 1447.9 | 950.5 | 1063.4 | 1063.4 | 1812.3 | 1230.4 | $1.23 \mathrm{E}-12$ |
|  | B1 | 398.3 | 376.5 | 421.3 | 208.9 | 214.8 | 309.4 | $3.09 \mathrm{E}-13$ |
|  | C1 | 2085.2 | 3759.0 | 3975.9 | 3975.9 | 5413.6 | 3675.5 | $3.68 \mathrm{E}-12$ |


| J2-213-2-R1 | A1 | 803.2 | 1005.3 | 698.1 | 678.7 | 898.6 | 807.7 | 8.08E-13 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | B1 | 80.5 | 176.5 | 62.5 | 162.3 | 186.7 | 121.9 | $1.22 \mathrm{E}-13$ |
|  | B2 | 458.2 | 458.2 | 318.2 | 512.7 | -- | 430.2 | $4.30 \mathrm{E}-13$ |
|  | B3 | 327.2 | 300.8 | 261.4 | -- | -- | 295.2 | $2.95 \mathrm{E}-13$ |
|  | C1 | 27.7 | 34.7 | 35.7 | -- | -- | 32.5 | $3.25 \mathrm{E}-14$ |
|  | D1 | 1407.8 | 1713.4 | 3553.8 | 1916.9 | 1063.4 | 1772.1 | 1.77E-12 |
|  | D2 | 3759.0 | 4976.6 | 4205.4 | 3088.6 | 2760.7 | 3675.5 | $3.68 \mathrm{E}-12$ |
| J2-213-3-R1 | A1 | 13665.7 | 10322.2 | 18092.3 | 10036.5 | 10322.2 | 12146.5 | $1.21 \mathrm{E}-11$ |
|  | B1 | 4705.0 | 4089.1 | 5726.1 | 4838.9 | 7167.2 | 5205.1 | $5.21 \mathrm{E}-12$ |
|  | B2 | 3553.8 | 4838.9 | 4574.8 | 4705.0 | 4838.9 | 4473.2 | $4.47 \mathrm{E}-12$ |
|  | B3 | 5726.1 | 4838.9 | 4205.4 | 5567.7 | 4574.8 | 4948.7 | $4.95 \mathrm{E}-12$ |
|  | C1 | 471.3 | 433.2 | 387.2 | 421.3 | 356.0 | 411.9 | $4.12 \mathrm{E}-13$ |
|  | C2 | 781.0 | 950.5 | 898.6 | 1063.4 | 1189.7 | 966.6 | $9.67 \mathrm{E}-13$ |
|  | C3 | 112.7 | 149.2 | 137.1 | -- | -- | 132.1 | $1.32 \mathrm{E}-13$ |
|  | D1 | 22.8 | 27.7 | -- | -- | -- | 25.1 | $2.51 \mathrm{E}-14$ |
|  | E1 | 5889.1 | 5413.6 | 7167.2 | 4089.1 | -- | 5528.8 | $5.53 \mathrm{E}-12$ |
|  | F1 | 624.0 | 1124.8 | 803.2 | 366.1 | 527.3 | 641.7 | $6.42 \mathrm{E}-13$ |
| J2-137-3-R1 | A1 | 13665.7 | 8722.7 | 9758.8 | 6229.0 | 5726.1 | 8386.7 | $8.39 \mathrm{E}-12$ |
|  | A2 | 13287.6 | 9758.8 | 5118.2 | 5413.6 | 5118.2 | 7127.1 | $7.13 \mathrm{E}-12$ |
|  | A3 | 44407.2 | 108996.3 | 87080.7 | 71551.5 | 118568.8 | 81409.5 | 8.14E-11 |
|  | A4 | 22645.6 | 21409.8 | 15724.0 | -- | -- | 19681.3 | $1.97 \mathrm{E}-11$ |
|  | B1 | 387.2 | 214.8 | 197.5 | -- | -- | 254.2 | $2.54 \mathrm{E}-13$ |
|  | C1 | 7371.2 | 3266.9 | 6229.0 | 4325.1 | 4205.4 | 4866.1 | $4.87 \mathrm{E}-12$ |
|  | C2 | 3759.0 | 3975.9 | 4205.4 | 4205.4 | 3266.9 | 3865.9 | $3.87 \mathrm{E}-12$ |
|  | D1 | 115.9 | 100.7 | 119.2 | -- | -- | 111.6 | $1.12 \mathrm{E}-13$ |
|  | D2 | 48.6 | 87.5 | 54.3 | -- | -- | 61.3 | $6.13 \mathrm{E}-14$ |
|  | D3 | 268.9 | 203.1 | 203.1 | -- | -- | 223.0 | $2.23 \mathrm{E}-13$ |
|  | D4 | 126.0 | 62.5 | 92.6 | -- | -- | 90.0 | $9.00 \mathrm{E}-14$ |
|  | E1 | 45670.9 | 36488.0 | 48307.2 | -- | -- | 43178.4 | $4.32 \mathrm{E}-11$ |
|  | E2 | 2839.3 | 2610.0 | 2467.6 | -- | -- | 2634.6 | $2.63 \mathrm{E}-12$ |
|  | E3 | 115.9 | 171.6 | 112.7 | 106.5 | 103.6 | 119.8 | $1.20 \mathrm{E}-13$ |
|  | E4 | 8971.0 | 11228.7 | 10322.2 | -- | -- | 10130.9 | $1.01 \mathrm{E}-11$ |
|  | F1 | 2144.6 | 1863.9 | 2268.4 | -- | -- | 2085.2 | $2.09 \mathrm{E}-12$ |
|  | F2 | 192.0 | 171.6 | 220.9 | -- | -- | 193.8 | $1.94 \mathrm{E}-13$ |
|  | G1 | 300.8 | 336.6 | 233.7 | -- | -- | 287.1 | $2.87 \mathrm{E}-13$ |
|  | G2 | 57.5 | 41.0 | 51.4 | -- | -- | 49.5 | $4.95 \mathrm{E}-14$ |
|  | H1 | 327.2 | 387.2 | 366.1 | -- | -- | 359.3 | $3.59 \mathrm{E}-13$ |
|  | H2 | 1971.4 | 803.2 | 398.3 | 484.7 | -- | 743.6 | 7.44E-13 |

## A1.5 Relict Spire Data

| Sample | Site | Permeability (mD) |  |  |  |  | Mean (mD) | Mean ( $\mathrm{m}^{\mathbf{2}}$ ) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | 1 | 2 | 3 | 4 | 5 |  |  |
| J2-129-1-R3 | A1 | 292.5 | 122.6 | 376.5 | 32.8 | 129.6 | 141.8 | $1.42 \mathrm{E}-13$ |
|  | A2 | 157.8 | 233.7 | 203.1 | 214.8 | 192.0 | 198.6 | $1.99 \mathrm{E}-13$ |
|  | A3 | 268.9 | 186.7 | 227.2 | 149.2 | 220.9 | 206.5 | $2.07 \mathrm{E}-13$ |
|  | B1 | 276.5 | 157.8 | 284.4 | 284.4 | 186.7 | 231.1 | $2.31 \mathrm{E}-13$ |
|  | B2 | 137.1 | 74.0 | 103.6 | 74.0 | 68.0 | 88.0 | $8.80 \mathrm{E}-14$ |
|  | B3 | 51.4 | 68.0 | 47.2 | 38.8 | 25.5 | 43.9 | $4.39 \mathrm{E}-14$ |
|  | C1 | 144302.4 | 112098.0 | 239125.2 | 170761.9 | 267528.6 | 177603.5 | $1.78 \mathrm{E}-10$ |
|  | C2 | 121942.9 | 92107.4 | 166036.9 | 89558.8 | 100196.6 | 110846.9 | $1.11 \mathrm{E}-10$ |
| J2-136-6-R1 | A1 | 181.5 | 171.6 | 133.3 | 129.6 | 153.4 | 152.5 | $1.53 \mathrm{E}-13$ |
|  | A2 | 122.6 | 95.2 | 92.6 | 106.5 | 80.5 | 98.5 | $9.85 \mathrm{E}-14$ |
|  | A3 | 2085.2 | 3088.6 | 1916.9 | 1971.4 | 1156.8 | 1949.4 | $1.95 \mathrm{E}-12$ |
|  | A4 | 1005.3 | 498.5 | 233.7 | 409.6 | 660.0 | 501.3 | $5.01 \mathrm{E}-13$ |
|  | B1 | 162.3 | 126.0 | 133.3 | 57.5 | 74.0 | 103.0 | $1.03 \mathrm{E}-13$ |
|  | B2 | 247.2 | 300.8 | 261.4 | 220.9 | 318.2 | 267.4 | $2.67 \mathrm{E}-13$ |
|  | B3 | 6776.1 | 4205.4 | 5889.1 | 8018.5 | 7167.2 | 6264.1 | $6.26 \mathrm{E}-12$ |
|  | C1 | 387.2 | 421.3 | 606.7 | 471.3 | 542.3 | 479.3 | $4.79 \mathrm{E}-13$ |
|  | C2 | 589.9 | 220.9 | 498.5 | 641.7 | 433.2 | 448.1 | $4.48 \mathrm{E}-13$ |
|  | C3 | 119.2 | 62.5 | 74.0 | 64.3 | 69.9 | 75.6 | $7.56 \mathrm{E}-14$ |
| $\begin{gathered} \text { ALV 2944-3-S1 } \\ \text { Pc } 1 \end{gathered}$ | A1 | 63954.9 | 49681.9 | 43178.4 | 48307.2 | 52549.7 | 51095.7 | $5.11 \mathrm{E}-11$ |
|  | A2 | 44407.2 | 57164.8 | 49681.9 | 45670.9 | 54045.1 | 49961.5 | $5.00 \mathrm{E}-11$ |
|  | A3 | 44407.2 | 43178.4 | 36488.0 | 52549.7 | 34496.7 | 41748.7 | $4.17 \mathrm{E}-11$ |
|  | A4 | 37526.3 | 57164.8 | 46970.5 | 25335.5 | 46970.5 | 41282.8 | $4.13 \mathrm{E}-11$ |
|  | B1 | 292.5 | 214.8 | 176.5 | 247.2 | 227.2 | 228.5 | $2.28 \mathrm{E}-13$ |
|  | B2 | 781.0 | 268.9 | 1156.8 | 898.6 | 387.2 | 610.1 | $6.10 \mathrm{E}-13$ |
|  | B3 | 29151.4 | 25335.5 | 24634.4 | 29151.4 | 27560.5 | 27100.4 | $2.71 \mathrm{E}-11$ |
| ALV 2941-6-S1 | A1 | 144302.4 | 191045.0 | 75681.7 | 152632.2 | 103047.9 | 126828.6 | $1.27 \mathrm{E}-10$ |
|  | A2 | 38594.2 | 22019.0 | 63954.9 | 29981.0 | 34496.7 | 35478.4 | $3.55 \mathrm{E}-11$ |
|  | B1 | 21409.8 | 16171.5 | 41983.7 | 15724.0 | 21409.8 | 21773.3 | $2.18 \mathrm{E}-11$ |
|  | B2 | 25335.5 | 26797.9 | 18607.2 | 19681.3 | 27560.5 | 23290.0 | $2.33 \mathrm{E}-11$ |
|  | C1 | 11228.7 | 12919.9 | 4574.8 | 5726.1 | 5726.1 | 7371.2 | $7.37 \mathrm{E}-12$ |
| J2-125-3-B1 | A1 | 4325.1 | 6229.0 | 5889.1 | 3266.9 | 6776.1 | 5118.2 | $5.12 \mathrm{E}-12$ |
|  | A2 | 220.9 | 284.4 | 309.4 | 261.4 | 149.2 | 237.7 | $2.38 \mathrm{E}-13$ |
|  | A3 | 1223.5 | 1575.1 | 1447.9 | 1863.9 | 1124.8 | 1423.7 | $1.42 \mathrm{E}-12$ |
|  | A4 | 1812.3 | 1863.9 | 1812.3 | 717.9 | 1447.9 | 1447.9 | $1.45 \mathrm{E}-12$ |
|  | B1 | 18607.2 | 18092.3 | 11228.7 | 8018.5 | 14865.9 | 13513.2 | $1.35 \mathrm{E}-11$ |
|  | B2 | 3003.2 | 4838.9 | 3654.9 | 5118.2 | 3553.8 | 3953.7 | $3.95 \mathrm{E}-12$ |
|  | B3 | 1223.5 | 1368.9 | 1189.7 | 1034.0 | 1156.8 | 1189.7 | $1.19 \mathrm{E}-12$ |
|  | C1 | 115.9 | 145.0 | 97.9 | 176.5 | 90.0 | 121.2 | $1.21 \mathrm{E}-13$ |
|  | C2 | 87.5 | 55.9 | 28.5 | 45.9 | 35.7 | 46.9 | $4.69 \mathrm{E}-14$ |
|  | C3 | 87.5 | 49.9 | 64.3 | 71.9 | 49.9 | 63.2 | $6.32 \mathrm{E}-14$ |
| ALV 2178-4-1 | A1 | 14.9 | 15.8 | 119.2 | 122.6 | 78.2 | 48.6 | $4.86 \mathrm{E}-14$ |
|  | A2 | 1619.9 | 1368.9 | 1005.3 | 1034.0 | 1093.6 | 1203.1 | $1.20 \mathrm{E}-12$ |
|  | A3 | 557.7 | 1005.3 | 826.1 | 898.6 | 977.5 | 835.4 | $8.35 \mathrm{E}-13$ |
|  | A4 | 7580.9 | 8018.5 | 6056.7 | 7796.7 | 6776.1 | 7207.5 | $7.21 \mathrm{E}-12$ |
|  | A5 | 30834.2 | 20817.4 | 39692.5 | 29151.4 | 36488.0 | 30661.6 | $3.07 \mathrm{E}-11$ |
|  | B1 | 18607.2 | 11228.7 | 15724.0 | 17591.7 | 21409.8 | 16538.6 | $1.65 \mathrm{E}-11$ |
|  | B2 | 3654.9 | 3865.9 | 3654.9 | 2839.3 | 3553.8 | 3494.5 | $3.49 \mathrm{E}-12$ |
|  | B3 | 498.5 | 356.0 | 247.2 | 268.9 | 247.2 | 311.1 | $3.11 \mathrm{E}-13$ |
|  | B4 | 2027.5 | 1447.9 | 924.2 | 1531.5 | 1189.7 | 1376.6 | $1.38 \mathrm{E}-12$ |

## Appendix 2: Core Permeability Data

## A2.1 Massive Anhydrite Data



| length: 1.142 cm width: 2.533 cm | 250 | 227.757 | 227.715 | 227.538 | 228.202 | 227.604 | 227.76 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 200 | 230.622 | 230.223 | 230.201 | 230.386 | 230.315 | 230.35 |
|  | 150 | 233.764 | 233.847 | 234.236 | 234.051 | 233.975 | 233.97 |
|  | 100 | 239.299 | 239.613 | 239.999 | 240.511 | 240.297 | 239.94 |
|  | 70 | 246.349 | 246.666 | 247.122 | 247.279 | 247.447 | 246.97 |
| $\begin{gathered} \text { J2-216-14-R1 } \\ \text { Core A } \end{gathered}$ | 70 | 0.365 | 0.324 | 0.343 | 0.349 | 0.352 | 0.35 |
|  | 100 | 0.329 | 0.334 | 0.336 | 0.340 | 0.333 | 0.33 |
|  | 150 | 0.334 | 0.326 | 0.325 | 0.344 | 0.326 | 0.33 |
|  | 200 | 0.310 | 0.333 | 0.333 | 0.334 | 0.325 | 0.33 |
|  | 250 | 0.332 | 0.338 | 0.331 | 0.315 | 0.324 | 0.33 |
|  | 300 | 0.331 | 0.323 | 0.316 | 0.323 | 0.329 | 0.32 |
|  | 350 | 0.329 | 0.330 | 0.322 | 0.330 | 0.323 | 0.33 |
|  | 400 | 0.323 | 0.321 | 0.322 | 0.315 | 0.322 | 0.32 |
|  | 350 | 0.306 | 0.323 | 0.323 | 0.324 | 0.324 | 0.32 |
|  | 300 | 0.314 | 0.323 | 0.333 | 0.322 | 0.324 | 0.32 |
|  | 250 | 0.332 | 0.341 | 0.333 | 0.324 | 0.332 | 0.33 |
|  | 200 | 0.349 | 0.340 | 0.331 | 0.350 | 0.335 | 0.34 |
|  | 150 | 0.332 | 0.342 | 0.350 | 0.335 | 0.350 | 0.34 |
| length: 1.841 cm width: 2.530 cm | 100 | 0.341 | 0.351 | 0.349 | 0.351 | 0.342 | 0.35 |
|  | 70 | 0.343 | 0.351 | 0.353 | 0.359 | 0.353 | 0.35 |
| $\begin{gathered} \mathrm{J} 2-216-14-\mathrm{R} 1 \\ \text { Core B } \end{gathered}$ | 70 | 0.380 | 0.392 | 0.371 | 0.390 | 0.376 | 0.38 |
|  | 100 | 0.389 | 0.378 | 0.388 | 0.361 | 0.373 | 0.38 |
|  | 150 | 0.371 | 0.377 | 0.386 | 0.386 | 0.364 | 0.38 |
|  | 200 | 0.365 | 0.391 | 0.371 | 0.355 | 0.344 | 0.37 |
|  | 250 | 0.331 | 0.347 | 0.330 | 0.354 | 0.336 | 0.34 |
|  | 300 | 0.343 | 0.342 | 0.336 | 0.341 | 0.333 | 0.34 |
|  | 350 | 0.324 | 0.324 | 0.316 | 0.332 | 0.336 | 0.33 |
|  | 400 | 0.328 | 0.335 | 0.329 | 0.326 | 0.330 | 0.33 |
|  | 350 | 0.342 | 0.335 | 0.326 | 0.345 | 0.347 | 0.34 |
|  | 300 | 0.341 | 0.348 | 0.348 | 0.348 | 0.345 | 0.35 |
|  | 250 | 0.365 | 0.355 | 0.363 | 0.363 | 0.355 | 0.36 |
|  | 200 | 0.374 | 0.365 | 0.367 | 0.375 | 0.366 | 0.37 |
|  | 150 | 0.373 | 0.362 | 0.384 | 0.374 | 0.375 | 0.37 |
| length: 2.243 cm width: 2.536 cm | 100 | 0.386 | 0.386 | 0.396 | 0.380 | 0.380 | 0.39 |
|  | 70 | 0.398 | 0.388 | 0.370 | 0.389 | 0.388 | 0.39 |
| J301-3 <br> Core A <br> length: 3.205 cm width: 2.546 cm | 70 | 44.519 | 44.719 | 44.820 | 44.789 | 44.714 | 44.71 |
|  | 100 | 43.244 | 42.910 | 42.915 | 42.990 | 42.865 | 42.99 |
|  | 150 | 40.907 | 40.837 | 40.942 | 40.751 | 40.928 | 40.87 |
|  | 200 | 39.711 | 39.921 | 39.959 | 39.651 | 39.784 | 39.81 |
|  | 250 | 38.926 | 38.955 | 39.029 | 38.942 | 38.913 | 38.95 |
|  | 300 | 38.632 | 38.308 | 38.730 | 38.592 | 38.377 | 38.53 |
|  | 350 | 38.093 | 38.303 | 37.967 | 38.138 | 37.996 | 38.10 |
|  | 400 | 37.959 | 37.832 | 37.962 | 38.224 | 37.860 | 37.97 |
|  | 350 | 37.888 | 37.916 | 38.050 | 37.933 | 37.989 | 37.96 |
|  | 300 | 38.558 | 38.089 | 38.118 | 38.249 | 38.469 | 38.30 |
|  | 250 | 38.303 | 38.524 | 38.362 | 38.304 | 38.361 | 38.37 |
|  | 200 | 38.765 | 38.720 | 38.765 | 38.939 | 39.177 | 38.87 |
|  | 150 | 39.377 | 39.376 | 39.497 | 39.450 | 39.254 | 39.39 |
|  | 100 | 40.232 | 40.297 | 40.411 | 40.267 | 40.207 | 40.28 |
|  | 70 | 41.201 | 41.325 | 41.416 | 41.435 | 41.547 | 41.39 |
| J301-3 <br> Core B | 70 | 3.089 | 3.066 | 3.026 | 3.024 | 2.968 | 3.03 |
|  | 100 | 2.150 | 2.146 | 2.135 | 2.109 | 2.115 | 2.13 |
|  | 150 | 1.455 | 1.449 | 1.430 | 1.458 | 1.424 | 1.44 |
|  | 200 | 1.136 | 1.110 | 1.139 | 1.101 | 1.106 | 1.12 |
|  | 250 | 0.896 | 0.865 | 0.861 | 0.831 | 0.844 | 0.86 |
|  | 300 | 0.749 | 0.744 | 0.702 | 0.730 | 0.710 | 0.73 |
|  | 350 | 0.647 | 0.614 | 0.630 | 0.598 | 0.613 | 0.62 |
|  | 400 | 0.585 | 0.540 | 0.519 | 0.570 | 0.543 | 0.55 |
|  | 350 | 0.536 | 0.545 | 0.545 | 0.526 | 0.552 | 0.54 |
|  | 300 | 0.603 | 0.581 | 0.617 | 0.594 | 0.609 | 0.60 |
|  | 250 | 0.616 | 0.639 | 0.619 | 0.639 | 0.620 | 0.63 |
|  | 200 | 0.711 | 0.710 | 0.694 | 0.700 | 0.699 | 0.70 |
|  | 150 | 0.785 | 0.830 | 0.843 | 0.815 | 0.839 | 0.82 |
| length: 5.196 cm width: 2.530 cm | 100 | 1.090 | 1.078 | 1.113 | 1.089 | 1.113 | 1.10 |
|  | 70 | 1.503 | 1.495 | 1.531 | 1.506 | 1.520 | 1.51 |




## A2.2 Flange, Slab and Crust Data



| length: 1.549 cm width: 2.506 cm | 100 | 1302.646 | 1303.591 | 1304.986 | 1316.100 | 1312.943 | 1308.04 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 70 | 1338.587 | 1341.667 | 1335.947 | 1336.308 | 1335.224 | 1337.54 |
| ALV 3521-R2 Core Ex 2 | 70 | 1013.176 | 1012.378 | 1022.284 | 1011.949 | 1017.531 | 1015.46 |
|  | 100 | 1006.855 | 1004.269 | 1001.139 | 1004.820 | 1000.574 | 1003.53 |
|  | 150 | 974.762 | 981.290 | 981.475 | 979.969 | 984.304 | 980.35 |
|  | 200 | 971.354 | 969.261 | 966.181 | 971.116 | 970.228 | 969.63 |
|  | 250 | 958.035 | 958.814 | 960.185 | 961.126 | 960.605 | 959.75 |
|  | 300 | 954.072 | 957.787 | 956.848 | 950.985 | 953.788 | 954.69 |
|  | 350 | 953.184 | 954.635 | 952.004 | 954.165 | 954.336 | 953.66 |
|  | 400 | 954.974 | 952.534 | 953.422 | 954.716 | 953.275 | 953.78 |
|  | 350 | 952.618 | 955.944 | 952.412 | 955.546 | 956.446 | 954.59 |
|  | 300 | 950.459 | 949.099 | 954.063 | 954.352 | 956.801 | 952.95 |
|  | 250 | 954.661 | 955.023 | 955.972 | 952.357 | 956.625 | 954.93 |
|  | 200 | 955.066 | 955.984 | 956.841 | 954.333 | 956.917 | 955.83 |
|  | 150 | 971.874 | 970.979 | 963.867 | 967.613 | 968.145 | 968.49 |
| length: 1.014 cm <br> width: 2.517 cm | 100 | 974.243 | 972.702 | 977.035 | 979.075 | 979.594 | 976.53 |
|  | 70 | 984.051 | 984.049 | 988.846 | 984.498 | 989.657 | 986.22 |
| $\begin{aligned} & \text { ALV 2415-1B } \\ & \text { Core A1 } \end{aligned}$ | 70 | 68.571 | 68.423 | 68.871 | 68.180 | 68.119 | 68.43 |
|  | 100 | 63.621 | 63.567 | 63.575 | 63.585 | 63.501 | 63.57 |
|  | 150 | 57.862 | 58.024 | 57.944 | 57.796 | 57.926 | 57.91 |
|  | 200 | 56.608 | 56.730 | 56.434 | 56.485 | 56.399 | 56.53 |
|  | 250 | 55.865 | 55.821 | 55.957 | 55.731 | 55.799 | 55.83 |
|  | 300 | 55.432 | 55.202 | 55.374 | 55.306 | 55.675 | 55.40 |
|  | 350 | 55.240 | 55.095 | 55.064 | 54.881 | 54.913 | 55.04 |
|  | 400 | 54.756 | 54.597 | 54.604 | 54.411 | 55.120 | 54.70 |
|  | 350 | 54.742 | 54.755 | 54.877 | 54.928 | 54.797 | 54.82 |
|  | 300 | 54.926 | 55.479 | 55.094 | 55.313 | 55.169 | 55.20 |
|  | 250 | 55.701 | 55.558 | 55.400 | 55.541 | 55.690 | 55.58 |
|  | 200 | 55.686 | 55.673 | 55.728 | 55.755 | 55.872 | 55.74 |
|  | 150 | 56.401 | 56.215 | 56.317 | 56.300 | 56.195 | 56.29 |
| length: 2.084 cm width: 2.480 cm | 100 | 57.992 | 57.659 | 57.656 | 57.678 | 57.699 | 57.74 |
|  | 70 | 62.883 | 62.735 | 62.855 | 62.786 | 62.703 | 62.79 |
| ALV 2415-1BCore B1 | 70 | 1731.305 | 1727.945 | 1726.755 | 1720.326 | 1726.610 | 1726.58 |
|  | 100 | 1661.130 | 1662.277 | 1659.746 | 1656.780 | 1661.611 | 1660.31 |
|  | 150 | 1621.342 | 1621.304 | 1620.372 | 1619.277 | 1617.850 | 1620.03 |
|  | 200 | 1598.285 | 1593.126 | 1595.238 | 1594.572 | 1591.969 | 1594.64 |
|  | 250 | 1572.853 | 1573.635 | 1570.254 | 1570.673 | 1573.056 | 1572.09 |
|  | 300 | 1551.429 | 1555.157 | 1550.990 | 1551.450 | 1552.245 | 1552.25 |
|  | 350 | 1534.874 | 1533.078 | 1535.014 | 1534.458 | 1533.041 | 1534.09 |
|  | 400 | 1517.890 | 1513.684 | 1512.048 | 1515.330 | 1512.577 | 1514.30 |
|  | 350 | 1517.067 | 1519.251 | 1517.333 | 1518.957 | 1517.523 | 1518.03 |
|  | 300 | 1522.546 | 1525.321 | 1525.635 | 1525.267 | 1521.020 | 1523.96 |
|  | 250 | 1531.960 | 1532.006 | 1533.799 | 1531.695 | 1527.531 | 1531.40 |
|  | 200 | 1545.121 | 1547.289 | 1545.072 | 1547.390 | 1545.990 | 1546.17 |
|  | 150 | 1559.269 | 1561.854 | 1562.734 | 1562.881 | 1562.708 | 1561.89 |
| length: 5.838 cm width: 2.490 cm | 100 | 1577.023 | 1576.552 | 1580.638 | 1578.798 | 1580.188 | 1578.64 |
|  | 70 | 1595.937 | 1593.497 | 1596.318 | 1592.738 | 1598.929 | 1595.48 |
| $\begin{aligned} & \text { ALV 2415-1B } \\ & \text { Core } 1 \end{aligned}$ | 70 | 66.8 | -- | -- | -- | -- | 66.8 |
|  | 100 | 59.3 | -- | -- | -- | -- | 59.3 |
|  | 150 | 52.0 | -- | -- | -- | -- | 52.0 |
|  | 200 | 46.1 | -- | -- | -- | -- | 46.1 |
|  | 250 | 41.6 | -- | -- | -- | -- | 41.6 |
|  | 300 | 39.2 | -- | -- | -- | -- | 39.2 |
|  | 350 | 37.3 | -- | -- | -- | -- | 37.3 |
|  | 400 | 36.0 | -- | -- | -- | -- | 36.0 |
| ALV 2415-1B Core 2 | 70 | 11.8 | -- | -- | -- | -- | 11.8 |
|  | 100 | 9.8 | -- | -- | -- | -- | 9.8 |
|  | 150 | 7.1 | -- | -- | -- | -- | 7.1 |
|  | 200 | 5.3 | -- | -- | -- | -- | 5.3 |
|  | 250 | 4.5 | -- | -- | -- | -- | 4.5 |
|  | 300 | 3.5 | -- | -- | -- | -- | 3.5 |
|  | 350 | 2.9 | -- | -- | -- | -- | 2.9 |
|  | 400 | 2.5 | -- | -- | -- | -- | 2.5 |
| $\begin{gathered} \text { ALV } 2727-3 \\ \text { Core B1 } \end{gathered}$ | 70 | 1998.995 | 1995.040 | 1989.075 | 1994.736 | 1993.444 | 1994.26 |
|  | 100 | 1973.888 | 1979.961 | 1979.924 | 1981.849 | 1977.613 | 1978.65 |
|  | 150 | 1956.234 | 1955.954 | 1948.347 | 1956.622 | 1958.593 | 1955.15 |


| length: 2.658 cm width: 2.485 cm | 200 | 1931.790 | 1935.533 | 1928.452 | 1934.961 | 1922.350 | 1930.61 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 250 | 1902.452 | 1904.822 | 1902.014 | 1898.109 | 1906.368 | 1902.75 |
|  | 300 | 1887.248 | 1894.701 | 1896.360 | 1881.084 | 1885.969 | 1889.06 |
|  | 350 | 1872.134 | 1874.619 | 1871.287 | 1872.729 | 1875.124 | 1873.18 |
|  | 400 | 1862.190 | 1866.476 | 1866.930 | 1862.102 | 1859.724 | 1863.48 |
|  | 350 | 1860.102 | 1877.249 | 1860.662 | 1863.453 | 1867.762 | 1865.83 |
|  | 300 | 1867.857 | 1867.034 | 1863.636 | 1870.709 | 1873.502 | 1868.54 |
|  | 250 | 1874.716 | 1876.572 | 1878.228 | 1883.035 | 1881.153 | 1878.74 |
|  | 200 | 1892.253 | 1893.593 | 1895.523 | 1897.051 | 1887.724 | 1893.23 |
|  | 150 | 1905.288 | 1906.861 | 1908.943 | 1906.681 | 1905.719 | 1906.70 |
|  | 100 | 1929.958 | 1920.503 | 1927.836 | 1929.570 | 1932.740 | 1928.12 |
|  | 70 | 1951.330 | 1943.875 | 1939.239 | 1947.789 | 1950.196 | 1946.48 |
| ALV 2927-3 Core 1 | 70 | 2631.6 | -- | -- | -- | -- | 2631.6 |
|  | 100 | 2601.1 | -- | -- | -- | -- | 2601.1 |
|  | 150 | 2295.1 | -- | -- | -- | -- | 2295.1 |
|  | 200 | 1950.2 | -- | -- | -- | -- | 1950.2 |
|  | 250 | 1816.9 | -- | -- | -- | -- | 1816.9 |
|  | 300 | 1715.5 | -- | -- | -- | -- | 1715.5 |
|  | 350 | 1653.1 | -- | -- | -- | -- | 1653.1 |
|  | 400 | 1613.7 | -- | -- | -- | -- | 1613.7 |
| $\begin{gathered} \text { ALV 2927-3 } \\ \text { Core } 2 \end{gathered}$ | 70 | 182.4 | -- | -- | -- | -- | 182.4 |
|  | 100 | 172.8 | -- | -- | -- | -- | 172.8 |
|  | 150 | 161.8 | -- | -- | -- | -- | 161.8 |
|  | 200 | 155.6 | -- | -- | -- | -- | 155.6 |
|  | 250 | 151.5 | -- | -- | -- | -- | 151.5 |
|  | 300 | 148.7 | -- | -- | -- | -- | 148.7 |
|  | 350 | 146.1 | -- | -- | -- | -- | 146.1 |
|  | 400 | 145.3 | -- | -- | -- | -- | 145.3 |
| ALV 2927-3 <br> Core 3 | 70 | 1081.4 | -- | -- | -- | -- | 1081.4 |
|  | 100 | 1062.7 | -- | -- | -- | -- | 1062.7 |
|  | 150 | 1010.0 | -- | -- | -- | -- | 1010.0 |
|  | 200 | 994.2 | -- | -- | -- | -- | 994.2 |
|  | 250 | 966.1 | -- | -- | -- | -- | 966.1 |
|  | 300 | 968.9 | -- | -- | -- | -- | 968.9 |
|  | 350 | 955.6 | -- | -- | -- | -- | 955.6 |
|  | 400 | 948.7 | -- | -- | -- | -- | 948.7 |
| J2-286Core A1 | 70 | 5.003 | 4.930 | 4.920 | 4.856 | 4.852 | 4.91 |
|  | 100 | 3.726 | 3.716 | 3.686 | 3.681 | 3.696 | 3.70 |
|  | 150 | 2.215 | 2.204 | 2.183 | 2.145 | 2.144 | 2.18 |
|  | 200 | 1.375 | 1.346 | 1.322 | 1.295 | 1.274 | 1.32 |
|  | 250 | 0.879 | 0.845 | 0.839 | 0.824 | 0.811 | 0.84 |
|  | 300 | 0.573 | 0.553 | 0.545 | 0.543 | 0.532 | 0.55 |
|  | 350 | 0.372 | 0.364 | 0.363 | 0.347 | 0.347 | 0.36 |
|  | 400 | 0.230 | 0.205 | 0.222 | 0.211 | 0.228 | 0.22 |
|  | 350 | 0.251 | 0.243 | 0.252 | 0.245 | 0.242 | 0.25 |
|  | 300 | 0.309 | 0.313 | 0.318 | 0.317 | 0.314 | 0.31 |
|  | 250 | 0.373 | 0.386 | 0.390 | 0.388 | 0.384 | 0.38 |
|  | 200 | 0.505 | 0.502 | 0.526 | 0.525 | 0.530 | 0.52 |
|  | 150 | 0.770 | 0.795 | 0.799 | 0.808 | 0.816 | 0.80 |
|  | 100 | 1.432 | 1.469 | 1.459 | 1.493 | 1.504 | 1.47 |
|  | 70 | 2.059 | 2.117 | 2.142 | 2.152 | 2.197 | 2.13 |
| $\begin{gathered} \text { J2-286 } \\ \text { Core A3 } \end{gathered}$ | 70 | 439.838 | 438.832 | 438.354 | 438.718 | 437.868 | 438.72 |
|  | 100 | 368.859 | 368.805 | 370.881 | 371.413 | 369.797 | 369.95 |
|  | 150 | 322.158 | 322.574 | 317.979 | 319.613 | 319.194 | 320.30 |
|  | 200 | 279.324 | 280.227 | 279.470 | 279.646 | 278.441 | 279.42 |
|  | 250 | 255.609 | 254.216 | 251.834 | 254.330 | 251.736 | 253.54 |
|  | 300 | 233.886 | 233.420 | 232.187 | 233.161 | 232.871 | 233.10 |
|  | 350 | 213.140 | 211.920 | 212.417 | 211.869 | 214.545 | 212.78 |
|  | 400 | 196.954 | 196.538 | 195.662 | 195.956 | 194.888 | 196.00 |
|  | 350 | 203.158 | 203.336 | 202.614 | 202.411 | 203.359 | 202.98 |
|  | 300 | 215.031 | 215.563 | 214.968 | 215.484 | 214.721 | 215.15 |
|  | 250 | 230.742 | 232.038 | 230.680 | 231.189 | 230.912 | 231.11 |
|  | 200 | 249.850 | 248.087 | 249.670 | 249.983 | 250.200 | 249.56 |
|  | 150 | 274.082 | 272.449 | 273.810 | 273.139 | 274.139 | 273.52 |
| length: 5.280 cm width: 2.525 cm | 100 | 302.767 | 302.318 | 304.170 | 307.748 | 303.678 | 304.13 |
|  | 70 | 325.777 | 326.183 | 326.831 | 328.068 | 331.561 | 327.68 |


| $\begin{gathered} \text { J2-286 } \\ \text { Core C2-1 } \end{gathered}$ | 70 | 0.917 | 0.896 | 0.865 | 0.862 | 0.916 | 0.89 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 100 | 0.910 | 0.906 | 0.869 | 0.864 | 0.882 | 0.89 |
|  | 150 | 0.853 | 0.808 | 0.839 | 0.882 | 0.889 | 0.85 |
|  | 200 | 0.807 | 0.858 | 0.851 | 0.834 | 0.818 | 0.83 |
|  | 250 | 0.880 | 0.827 | 0.859 | 0.839 | 0.811 | 0.84 |
|  | 300 | 0.814 | 0.820 | 0.829 | 0.793 | 0.817 | 0.81 |
|  | 350 | 0.783 | 0.778 | 0.754 | 0.786 | 0.791 | 0.78 |
|  | 400 | 0.766 | 0.769 | 0.789 | 0.774 | 0.743 | 0.77 |
|  | 350 | 0.786 | 0.753 | 0.797 | 0.785 | 0.761 | 0.78 |
|  | 300 | 0.789 | 0.819 | 0.793 | 0.803 | 0.822 | 0.81 |
|  | 250 | 0.781 | 0.773 | 0.824 | 0.811 | 0.836 | 0.80 |
|  | 200 | 0.823 | 0.777 | 0.789 | 0.810 | 0.825 | 0.80 |
|  | 150 | 0.855 | 0.847 | 0.833 | 0.783 | 0.852 | 0.83 |
| length: 1.960 cm width: 2.520 cm | 100 | 0.878 | 0.828 | 0.826 | 0.844 | 0.830 | 0.84 |
|  | 70 | 0.915 | 0.863 | 0.865 | 0.902 | 0.890 | 0.89 |
| $\begin{gathered} \text { J2-286 } \\ \text { Core C2-2 } \end{gathered}$ | 70 | 1.376 | 1.316 | 1.320 | 1.341 | 1.398 | 1.35 |
|  | 100 | 1.385 | 1.308 | 1.357 | 1.322 | 1.335 | 1.34 |
|  | 150 | 1.258 | 1.268 | 1.254 | 1.228 | 1.278 | 1.26 |
|  | 200 | 1.243 | 1.280 | 1.242 | 1.282 | 1.250 | 1.26 |
|  | 250 | 1.211 | 1.218 | 1.202 | 1.209 | 1.227 | 1.21 |
|  | 300 | 1.198 | 1.172 | 1.216 | 1.178 | 1.198 | 1.19 |
|  | 350 | 1.215 | 1.167 | 1.222 | 1.170 | 1.189 | 1.19 |
|  | 400 | 1.140 | 1.179 | 1.218 | 1.139 | 1.199 | 1.17 |
|  | 350 | 1.236 | 1.204 | 1.220 | 1.244 | 1.222 | 1.23 |
|  | 300 | 1.252 | 1.228 | 1.224 | 1.254 | 1.258 | 1.24 |
|  | 250 | 1.355 | 1.308 | 1.356 | 1.335 | 1.368 | 1.34 |
|  | 200 | 1.326 | 1.353 | 1.358 | 1.366 | 1.389 | 1.36 |
|  | 150 | 1.394 | 1.360 | 1.357 | 1.342 | 1.356 | 1.36 |
| length: 3.137 cm width: 2.530 cm | 100 | 1.372 | 1.391 | 1.390 | 1.341 | 1.354 | 1.37 |
|  | 70 | 1.391 | 1.372 | 1.392 | 1.380 | 1.396 | 1.39 |
| $\begin{gathered} \text { J2-286 } \\ \text { Core C3-1 } \end{gathered}$ | 70 | 185.271 | 185.176 | 184.815 | 184.534 | 185.821 | 185.12 |
|  | 100 | 175.694 | 176.695 | 177.123 | 177.993 | 177.633 | 177.03 |
|  | 150 | 170.851 | 170.930 | 171.989 | 171.571 | 171.891 | 171.45 |
|  | 200 | 163.143 | 162.622 | 162.664 | 162.877 | 163.017 | 162.86 |
|  | 250 | 152.794 | 153.467 | 153.257 | 152.485 | 152.454 | 152.89 |
|  | 300 | 145.672 | 145.911 | 145.358 | 147.131 | 145.947 | 146.00 |
|  | 350 | 141.506 | 142.266 | 141.051 | 140.182 | 141.250 | 141.25 |
|  | 400 | 139.136 | 138.557 | 138.575 | 137.348 | 137.831 | 138.29 |
|  | 350 | 139.164 | 138.605 | 138.142 | 139.521 | 138.503 | 138.79 |
|  | 300 | 140.467 | 141.136 | 140.875 | 139.761 | 141.474 | 140.74 |
|  | 250 | 143.716 | 142.803 | 142.978 | 143.609 | 143.316 | 143.28 |
|  | 200 | 148.219 | 147.587 | 149.246 | 148.690 | 148.001 | 148.35 |
|  | 150 | 157.364 | 157.677 | 158.338 | 156.804 | 157.850 | 157.61 |
| length: 1.320 cm width: 2.517 cm | 100 | 172.424 | 173.504 | 171.991 | 174.596 | 173.699 | 173.24 |
|  | 70 | 179.532 | 180.064 | 178.833 | 179.368 | 179.895 | 179.54 |
| $\begin{gathered} \text { J2-286 } \\ \text { Core C3-2 } \end{gathered}$ | 70 | 57.872 | 58.003 | 58.184 | 57.857 | 58.008 | 57.98 |
|  | 100 | 54.417 | 54.526 | 54.147 | 54.344 | 54.146 | 54.32 |
|  | 150 | 51.471 | 51.068 | 51.131 | 51.274 | 51.009 | 51.19 |
|  | 200 | 49.982 | 49.641 | 50.102 | 50.041 | 49.565 | 49.87 |
|  | 250 | 48.993 | 49.000 | 49.006 | 48.933 | 48.843 | 48.95 |
|  | 300 | 48.733 | 48.869 | 48.828 | 48.586 | 48.664 | 48.74 |
|  | 350 | 48.353 | 48.706 | 48.272 | 48.409 | 48.049 | 48.36 |
|  | 400 | 48.325 | 48.453 | 48.062 | 48.091 | 48.338 | 48.25 |
|  | 350 | 48.485 | 48.209 | 48.405 | 48.261 | 48.379 | 48.35 |
|  | 300 | 48.409 | 48.490 | 48.171 | 48.187 | 48.509 | 48.35 |
|  | 250 | 48.651 | 48.433 | 48.269 | 48.533 | 48.379 | 48.45 |
|  | 200 | 48.541 | 48.464 | 48.362 | 48.587 | 48.740 | 48.54 |
| length: 2.166 cm width: 2.530 cm | 150 | 48.790 | 48.988 | 48.960 | 48.943 | 48.764 | 48.89 |
|  | 100 | 49.496 | 49.097 | 49.544 | 49.443 | 49.586 | 49.43 |
|  | 70 | 50.230 | 50.391 | 50.858 | 50.457 | 50.515 | 50.49 |
| $\begin{gathered} \mathrm{J} 2-286 \\ \text { Core C4 } \end{gathered}$ | 70 | 22.544 | 22.713 | 22.845 | 22.920 | 22.557 | 22.72 |
|  | 100 | 21.136 | 21.159 | 21.172 | 21.031 | 21.045 | 21.11 |
|  | 150 | 19.479 | 19.738 | 19.524 | 19.533 | 19.430 | 19.54 |
|  | 200 | 18.737 | 18.829 | 18.635 | 18.738 | 18.771 | 18.74 |
|  | 250 | 18.234 | 18.215 | 18.267 | 18.179 | 18.180 | 18.21 |
|  | 300 | 17.665 | 17.724 | 17.912 | 17.633 | 17.797 | 17.75 |


|  | 350 | 17.335 | 17.286 | 17.239 | 17.272 | 17.215 | 17.27 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 400 | 16.973 | 16.987 | 16.979 | 17.125 | 17.064 | 17.03 |
|  | 350 | 17.282 | 17.309 | 17.257 | 17.237 | 17.223 | 17.26 |
|  | 300 | 17.364 | 17.347 | 17.315 | 17.350 | 17.350 | 17.35 |
|  | 250 | 17.551 | 17.581 | 17.392 | 17.502 | 17.455 | 17.50 |
|  | 200 | 17.513 | 17.635 | 17.770 | 17.751 | 17.659 | 17.67 |
|  | 150 | 18.164 | 18.117 | 18.011 | 18.113 | 18.096 | 18.10 |
| length: 2.242 cm | 100 | 18.457 | 18.419 | 18.517 | 18.750 | 18.636 | 18.56 |
| width: 2.534 cm | 70 | 19.155 | 19.191 | 19.282 | 19.340 | 19.185 | 19.23 |
|  | 70 | 2775.457 | 2753.390 | 2750.327 | 2750.043 | 2751.290 | 2756.08 |
| ALV 2608-3-3 | 100 | 2731.965 | 2741.756 | 2748.601 | 2744.529 | 2743.458 | 2742.06 |
| Core B2 | 150 | 2737.428 | 2746.376 | 2745.481 | 2743.131 | 2740.446 | 2742.57 |
|  | 200 | 2739.747 | 2742.163 | 2747.917 | 2754.997 | 2738.889 | 2744.74 |
|  | 250 | 2738.028 | 2745.671 | 2739.575 | 2742.961 | 2739.434 | 2741.13 |
|  | 300 | 2731.569 | 2730.212 | 2744.337 | 2739.639 | 2739.495 | 2737.05 |
|  | 350 | 2742.793 | 2747.108 | 2741.122 | 2746.627 | 2745.311 | 2744.59 |
|  | 400 | 2743.230 | 2749.732 | 2743.436 | 2747.952 | 2746.149 | 2746.10 |
|  | 350 | 2751.994 | 2753.703 | 2750.475 | 2750.154 | 2745.641 | 2750.39 |
|  | 300 | 2748.038 | 2748.589 | 2752.861 | 2748.051 | 2753.687 | 2750.24 |
|  | 250 | 2752.909 | 2757.370 | 2756.729 | 2754.918 | 2752.377 | 2754.86 |
|  | 200 | 2752.351 | 2752.451 | 2759.157 | 2759.967 | 2753.436 | 2755.47 |
|  | 150 | 2760.591 | 2762.581 | 2754.112 | 2754.220 | 2762.390 | 2758.78 |
| length: 3.127 cm | 100 | 2765.729 | 2766.051 | 2765.017 | 2764.154 | 2768.542 | 2765.90 |
| width: 2.500 cm | 70 | 2774.483 | 2772.973 | 2767.481 | 2773.497 | 2771.457 | 2771.98 |
|  | 70 | 1463.134 | 1458.002 | 1452.060 | 1453.063 | 1455.805 | 1456.41 |
| ALV 2608-3-3 | 100 | 1445.711 | 1447.873 | 1450.619 | 1446.636 | 1448.560 | 1447.88 |
| Core C1 | 150 | 1442.011 | 1447.001 | 1448.876 | 1445.477 | 1447.688 | 1446.21 |
|  | 200 | 1441.697 | 1444.850 | 1449.962 | 1444.873 | 1444.956 | 1445.27 |
|  | 250 | 1443.511 | 1446.972 | 1446.920 | 1447.800 | 1446.114 | 1446.26 |
|  | 300 | 1447.954 | 1446.510 | 1446.235 | 1445.371 | 1449.270 | 1447.07 |
|  | 350 | 1438.684 | 1433.104 | 1441.381 | 1435.182 | 1436.227 | 1436.91 |
|  | 400 | 1435.555 | 1439.549 | 1436.163 | 1435.085 | 1436.500 | 1436.57 |
|  | 350 | 1438.654 | 1441.551 | 1435.478 | 1439.569 | 1440.906 | 1439.23 |
|  | 300 | 1443.227 | 1443.663 | 1441.331 | 1439.345 | 1443.221 | 1442.16 |
|  | 250 | 1444.857 | 1444.730 | 1445.591 | 1448.387 | 1446.511 | 1446.01 |
|  | 200 | 1442.243 | 1451.638 | 1451.012 | 1449.685 | 1446.193 | 1448.15 |
|  | 150 | 1450.345 | 1443.050 | 1448.496 | 1450.353 | 1455.514 | 1449.55 |
| length: 2.092 cm | 100 | 1457.652 | 1456.714 | 1456.325 | 1456.357 | 1454.295 | 1456.27 |
| width: 2.496 cm | 70 | 1456.331 | 1458.631 | 1458.978 | 1453.156 | 1459.316 | 1457.28 |
|  | 70 | 896.556 | 896.976 | 898.147 | 898.717 | 899.208 | 897.92 |
| ALV 2608-4-1 | 100 | 867.383 | 866.905 | 866.397 | 865.442 | 865.227 | 866.27 |
| Pc 1 | 150 | 817.947 | 818.801 | 817.289 | 816.501 | 817.137 | 817.53 |
| Core A1 | 200 | 789.223 | 789.364 | 786.787 | 787.071 | 786.689 | 787.83 |
|  | 250 | 764.867 | 763.899 | 762.535 | 761.835 | 761.444 | 762.91 |
|  | 300 | 742.977 | 741.256 | 740.065 | 740.453 | 738.363 | 740.62 |
|  | 350 | 724.113 | 723.107 | 722.998 | 722.381 | 720.703 | 722.66 |
|  | 400 | 707.246 | 706.542 | 705.728 | 706.282 | 706.539 | 706.47 |
|  | 350 | 727.239 | 726.888 | 727.994 | 726.880 | 727.698 | 727.34 |
|  | 300 | 729.733 | 728.912 | 729.021 | 728.698 | 729.093 | 729.09 |
|  | 250 | 731.548 | 732.563 | 731.056 | 731.892 | 732.966 | 732.00 |
|  | 200 | 738.577 | 738.613 | 741.350 | 740.166 | 741.369 | 740.01 |
|  | 150 | 753.186 | 755.743 | 754.947 | 756.566 | 756.635 | 755.41 |
| length: 5.837 cm | 100 | 775.990 | 778.436 | 779.024 | 780.261 | 782.365 | 779.21 |
| width: 2.525 cm | 70 | 806.146 | 805.729 | 806.593 | 807.904 | 806.143 | 806.59 |
|  | 70 | 688.516 | 687.165 | 687.920 | 687.060 | 687.454 | 687.62 |
| ALV 2608-4-1 | 100 | 626.401 | 625.620 | 626.065 | 625.729 | 626.057 | 625.97 |
| Pc 1 | 150 | 565.271 | 563.772 | 564.854 | 564.410 | 563.430 | 564.35 |
| Core C1-1 | 200 | 531.547 | 529.972 | 529.863 | 529.831 | 528.912 | 530.02 |
|  | 250 | 507.096 | 505.348 | 505.012 | 504.579 | 505.245 | 505.46 |
|  | 300 | 486.959 | 486.771 | 486.493 | 486.415 | 486.727 | 486.67 |
|  | 350 | 472.222 | 472.311 | 471.721 | 471.045 | 471.964 | 471.85 |
|  | 400 | 462.210 | 461.304 | 461.877 | 459.981 | 459.623 | 461.00 |
|  | 350 | 461.322 | 462.435 | 460.961 | 461.092 | 461.560 | 461.47 |
|  | 300 | 465.128 | 464.707 | 465.331 | 464.802 | 465.701 | 465.13 |
|  | 250 | 471.508 | 471.021 | 471.629 | 472.336 | 471.264 | 471.55 |
|  | 200 | 482.808 | 483.077 | 482.236 | 483.054 | 482.665 | 482.77 |



\begin{tabular}{|c|c|c|c|c|c|c|c|}
\hline \multirow[t]{10}{*}{Core B3} \& 200 \& 562.977 \& 561.826 \& 562.064 \& 560.452 \& 560.423 \& 561.55 <br>
\hline \& 250 \& 549.962 \& 549.006 \& 547.481 \& 549.622 \& 547.318 \& 548.68 <br>
\hline \& 300 \& 539.826 \& 539.616 \& 537.957 \& 537.254 \& 536.649 \& 538.26 <br>
\hline \& 350 \& 531.319 \& 530.001 \& 530.137 \& 529.579 \& 528.636 \& 529.93 <br>
\hline \& 400 \& 523.424 \& 523.514 \& 522.722 \& 522.516 \& 522.860 \& 523.01 <br>
\hline \& 350 \& 523.633 \& 524.508 \& 523.441 \& 523.700 \& 523.315 \& 523.72 <br>
\hline \& 300 \& 525.251 \& 524.845 \& 525.165 \& 525.892 \& 525.743 \& 525.38 <br>
\hline \& 250 \& 529.684 \& 529.349 \& 529.007 \& 529.828 \& 529.732 \& 529.52 <br>
\hline \& 200 \& 534.243 \& 533.983 \& 534.272 \& 533.413 \& 533.591 \& 533.90 <br>
\hline \& 150 \& 542.806 \& 543.046 \& 544.963 \& 544.414 \& 543.595 \& 543.76 <br>
\hline \multirow[t]{2}{*}{length: 2.768 cm width: 2.513 cm} \& 100 \& 557.040 \& 562.183 \& 560.094 \& 563.664 \& 563.511 \& 561.29 <br>
\hline \& 70 \& 577.601 \& 577.669 \& 577.834 \& 577.198 \& 578.211 \& 577.70 <br>
\hline \multirow{8}{*}{$$
\begin{gathered}
\text { ALV } 2608 \\
\operatorname{Pc} 2 \\
\text { Core } 1
\end{gathered}
$$} \& 70 \& 6366.8 \& -- \& -- \& -- \& -- \& 6366.8 <br>
\hline \& 100 \& 5838.7 \& -- \& -- \& -- \& -- \& 5838.7 <br>
\hline \& 150 \& 5607.6 \& -- \& -- \& -- \& -- \& 5607.6 <br>
\hline \& 200 \& 5475.2 \& -- \& -- \& -- \& -- \& 5475.2 <br>
\hline \& 250 \& 5216.6 \& -- \& -- \& -- \& -- \& 5216.6 <br>
\hline \& 300 \& 5051.5 \& -- \& -- \& -- \& -- \& 5051.5 <br>
\hline \& 350 \& 5000.4 \& -- \& -- \& -- \& -- \& 5000.4 <br>
\hline \& 400 \& 4953.8 \& -- \& -- \& -- \& -- \& 4953.8 <br>
\hline \multirow{8}{*}{$$
\begin{gathered}
\text { ALV } 2608-4-1 \\
\text { Pc } 2 \\
\text { Core } 2
\end{gathered}
$$} \& 70 \& 927.5 \& -- \& -- \& -- \& -- \& 927.5 <br>
\hline \& 100 \& 862.5 \& -- \& -- \& -- \& -- \& 862.5 <br>
\hline \& 150 \& 768.4 \& -- \& -- \& -- \& -- \& 768.4 <br>
\hline \& 200 \& 661.4 \& -- \& -- \& -- \& -- \& 661.4 <br>
\hline \& 250 \& 620.5 \& -- \& -- \& -- \& -- \& 620.5 <br>
\hline \& 300 \& 578.5 \& -- \& -- \& -- \& -- \& 578.5 <br>
\hline \& 350 \& 546.8 \& -- \& -- \& -- \& -- \& 546.8 <br>
\hline \& 400 \& 519.7 \& -- \& -- \& -- \& -- \& 519.7 <br>
\hline \multirow{8}{*}{$$
\begin{gathered}
\text { ALV 2608-4-1 } \\
\text { Pc } 2 \\
\text { Core } 4
\end{gathered}
$$} \& 70 \& 2989.3 \& -- \& -- \& -- \& -- \& 2989.3 <br>
\hline \& 100 \& 2522.4 \& -- \& -- \& -- \& -- \& 2522.4 <br>
\hline \& 150 \& 1329.1 \& -- \& -- \& -- \& -- \& 1329.1 <br>
\hline \& 200 \& 962.2 \& -- \& -- \& -- \& -- \& 962.2 <br>
\hline \& 250 \& 824.0 \& -- \& -- \& -- \& -- \& 824.0 <br>
\hline \& 300 \& 686.7 \& -- \& -- \& -- \& -- \& 686.7 <br>
\hline \& 350 \& 600.3 \& -- \& -- \& -- \& -- \& 600.3 <br>
\hline \& 400 \& 516.4 \& -- \& -- \& -- \& -- \& 516.4 <br>
\hline \multirow[b]{15}{*}{JAS 177-2-1
Core A2

length: 2.218 cm
width: 2.514 cm} \& 70 \& 314.626 \& 313.346 \& 312.126 \& 310.303 \& 309.883 \& 312.05 <br>
\hline \& 100 \& 243.739 \& 243.397 \& 243.349 \& 243.099 \& 243.725 \& 243.46 <br>
\hline \& 150 \& 215.481 \& 214.521 \& 214.077 \& 213.719 \& 213.385 \& 214.24 <br>
\hline \& 200 \& 199.998 \& 198.496 \& 198.091 \& 198.303 \& 197.660 \& 198.51 <br>
\hline \& 250 \& 187.728 \& 186.088 \& 184.883 \& 184.316 \& 184.134 \& 185.42 <br>
\hline \& 300 \& 171.939 \& 170.666 \& 170.140 \& 169.900 \& 169.130 \& 170.35 <br>
\hline \& 350 \& 160.626 \& 159.590 \& 159.027 \& 159.089 \& 157.934 \& 159.25 <br>
\hline \& 400 \& 153.294 \& 152.328 \& 152.055 \& 151.545 \& 151.417 \& 152.13 <br>
\hline \& 350 \& 152.204 \& 152.165 \& 152.138 \& 152.092 \& 151.725 \& 152.06 <br>
\hline \& 300 \& 154.280 \& 154.130 \& 153.832 \& 154.500 \& 153.939 \& 154.14 <br>
\hline \& 250 \& 160.107 \& 160.671 \& 160.192 \& 160.762 \& 160.773 \& 160.50 <br>
\hline \& 200 \& 168.698 \& 170.030 \& 169.869 \& 170.306 \& 170.118 \& 169.80 <br>
\hline \& 150 \& 185.603 \& 187.132 \& 187.371 \& 187.428 \& 187.657 \& 187.04 <br>
\hline \& 100 \& 199.953 \& 200.603 \& 200.621 \& 201.397 \& 200.972 \& 200.71 <br>
\hline \& 70 \& 218.030 \& 218.950 \& 220.015 \& 220.015 \& 220.404 \& 219.48 <br>

\hline \multirow{13}{*}{$$
\begin{gathered}
\text { JAS } 177-2-1 \\
\text { Core B1 }
\end{gathered}
$$} \& 70 \& 1970.679 \& 1967.907 \& 1952.082 \& 1946.331 \& 1966.707 \& 1960.72 <br>

\hline \& 100 \& 1938.225 \& 1938.673 \& 1933.839 \& 1934.105 \& 1937.615 \& 1936.49 <br>
\hline \& 150 \& 1870.491 \& 1868.112 \& 1866.645 \& 1863.640 \& 1864.134 \& 1866.60 <br>
\hline \& 200 \& 1829.610 \& 1826.120 \& 1826.484 \& 1826.328 \& 1829.966 \& 1827.70 <br>
\hline \& 250 \& 1798.435 \& 1794.460 \& 1794.601 \& 1797.557 \& 1793.888 \& 1795.79 <br>
\hline \& 300 \& 1775.321 \& 1774.622 \& 1775.426 \& 1776.324 \& 1772.495 \& 1774.84 <br>
\hline \& 350 \& 1756.224 \& 1755.267 \& 1755.827 \& 1757.025 \& 1759.920 \& 1756.85 <br>
\hline \& 400 \& 1746.502 \& -- \& 1748.644 \& 1745.181 \& 1743.343 \& 1745.92 <br>
\hline \& 350 \& 1745.179 \& 1751.080 \& 1751.198 \& 1751.386 \& 1752.148 \& 1750.20 <br>
\hline \& 300 \& 1759.851 \& 1754.367 \& 1758.087 \& 1756.928 \& 1754.947 \& 1756.83 <br>
\hline \& 250 \& 1767.359 \& 1765.442 \& 1770.256 \& 1770.705 \& 1768.686 \& 1768.49 <br>
\hline \& 200 \& 1784.591 \& 1789.792 \& 1783.725 \& 1785.145 \& 1790.641 \& 1786.78 <br>
\hline \& 150 \& 1809.819 \& 1811.795 \& 1807.653 \& 1810.306 \& 1808.277 \& 1809.57 <br>
\hline \multirow[t]{2}{*}{length: 4.716 cm width: 2.530 cm} \& 100 \& 1844.081 \& 1850.868 \& 1848.891 \& 1847.695 \& 1850.864 \& 1848.48 <br>
\hline \& 70 \& 1890.155 \& 1892.118 \& 1896.345 \& 1890.910 \& 1895.441 \& 1892.99 <br>
\hline
\end{tabular}

|  | 70 | 3091.102 | 3084.953 | 3090.156 | 3075.858 | 3079.055 | 3084.22 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| JAS 177-2-1 | 100 | 3061.865 | 3053.692 | 3064.990 | 3061.396 | 3060.384 | 3060.46 |
| Core B2 | 150 | 3025.202 | 3031.239 | 3015.798 | 3036.012 | 3017.584 | 3025.16 |
|  | 200 | 3009.429 | 2990.444 | 2997.535 | 2998.335 | 2997.086 | 2998.56 |
|  | 250 | 2972.199 | 2972.043 | 2975.953 | 2970.287 | 2973.411 | 2972.78 |
|  | 300 | 2963.910 | 2950.067 | 2955.262 | 2949.312 | 2952.282 | 2954.16 |
|  | 350 | 2946.318 | 2924.595 | 2944.733 | 2941.581 | 2927.182 | 2936.87 |
|  | 400 | 2895.997 | 2900.238 | 2900.880 | 2898.089 | 2897.285 | 2898.50 |
|  | 350 | 2895.263 | 2896.184 | 2896.945 | 2900.734 | 2901.119 | 2898.05 |
|  | 300 | 2890.881 | 2894.065 | 2901.886 | 2895.291 | 2897.104 | 2895.84 |
|  | 250 | 2900.340 | 2914.090 | 2904.944 | 2903.531 | 2903.011 | 2905.18 |
|  | 200 | 2900.255 | 2902.641 | 2910.156 | 2904.491 | 2904.085 | 2904.32 |
|  | 150 | 2910.503 | 2900.946 | 2915.379 | 2908.103 | 2909.121 | 2908.81 |
| length: 3.940 cm width: 2.490 cm | 100 | 2908.726 | 2911.321 | 2904.708 | 2910.646 | 2913.575 | 2909.79 |
|  | 70 | 2915.579 | 2907.889 | 2912.434 | 2912.921 | 2914.079 | 2912.58 |
| $\begin{gathered} \text { JAS } 177-2-1 \\ \text { Core C2 } \end{gathered}$ | 70 | 227.249 | 226.566 | 226.662 | 226.368 | 226.327 | 226.63 |
|  | 100 | 217.294 | 217.596 | 217.714 | 216.642 | 217.136 | 217.28 |
|  | 150 | 203.239 | 202.206 | 202.509 | 202.303 | 201.895 | 202.43 |
|  | 200 | 194.519 | 193.915 | 194.100 | 193.986 | 194.010 | 194.11 |
|  | 250 | 192.707 | 192.840 | 192.731 | 192.418 | 192.439 | 192.63 |
|  | 300 | 191.615 | 192.347 | 192.282 | 192.324 | 192.058 | 192.13 |
|  | 350 | 191.594 | 191.662 | 191.394 | 191.861 | 191.618 | 191.63 |
|  | 400 | 191.821 | 191.713 | 191.999 | 191.959 | 191.851 | 191.87 |
|  | 350 | 192.400 | 192.625 | 192.448 | 192.009 | 191.940 | 192.28 |
|  | 300 | 192.295 | 192.829 | 192.629 | 192.629 | 192.562 | 192.59 |
|  | 250 | 192.588 | 192.607 | 192.874 | 193.428 | 192.696 | 192.84 |
|  | 200 | 193.275 | 193.233 | 193.118 | 192.832 | 193.428 | 193.18 |
|  | 150 | 194.036 | 194.147 | 194.058 | 193.834 | 194.526 | 194.12 |
| length: 1.684 cm width: 2.494 cm | 100 | 198.135 | 198.350 | 198.803 | 198.659 | 198.840 | 198.56 |
|  | 70 | 204.538 | 205.062 | 205.519 | 206.114 | 205.927 | 205.43 |
| $\begin{gathered} \text { ALV 2179-1-1 } \\ \text { Core A1 } \end{gathered}$ | 70 | 705.981 | 700.726 | 699.559 | 700.027 | 702.002 | 701.66 |
|  | 100 | 668.923 | 669.453 | 667.615 | 668.482 | 667.635 | 668.42 |
|  | 150 | 632.466 | 631.657 | 628.528 | 631.155 | 628.174 | 630.39 |
|  | 200 | 607.264 | 604.945 | 604.588 | 604.883 | 603.348 | 605.00 |
|  | 250 | 588.086 | 588.087 | 589.624 | 585.678 | 586.983 | 587.69 |
|  | 300 | 575.528 | 576.449 | 575.717 | 575.843 | 576.054 | 575.92 |
|  | 350 | 565.491 | 565.159 | 566.121 | 564.705 | 566.500 | 565.59 |
|  | 400 | 560.223 | 561.459 | 560.019 | 561.083 | 560.965 | 560.75 |
|  | 350 | 560.110 | 561.736 | 561.288 | 560.340 | 562.180 | 561.13 |
|  | 300 | 563.446 | 565.603 | 566.188 | 565.287 | 565.028 | 565.11 |
|  | 250 | 569.144 | 568.900 | 569.277 | 569.405 | 569.993 | 569.34 |
|  | 200 | 576.223 | 576.950 | 576.691 | 576.831 | 576.685 | 576.68 |
|  | 150 | 589.704 | 586.060 | 588.477 | 590.419 | 588.755 | 588.68 |
| length: 1.227 cm width: 2.500 cm | 100 | 610.599 | 610.206 | 610.488 | 608.705 | 609.243 | 609.85 |
|  | 70 | 636.543 | 633.144 | 637.494 | 637.501 | 638.534 | 636.64 |
| $\begin{gathered} \text { ALV 2179-1-1 } \\ \text { Core A2 } \end{gathered}$ | 70 | 1174.117 | 1170.142 | 1168.271 | 1159.571 | 1163.192 | 1167.05 |
|  | 100 | 1151.085 | 1143.705 | 1144.000 | 1151.837 | 1152.335 | 1148.59 |
|  | 150 | 1130.405 | 1127.753 | 1126.114 | 1119.097 | 1126.485 | 1125.96 |
|  | 200 | 1111.778 | 1103.340 | 1104.392 | 1105.404 | 1108.562 | 1106.69 |
|  | 250 | 1092.732 | 1095.101 | 1091.446 | 1093.232 | 1096.337 | 1093.77 |
|  | 300 | 1079.162 | 1075.249 | 1080.424 | 1077.993 | 1082.638 | 1079.09 |
|  | 350 | 1069.860 | 1070.511 | 1070.050 | 1065.363 | 1066.531 | 1068.46 |
|  | 400 | 1060.841 | 1063.095 | 1064.589 | 1063.225 | 1061.767 | 1062.70 |
|  | 350 | 1067.572 | 1065.166 | 1063.086 | 1063.529 | 1065.070 | 1064.88 |
|  | 300 | 1067.695 | 1069.764 | 1071.156 | 1064.323 | 1070.163 | 1068.62 |
|  | 250 | 1073.940 | 1074.398 | 1079.232 | 1073.927 | 1078.135 | 1075.92 |
|  | 200 | 1081.777 | 1087.969 | 1085.055 | 1085.629 | 1078.521 | 1083.79 |
|  | 150 | 1094.489 | 1093.103 | 1096.476 | 1095.134 | 1097.036 | 1095.25 |
| length: 1.512 cm width: 2.477 cm | 100 | 1118.171 | 1111.254 | 1116.128 | 1113.354 | 1117.432 | 1115.26 |
|  | 70 | 1129.681 | 1126.553 | 1127.663 | 1136.349 | 1136.837 | 1131.41 |
| $\begin{gathered} \text { ALV 2179-1-1 } \\ \text { Core A3 } \end{gathered}$ | 70 | 1068.020 | 1058.076 | 1045.073 | 1038.420 | 1041.216 | 1050.10 |
|  | 100 | 1001.936 | 1004.255 | 1003.405 | 1002.527 | 1002.657 | 1002.96 |
|  | 150 | 955.274 | 954.003 | 961.515 | 956.513 | 955.928 | 956.64 |
|  | 200 | 926.327 | 928.599 | 930.027 | 924.891 | 929.664 | 927.90 |
|  | 250 | 907.391 | 907.141 | 904.837 | 904.037 | 904.137 | 905.51 |
|  | 300 | 890.551 | 887.398 | 888.108 | 884.158 | 886.297 | 887.30 |


|  | 350 | 876.486 | 879.913 | 874.105 | 872.965 | 873.598 | 875.41 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 400 | 866.835 | 861.410 | 860.928 | 863.097 | 863.697 | 863.19 |
|  | 350 | 862.285 | 861.886 | 862.434 | 863.646 | 863.059 | 862.66 |
|  | 300 | 864.469 | 862.581 | 861.663 | 864.585 | 864.191 | 863.50 |
|  | 250 | 867.398 | 864.946 | 869.532 | 867.521 | 868.132 | 867.50 |
|  | 200 | 871.760 | 875.320 | 873.618 | 876.510 | 875.317 | 874.50 |
|  | 150 | 884.081 | 883.647 | 881.405 | 887.031 | 883.500 | 883.93 |
|  | 100 | 901.256 | 902.294 | 905.432 | 904.437 | 909.051 | 904.49 |
|  | 70 | 926.343 | 927.987 | 927.487 | 928.065 | 924.423 | 926.86 |
| $\begin{aligned} & \text { ALV 2179-1-1 } \\ & \text { Core B2 } \end{aligned}$ | 70 | 1525.934 | 1514.922 | 1517.709 | 1510.970 | 1513.303 | 1516.56 |
|  | 100 | 1514.446 | 1509.204 | 1512.182 | 1510.386 | 1512.859 | 1511.81 |
|  | 150 | 1496.321 | 1496.147 | 1505.554 | 1504.492 | 1502.829 | 1501.06 |
|  | 200 | 1486.043 | 1486.640 | 1484.575 | 1490.167 | 1489.973 | 1487.48 |
|  | 250 | 1486.181 | 1487.855 | 1488.447 | 1480.940 | 1481.430 | 1484.97 |
|  | 300 | 1480.368 | 1469.180 | 1479.771 | 1479.525 | 1473.946 | 1476.55 |
|  | 350 | 1473.449 | 1474.968 | 1471.480 | 1462.585 | 1465.123 | 1469.51 |
|  | 400 | 1461.807 | 1465.717 | 1461.805 | 1464.275 | 1459.899 | 1462.70 |
|  | 350 | 1463.376 | 1465.393 | 1467.523 | 1470.742 | 1464.395 | 1466.28 |
|  | 300 | 1467.562 | 1468.938 | 1471.480 | 1472.988 | 1465.040 | 1469.20 |
|  | 250 | 1466.576 | 1475.790 | 1470.244 | 1470.892 | 1475.539 | 1471.80 |
|  | 200 | 1475.733 | 1480.648 | 1481.299 | 1480.353 | 1475.737 | 1478.75 |
| length: 1.745 cm width: 2.462 cm | 150 | 1482.246 | 1492.937 | 1488.151 | 1484.985 | 1492.164 | 1488.09 |
|  | 100 | 1502.738 | 1500.069 | 1502.000 | 1500.531 | 1507.536 | 1502.57 |
|  | 70 | 1506.464 | 1504.310 | 1513.987 | 1511.296 | 1507.031 | 1508.61 |

## A2.3 Zn-Rich Actively Diffusing Spire Data








## A2.4 Black Smoker Chimney Data




| length: 1.889 cm width: 2.534 cm | 100 70 | $\begin{aligned} & 0.494 \\ & 0.573 \end{aligned}$ | $\begin{aligned} & 0.486 \\ & 0.562 \end{aligned}$ | 0.495 0.574 | 0.503 0.549 | 0.516 0.570 | $\begin{aligned} & 0.50 \\ & 0.57 \end{aligned}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 70 | 0.970 | 1.033 | 0.973 | 1.011 | 0.989 | 1.00 |
| J2-137-1-R1 | 100 | 0.941 | 0.935 | 0.901 | 0.934 | 0.871 | 0.92 |
| Core D3 | 150 | 0.763 | 0.803 | 0.783 | 0.767 | 0.817 | 0.79 |
|  | 200 | 0.774 | 0.722 | 0.775 | 0.742 | 0.775 | 0.76 |
|  | 250 | 0.744 | 0.753 | 0.764 | 0.740 | 0.754 | 0.75 |
|  | 300 | 0.722 | 0.711 | 0.722 | 0.715 | 0.702 | 0.71 |
|  | 350 | 0.676 | 0.691 | 0.696 | 0.704 | 0.699 | 0.69 |
|  | 400 | 0.670 | 0.701 | 0.684 | 0.692 | 0.684 | 0.69 |
|  | 350 | 0.696 | 0.713 | 0.683 | 0.695 | 0.715 | 0.70 |
|  | 300 | 0.701 | 0.714 | 0.736 | 0.716 | 0.718 | 0.72 |
|  | 250 | 0.750 | 0.726 | 0.727 | 0.731 | 0.730 | 0.73 |
|  | 200 | 0.764 | 0.714 | 0.724 | 0.745 | 0.757 | 0.74 |
|  | 150 | 0.755 | 0.766 | 0.788 | 0.784 | 0.772 | 0.77 |
| length: 2.123 cm width: 2.533 cm | 100 | 0.781 | 0.790 | 0.778 | 0.801 | 0.789 | 0.79 |
|  | 70 | 0.802 | 0.810 | 0.835 | 0.810 | 0.794 | 0.81 |
| $\begin{gathered} \mathrm{J} 2-137-1-\mathrm{R} 1 \\ \text { Core D4 } \end{gathered}$ | 70 | 0.894 | 0.908 | 0.869 | 0.918 | 0.905 | 0.90 |
|  | 100 | 0.755 | 0.732 | 0.758 | 0.736 | 0.758 | 0.75 |
|  | 150 | 0.689 | 0.667 | 0.658 | 0.637 | 0.661 | 0.66 |
|  | 200 | 0.641 | 0.611 | 0.612 | 0.632 | 0.622 | 0.62 |
|  | 250 | 0.615 | 0.663 | 0.612 | 0.605 | 0.641 | 0.63 |
|  | 300 | 0.604 | 0.613 | 0.607 | 0.618 | 0.593 | 0.61 |
|  | 350 | 0.593 | 0.593 | 0.605 | 0.599 | 0.604 | 0.60 |
|  | 400 | 0.599 | 0.607 | 0.590 | 0.587 | 0.600 | 0.60 |
|  | 350 | 0.603 | 0.578 | 0.610 | 0.600 | 0.622 | 0.60 |
|  | 300 | 0.624 | 0.612 | 0.613 | 0.593 | 0.615 | 0.61 |
|  | 250 | 0.617 | 0.614 | 0.636 | 0.626 | 0.615 | 0.62 |
|  | 200 | 0.651 | 0.617 | 0.628 | 0.626 | 0.636 | 0.63 |
|  | 150 | 0.630 | 0.652 | 0.652 | 0.632 | 0.657 | 0.65 |
| length: 2.139 cm width: 2.526 cm | 100 | 0.660 | 0.665 | 0.671 | 0.642 | 0.654 | 0.66 |
|  | 70 | 0.686 | 0.697 | 0.687 | 0.697 | 0.679 | 0.69 |
| $\begin{gathered} \text { J2-213-3-R1 } \\ \text { Core A1 } \end{gathered}$ | 70 | 29.573 | 29.698 | 30.018 | 29.730 | 29.741 | 29.75 |
|  | 100 | 25.614 | 25.203 | 25.234 | 25.114 | 25.153 | 25.26 |
|  | 150 | 21.951 | 21.891 | 21.836 | 22.030 | 21.795 | 21.90 |
|  | 200 | 20.371 | 20.347 | 20.248 | 20.217 | 20.221 | 20.28 |
|  | 250 | 19.196 | 19.305 | 19.472 | 19.242 | 19.212 | 19.29 |
|  | 300 | 18.536 | 18.428 | 18.376 | 18.471 | 18.434 | 18.45 |
|  | 350 | 17.922 | 17.957 | 17.873 | 17.813 | 17.939 | 17.90 |
|  | 400 | 17.537 | 17.563 | 17.454 | 17.359 | 17.440 | 17.47 |
|  | 350 | 17.447 | 17.485 | 17.539 | 17.439 | 17.507 | 17.48 |
|  | 300 | 17.575 | 17.604 | 17.761 | 17.672 | 17.640 | 17.65 |
|  | 250 | 17.956 | 17.833 | 17.860 | 17.907 | 17.884 | 17.89 |
|  | 200 | 18.251 | 18.100 | 18.301 | 18.219 | 18.325 | 18.24 |
|  | 150 | 18.864 | 18.728 | 18.713 | 18.744 | 18.837 | 18.78 |
| length: 2.287 cm width: 2.492 cm | 100 | 19.674 | 19.852 | 19.752 | 19.930 | 19.908 | 19.82 |
|  | 70 | 21.015 | 21.125 | 21.184 | 21.142 | 21.231 | 21.14 |
| $\begin{gathered} \text { J2-213-3-R1 } \\ \text { Core B1 } \end{gathered}$ | 70 | 48.768 | 48.880 | 48.898 | 49.026 | 48.972 | 48.91 |
|  | 100 | 45.986 | 45.305 | 45.418 | 45.442 | 45.399 | 45.51 |
|  | 150 | 42.314 | 42.275 | 42.364 | 42.090 | 42.166 | 42.24 |
|  | 200 | 40.704 | 40.712 | 40.296 | 40.307 | 40.334 | 40.47 |
|  | 250 | 39.230 | 38.910 | 39.001 | 38.968 | 39.086 | 39.04 |
|  | 300 | 37.503 | 37.663 | 37.684 | 37.603 | 37.940 | 37.68 |
|  | 350 | 36.207 | 36.242 | 36.348 | 36.000 | 36.002 | 36.16 |
|  | 400 | 35.086 | 35.063 | 34.779 | 34.917 | 34.839 | 34.94 |
|  | 350 | 35.011 | 35.157 | 35.150 | 35.242 | 35.331 | 35.18 |
|  | 300 | 35.806 | 35.684 | 35.989 | 35.769 | 35.889 | 35.83 |
|  | 250 | 36.595 | 36.466 | 36.581 | 36.524 | 36.601 | 36.55 |
|  | 200 | 37.253 | 37.486 | 37.189 | 37.520 | 37.184 | 37.33 |
|  | 150 | 38.419 | 38.277 | 38.343 | 38.123 | 38.233 | 38.28 |
| length: 1.526 cm width: 2.485 cm | 100 | 39.686 | 39.518 | 39.546 | 39.545 | 39.695 | 39.60 |
|  | 70 | 41.330 | 41.643 | 41.507 | 41.470 | 41.591 | 41.51 |
| $\begin{gathered} \text { J2-213-3-R1 } \\ \text { Core D1 } \end{gathered}$ | 70 | 145.398 | 144.73 | 146.158 | 145.144 | 144.465 | 145.18 |
|  | 100 | 135.884 | 135.292 | 135.557 | 134.619 | 136.561 | 135.58 |
|  | 150 | 123.745 | 122.970 | 123.249 | 121.982 | 122.768 | 122.94 |
|  | 200 | 115.644 | 115.183 | 115.355 | 114.749 | 115.242 | 115.23 |


|  | 250 | 106.472 | 106.354 | 106.853 | 106.308 | 105.595 | 105.60 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 300 | 99.981 | 100.422 | 99.220 | 99.071 | 98.800 | 99.50 |
|  | 350 | 92.481 | 92.160 | 91.900 | 92.009 | 91.986 | 92.11 |
|  | 400 | 87.036 | 86.522 | 86.651 | 86.247 | 86.657 | 86.62 |
|  | 350 | 87.533 | 87.292 | 87.312 | 87.492 | 87.623 | 87.45 |
|  | 300 | 89.576 | 89.806 | 89.764 | 90.623 | 89.867 | 89.93 |
|  | 250 | 94.961 | 95.042 | 95.249 | 95.322 | 95.471 | 95.21 |
| length: 2.140 cm | 150 | 110.934 | 110.159 | 109.454 | 110.320 | 110.540 | 101.86 |
| width: 2.534 cm | 100 | 118.407 | 119.333 | 118.587 | 118.251 | 119.122 | 110.28 |
|  | 70 | 125.588 | 126.324 | 125.950 | 124.985 | 125.134 | 118.74 |
|  |  | 200 | 101.409 | 101.525 | 101.889 | 102.534 | 101.938 |

## A2.5 Relict Spire Data










|  | 150 | 113.721 | 113.565 | 114.081 | 114.629 | 113.879 | 113.97 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| length: 3.389 cm <br> width: 2.564 cm | 100 | 116.575 | 116.777 | 116.954 | 116.368 | 116.928 | 116.72 |
|  | 70 | 120.350 | 120.522 | 120.478 | 120.523 | 121.252 | 120.62 |

## Appendix 3: Porosity Data

## A3.1 Massive Anhydrite Data

| Sample | Confining Pressure (psi) | 1 | 2 | rosity 3 | 4 | 5 | Mean $\phi$ (\%) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{gathered} \text { J2-210-8-R2 } \\ \text { Core A-1 } \end{gathered}$ | 300 | 2.964 | 2.173 | 1.540 | 2.426 | 2.807 | 2.32 |
|  | 400 | 3.315 | 1.287 | 3.062 | 3.062 | 3.347 | 2.66 |
|  | 450 | 3.570 | 0.404 | 2.046 | 2.299 | 3.473 | 1.88 |
| $\begin{gathered} \text { J2-210-8-R2 } \\ \text { Core A-2 } \end{gathered}$ | 300 | 3.125 | 3.742 | 2.198 | 3.125 | 0.276 | 1.86 |
|  | 400 | 3.051 | 0.583 | 2.198 | 2.508 | 3.361 | 2.01 |
|  | 450 | 4.288 | 2.198 | 3.125 | 3.125 | 0.276 | 1.91 |
| $\begin{gathered} \text { J2-210-8-R2 } \\ \text { Core A-3 } \end{gathered}$ | 300 | 5.945 | 5.945 | 0.847 | 2.537 | 2.203 | 2.78 |
|  | 400 | 5.737 | 0.100 | 3.079 | 5.527 | 4.645 | 2.14 |
|  | 450 | 5.527 | 2.203 | 3.893 | 5.465 | 1.723 | 3.39 |
| $\begin{gathered} \mathrm{J} 2-216-5-\mathrm{R} 1 \\ \text { Core A2 } \end{gathered}$ | 300 | 9.147 | 12.644 | 13.754 | 9.106 | 13.933 | 11.51 |
|  | 400 | 13.933 | 8.795 | 12.153 | 13.399 | 11.445 | 11.80 |
|  | 450 | 13.221 | 10.383 | 14.070 | 10.559 | 9.063 | 11.31 |
| $\begin{gathered} \text { J2-216-14-R1 } \\ \text { Core A } \end{gathered}$ | 300 | 3.643 | 3.191 | 1.317 | 1.194 | 2.009 | 2.06 |
|  | 400 | 3.316 | 2.537 | 3.807 | 3.480 | 2.537 | 3.09 |
|  | 450 | 3.519 | 3.972 | 1.557 | 3.846 | 3.480 | 3.11 |
| $\begin{gathered} \mathrm{J} 2-216-14-\mathrm{R} 1 \\ \text { Core B } \end{gathered}$ | 300 | 5.624 | 5.606 | 5.684 | 5.624 | 5.624 | 5.63 |
|  | 400 | 5.527 | 5.527 | 5.389 | 5.467 | 5.389 | 5.46 |
|  | 450 | 5.467 | 5.467 | 5.389 | 5.467 | 5.546 | 5.47 |
| $\begin{aligned} & \text { J301-3 } \\ & \text { Core A } \end{aligned}$ | 300 | 14.795 | 14.795 | 14.846 | 14.846 | 14.846 | 14.83 |
|  | 400 | 14.757 | 14.757 | 14.706 | 14.757 | 14.846 | 14.77 |
|  | 450 | 14.757 | 14.757 | 14.666 | 14.666 | 14.717 | 14.71 |
| $\begin{aligned} & \text { J301-3 } \\ & \text { Core B } \end{aligned}$ | 300 | 8.001 | 7.861 | 7.892 | 7.946 | 7.892 | 7.92 |
|  | 400 | 7.837 | 7.752 | 7.837 | 7.892 | 7.783 | 7.82 |
|  | 450 | 7.837 | 7.892 | 7.868 | 7.783 | 7.892 | 7.85 |
| $\begin{gathered} \text { ALV 2581-8 } \\ \text { Core A } \end{gathered}$ | 300 | 12.696 | 12.734 | 12.630 | 12.668 | 12.630 | 12.67 |
|  | 400 | 12.630 | 12.601 | 12.564 | 12.564 | 12.601 | 12.59 |
|  | 450 | 12.497 | 12.535 | 12.535 | 12.601 | 12.535 | 12.54 |
| MIR 1, 1/74 | 300 | 11.017 | 11.017 | 10.876 | 10.806 | 10.876 | 10.92 |
| Sta 2403 | 400 | 10.876 | 10.806 | 10.806 | 10.806 | 10.876 | 10.83 |
| Core A-1 | 450 | 10.946 | 10.946 | 10.876 | 10.876 | 10.806 | 10.89 |
| MIR 1, 1/74 | 300 | 12.322 | 12.173 | 12.173 | 12.130 | 12.023 | 12.16 |
| Sta 2403 | 400 | 12.023 | 12.066 | 12.023 | 11.949 | 11.800 | 11.97 |
| Core A-2 | 450 | 12.023 | 11.949 | 11.875 | 11.875 | 11.949 | 11.93 |
| MIR 1, 1/74 | 300 | 9.073 | 9.007 | 8.942 | 9.045 | 9.007 | 9.02 |
| Sta 2403 | 400 | 8.942 | 8.914 | 8.914 | 8.914 | 8.877 | 8.91 |
| Core B | 450 | 8.914 | 8.783 | 8.783 | 8.783 | 8.914 | 8.84 |
| MIR 1, 2/78 | 300 | 15.105 | 14.960 | 14.992 | 14.919 | 14.992 | 14.99 |
| Sta 2417 | 400 | 14.774 | 14.743 | 14.847 | 14.774 | 14.702 | 14.77 |
| Core A | 450 | 14.670 | 14.670 | 14.670 | 14.630 | 14.670 | 14.66 |
| MIR 1, 2/78 | 300 | 13.497 | 13.587 | 13.416 | 13.416 | 13.416 | 13.47 |
| Sta 2417 | 400 | 13.358 | 13.244 | 13.301 | 13.301 | 13.244 | 13.29 |
| Core B | 450 | 13.188 | 13.188 | 13.277 | 13.188 | 13.220 | 13.21 |
| ALV 21837-0 Core 2 | 300 | 5.020 | -- | -- | -- | -- | 5.02 |
|  | 400 | 3.641 | -- | -- | -- | -- | 3.64 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| ALV 21837-0 Core 3 | 300 | 6.115 | -- | -- | -- | -- | 6.12 |
|  | 400 | 2.395 | -- | -- | -- | -- | 2.40 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| ALV 21837-0 <br> Core B | 300 | 6.633 | -- | -- | -- | -- | 6.63 |
|  | 400 | 4.022 | -- | -- | -- | -- | 4.02 |
|  | 450 | -- | -- | -- | -- | -- | -- |

## A3.2 Flange, Slab and Crust Data

| Sample | Confining Pressure (psi) | 1 | 2 | rosity | 4 | 5 | Mean $\phi$ (\%) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{aligned} & \text { ALV } 3517-\mathrm{R} 1 \\ & \text { Core D1 } \end{aligned}$ | 300 | 25.098 | 21.125 | 21.801 | 21.171 | 21.664 | 22.12 |
|  | 400 | 24.330 | 21.486 | 21.084 | 21.084 | 20.409 | 21.64 |
|  | 450 | 24.147 | 20.768 | 20.587 | 20.368 | 19.295 | 20.97 |
| $\begin{gathered} \text { ALV 3517-R1 } \\ \text { Core D3 } \end{gathered}$ | 300 | 23.724 | 18.833 | 19.017 | 21.247 | 18.508 | 20.17 |
|  | 400 | 23.119 | 19.202 | 20.318 | 18.094 | 18.461 | 19.76 |
|  | 450 | 22.928 | 18.603 | 19.387 | 18.461 | 19.344 | 19.68 |
| $\begin{gathered} \text { ALV } 3521-\mathrm{R} 2 \\ \text { Core A2 } \end{gathered}$ | 300 | 42.067 | 41.983 | 41.983 | 41.983 | 41.983 | 42.00 |
|  | 400 | 40.570 | 40.570 | 40.570 | 40.570 | 40.369 | 40.53 |
|  | 450 | 39.968 | 39.968 | 40.079 | 40.079 | 40.079 | 40.03 |
| $\begin{aligned} & \text { ALV 3521-R2 } \\ & \text { Core Ex } 1 \end{aligned}$ | 300 | 45.559 | 45.470 | 45.074 | 45.272 | 45.272 | 45.33 |
|  | 400 | 44.671 | 44.671 | 44.671 | 44.869 | 44.869 | 44.75 |
|  | 450 | 44.296 | 44.160 | 44.473 | 44.473 | 44.269 | 44.33 |
| $\begin{gathered} \text { ALV 3521-R2 } \\ \text { Core Ex } 2 \end{gathered}$ | 300 | 46.020 | 45.803 | 46.239 | 45.747 | 45.747 | 45.91 |
|  | 400 | 44.878 | 44.659 | 44.659 | 44.451 | 44.659 | 44.66 |
|  | 450 | 44.017 | 43.800 | 44.396 | 43.745 | 43.800 | 43.95 |
| $\begin{aligned} & \text { ALV 2415-1B } \\ & \text { Core A1 } \end{aligned}$ | 300 | 19.230 | 20.608 | 20.529 | 19.977 | 20.667 | 20.20 |
|  | 400 | 20.806 | 20.390 | 20.667 | 20.390 | 20.390 | 20.53 |
|  | 450 | 20.390 | 20.943 | 20.529 | 20.529 | 20.667 | 20.61 |
| $\begin{aligned} & \text { ALV 2415-1B } \\ & \text { Core B1 } \end{aligned}$ | 300 | 31.238 | 31.141 | 31.372 | 31.105 | 31.008 | 31.17 |
|  | 400 | 31.105 | 30.971 | 30.971 | 30.971 | 30.638 | 30.93 |
|  | 450 | 30.807 | 30.874 | 30.771 | 30.673 | 30.638 | 30.75 |
| ALV 2415-1B Core 1 | 300 | 17.110 | -- | -- | -- | -- | 17.110 |
|  | 400 | 15.271 | -- | -- | -- | -- | 15.271 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| ALV 2415-1B Core 2 | 300 | 19.164 | -- | -- | -- | -- | 19.164 |
|  | 400 | 17.216 | -- | -- | -- | -- | 17.216 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| $\begin{gathered} \text { ALV 2927-3 } \\ \text { Core B1 } \end{gathered}$ | 300 | 38.150 | 38.219 | 38.219 | 38.277 | 38.219 | 38.22 |
|  | 400 | 37.899 | 37.969 | 37.969 | 37.899 | 37.969 | 37.94 |
|  | 450 | 37.775 | 37.845 | 37.845 | 37.969 | 37.845 | 37.86 |
| ALV 2927-3 Core 1 | 300 | 41.806 | -- | -- | -- | -- | 41.806 |
|  | 400 | 39.623 | -- | -- | -- | -- | 39.623 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| ALV 2927-3 Core 2 | 300 | 32.669 | -- | -- | -- | -- | 32.669 |
|  | 400 | 29.474 | -- | -- | -- | -- | 29.474 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| ALV 2927-3 Core 3 | 300 | 40.726 | -- | -- | -- | -- | 40.726 |
|  | 400 | 38.191 | -- | -- | -- | -- | 38.191 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| $\begin{gathered} \text { J2-286 } \\ \text { Core A1 } \end{gathered}$ | 300 | 20.686 | 20.686 | 20.686 | 20.581 | 20.686 | 20.66 |
|  | 400 | 20.686 | 20.371 | 20.371 | 20.476 | 20.476 | 20.48 |
|  | 450 | 20.476 | 20.371 | 20.371 | 20.371 | 20.208 | 20.36 |
| $\begin{gathered} \mathrm{J} 2-286 \\ \text { Core A3 } \end{gathered}$ | 300 | 24.025 | 25.195 | 25.225 | 25.860 | 26.063 | 25.26 |
|  | 400 | 26.100 | 25.255 | 25.023 | 24.957 | 25.225 | 25.31 |
|  | 450 | 26.131 | 25.225 | 25.121 | 25.225 | 25.157 | 25.37 |
| $\begin{gathered} \text { J2-286 } \\ \text { Core C2-1 } \end{gathered}$ | 300 | 27.407 | 27.494 | 26.959 | 27.043 | 27.257 | 27.23 |
|  | 400 | 27.494 | 27.107 | 27.343 | 27.193 | 27.193 | 27.27 |
|  | 450 | 27.193 | 27.043 | 27.193 | 27.127 | 27.043 | 27.12 |
| $\begin{gathered} \text { J2-286 } \\ \text { Core C2-2 } \end{gathered}$ | 300 | 24.243 | 23.990 | 24.143 | 24.089 | 24.143 | 24.12 |
|  | 400 | 24.539 | 24.143 | 24.102 | 24.143 | 23.947 | 24.17 |
|  | 450 | 24.341 | 24.200 | 24.046 | 24.200 | 24.200 | 24.20 |
| $\begin{gathered} \text { J2-286 } \\ \text { Core C3-1 } \end{gathered}$ | 300 | 31.056 | 28.503 | 27.191 | 25.301 | 26.118 | 27.56 |
|  | 400 | 30.168 | 28.945 | 26.879 | 27.097 | 25.989 | 27.77 |
|  | 450 | 30.263 | 27.532 | 26.443 | 28.945 | 29.731 | 28.55 |
| $\begin{gathered} \text { J2-286 } \\ \text { Core C3-2 } \end{gathered}$ | 300 | 22.005 | 20.927 | 20.523 | 21.466 | 21.272 | 21.23 |
|  | 400 | 21.677 | 20.466 | 19.929 | 20.122 | 21.330 | 20.69 |
|  | 450 | 21.523 | 19.929 | 21.196 | 19.854 | 20.332 | 20.56 |
| $\begin{gathered} \mathrm{J} 2-286 \\ \text { Core C4 } \end{gathered}$ | 300 | 22.267 | 20.833 | 20.443 | 19.152 | 19.667 | 20.44 |
|  | 400 | 21.688 | 19.797 | 20.056 | 18.951 | 19.281 | 19.93 |
|  | 450 | 21.484 | 19.152 | 19.355 | 18.839 | 19.409 | 19.63 |


| ALV 2608-3-3 Core B2 | 300 | 43.155 | 42.980 | 42.760 | 42.760 | 43.143 | 42.96 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 400 | 42.599 | 42.647 | 42.314 | 42.426 | 42.709 | 42.54 |
|  | 450 | 42.426 | 42.647 | 41.872 | 42.154 | 42.204 | 42.26 |
| $\begin{gathered} \text { ALV 2608-3-3 } \\ \text { Core C1 } \end{gathered}$ | 300 | 29.401 | 29.252 | 28.955 | 29.103 | 29.103 | 29.16 |
|  | 400 | 28.870 | 28.658 | 28.955 | 28.870 | 28.573 | 28.79 |
|  | 450 | 28.807 | 28.573 | 28.573 | 28.510 | 28.807 | 28.65 |
| $\begin{gathered} \hline \text { ALV } 2608-4-1 \\ \text { Pc } 1 \\ \text { Core A1 } \\ \hline \end{gathered}$ | 300 | 35.109 | 34.934 | 34.865 | 34.934 | 34.865 | 34.94 |
|  | 400 | 34.971 | 34.794 | 34.865 | 34.587 | 34.725 | 34.79 |
|  | 450 | 34.832 | 34.556 | 34.656 | 34.794 | 34.794 | 34.73 |
| $\begin{gathered} \hline \text { ALV 2608-4-1 } \\ \text { Pc } 1 \\ \text { Core C1-1 } \\ \hline \end{gathered}$ | 300 | 37.501 | 37.326 | 36.938 | 37.840 | 36.538 | 37.23 |
|  | 400 | 37.050 | 36.650 | 36.825 | 36.663 | 35.370 | 36.51 |
|  | 450 | 37.275 | 35.816 | 35.705 | 34.876 | 35.147 | 35.75 |
| $\begin{gathered} \hline \text { ALV 2608-4-1 } \\ \text { Pc } 1 \\ \text { Core C1-2 } \\ \hline \end{gathered}$ | 300 | 46.627 | 46.561 | 45.672 | 46.413 | 46.644 | 46.38 |
|  | 400 | 46.561 | 46.116 | 46.495 | 46.264 | 46.199 | 46.33 |
|  | 450 | 45.902 | 45.968 | 46.199 | 45.672 | 45.754 | 45.90 |
| $\begin{gathered} \hline \text { ALV 2608-4-1 } \\ \text { Pc } 1 \\ \text { Core C3-1 } \\ \hline \end{gathered}$ | 300 | 35.624 | 34.357 | 34.902 | 35.093 | 35.226 | 35.04 |
|  | 400 | 35.093 | 34.828 | 34.828 | 34.638 | 34.770 | 34.77 |
|  | 450 | 35.035 | 35.300 | 34.373 | 35.300 | 35.300 | 35.06 |
| $\begin{gathered} \hline \text { ALV 2608-4-1 } \\ \text { Pc } 1 \\ \text { Core C3-2 } \end{gathered}$ | 300 | 37.516 | 36.459 | 37.255 | 37.516 | 37.182 | 37.18 |
|  | 400 | 37.124 | 37.573 | 36.995 | 36.735 | 37.124 | 37.11 |
|  | 450 | 37.182 | 36.662 | 36.402 | 35.827 | 36.995 | 36.61 |
| $\begin{gathered} \text { ALV 2608-4-1 } \\ \text { Pc } 2 \\ \text { Core A3 } \\ \hline \end{gathered}$ | 300 | 39.877 | 39.660 | 40.022 | 39.877 | 39.733 | 39.83 |
|  | 400 | 39.699 | 39.626 | 39.589 | 39.660 | 39.555 | 39.63 |
|  | 450 | 39.771 | 39.660 | 39.555 | 39.555 | 39.555 | 39.62 |
| $\begin{gathered} \hline \text { ALV } 2608-4-1 \\ \text { Pc } 2 \\ \text { Core B3 } \\ \hline \end{gathered}$ | 300 | 43.230 | 43.159 | 43.105 | 43.105 | 43.105 | 43.14 |
|  | 400 | 42.923 | 42.727 | 42.980 | 42.852 | 42.727 | 42.84 |
|  | 450 | 42.727 | 42.478 | 42.673 | 42.798 | 42.798 | 42.69 |
| ALV 2608-4-1 | 300 | 45.569 | -- | -- | -- | -- | 45.569 |
| Pc 2 <br> Core 1 | 400 | 44.132 | -- | -- | -- | -- | 44.132 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| ALV 2608-4-1 | 300 | 47.969 | -- | -- | -- | -- | 47.969 |
| Pc 2 <br> Core 2 | 400 | 46.249 | -- | -- | -- | -- | 46.249 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| $\begin{gathered} \text { ALV 2608-4-1 } \\ \text { Pc } 2 \\ \text { Core } 4 \\ \hline \end{gathered}$ | 300 | 44.009 | -- | -- | -- | -- | 44.009 |
|  | 400 | 42.376 | -- | -- | -- | -- | 42.376 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| $\begin{gathered} \text { JAS } 177-2-1 \\ \text { Core A2 } \end{gathered}$ | 300 | 46.347 | 46.218 | 46.071 | 46.071 | 45.754 | 46.09 |
|  | 400 | 45.924 | 45.772 | 45.543 | 45.772 | 45.478 | 45.70 |
|  | 450 | 45.395 | 45.772 | 45.327 | 45.395 | 45.327 | 45.44 |
| $\begin{gathered} \text { JAS 177-2-1 } \\ \text { Core B1 } \end{gathered}$ | 300 | 39.921 | 38.307 | 38.472 | 39.337 | 38.179 | 38.84 |
|  | 400 | 39.383 | 38.390 | 38.472 | 38.472 | 39.588 | 38.86 |
|  | 450 | 39.217 | 37.935 | 38.472 | 38.263 | 38.060 | 38.39 |
| $\begin{gathered} \text { JAS 177-2-1 } \\ \text { Core B2 } \end{gathered}$ | 300 | 42.600 | 40.598 | 40.167 | 41.171 | 41.267 | 41.15 |
|  | 400 | 42.128 | 40.547 | 40.315 | 41.789 | 40.453 | 41.04 |
|  | 450 | 41.692 | 40.505 | 40.409 | 40.598 | 39.789 | 40.59 |
| $\begin{gathered} \text { JAS } 177-2-1 \\ \text { Core C2 } \end{gathered}$ | 300 | 37.972 | 41.246 | 42.354 | 42.354 | 42.088 | 41.17 |
|  | 400 | 41.615 | 41.062 | 38.781 | 38.051 | 38.051 | 39.48 |
|  | 450 | 41.062 | 38.051 | 38.051 | 39.147 | 36.959 | 38.63 |
| $\begin{gathered} \text { ALV 2179-4-1 } \\ \text { Core A1 } \end{gathered}$ | 300 | 38.877 | 38.396 | 35.518 | 35.892 | 36.131 | 36.94 |
|  | 400 | 38.052 | 34.800 | 37.090 | 34.358 | 32.661 | 35.34 |
|  | 450 | 36.475 | 32.322 | 34.358 | 29.859 | 32.456 | 33.02 |
| $\begin{gathered} \text { ALV 2179-4-1 } \\ \text { Core A2 } \end{gathered}$ | 300 | 41.325 | 36.814 | 35.705 | 39.910 | 36.102 | 37.91 |
|  | 400 | 39.507 | 35.817 | 37.014 | 37.414 | 39.419 | 37.81 |
|  | 450 | 38.816 | 37.100 | 36.216 | 37.100 | 36.814 | 37.20 |
| $\begin{gathered} \text { ALV 2179-4-1 } \\ \text { Core A3 } \end{gathered}$ | 300 | 40.768 | 39.392 | 39.981 | 36.275 | 35.887 | 38.41 |
|  | 400 | 39.392 | 36.662 | 36.468 | 37.161 | 36.275 | 37.17 |
|  | 450 | 38.805 | 36.857 | 34.837 | 36.578 | 37.356 | 36.86 |
| $\begin{gathered} \text { ALV 2179-4-1 } \\ \text { Core B2 } \end{gathered}$ | 300 | 43.273 | 43.191 | 43.191 | 43.011 | 43.273 | 43.19 |
|  | 400 | 42.622 | 42.256 | 42.256 | 42.176 | 42.541 | 42.37 |
|  | 450 | 42.256 | 42.071 | 42.071 | 41.996 | 42.176 | 42.11 |

## A3.3 Zn-Rich Actively Diffusing Spire Data

|  | Confining | Porosity (\%) |  |  |  |  | $\begin{gathered} \text { Mean } \phi \\ (\%) \end{gathered}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Sample | Pressure <br> (psi) | 1 | 2 | 3 | 4 | 5 |  |
| ALV 2187-1-1 | 300 | 45.72 | 45.57 | 45.41 | 45.57 | 45.37 | 45.53 |
| top | 400 | 44.20 | 44.00 | 44.01 | 44.11 | 44.11 | 44.08 |
| Core A2 | 450 | 43.60 | 43.35 | 43.30 | 43.60 | 43.40 | 43.45 |
| ALV 2187-1-1 | 300 | 36.72 | 36.51 | 36.40 | 36.56 | 36.51 | 36.54 |
| top | 400 | 35.93 | 35.93 | 35.87 | 35.77 | 35.72 | 35.84 |
| Core A4 | 450 | 35.62 | 35.56 | 35.66 | 35.45 | 35.66 | 35.59 |
| ALV 2187-1-1 | 300 | 45.93 | 45.48 | 45.48 | 45.02 | 45.38 | 45.46 |
| top | 400 | 44.80 | 43.89 | 44.12 | 44.12 | 44.02 | 44.19 |
| Core B1 | 450 | 43.89 | 43.67 | 43.44 | 43.67 | 43.44 | 43.62 |
| ALV 2187-1-1 | 300 | 39.15 | 38.87 | 38.61 | 38.79 | 39.15 | 38.92 |
| bottom | 400 | 38.44 | 38.44 | 38.51 | 38.16 | 38.26 | 38.36 |
| Core A2 | 450 | 37.98 | 38.08 | 38.08 | 38.08 | 38.08 | 38.06 |
| ALV 2187-1-1 | 300 | 42.28 | 42.12 | 42.12 | 41.95 | 42.03 | 42.10 |
| bottom | 400 | 41.63 | 41.79 | 41.70 | 42.03 | 41.79 | 41.79 |
| Core B1 | 450 | 41.21 | 41.04 | 41.04 | 41.21 | 40.79 | 41.06 |
| ALV 2187-1-1 | 300 | 41.04 | 41.11 | 40.91 | 40.98 | 41.09 | 41.03 |
| bottom | 400 | 40.80 | 40.44 | 40.44 | 40.44 | 40.44 | 40.51 |
| Core C2 | 450 | 40.34 | 40.34 | 40.34 | 40.44 | 40.52 | 40.39 |
| ALV 2187-1-1 | 300 | 40.81 | 40.81 | 40.59 | 40.59 | 40.37 | 40.63 |
| bottom | 400 | 40.25 | 40.00 | 40.03 | 40.03 | 39.71 | 40.00 |
| Core C3 | 450 | 40.09 | 39.71 | 40.03 | 39.15 | 39.81 | 39.76 |
|  | 300 | 41.88 | 41.63 | 41.82 | 41.78 | 41.78 | 41.78 |
| ALV 2187-1-2 | 400 | 41.39 | 41.39 | 41.29 | 41.19 | 41.19 | 41.29 |
| Core B2-2 | 450 | 41.00 | 40.90 | 40.80 | 40.86 | 40.96 | 40.90 |
|  | 300 | 39.47 | 39.33 | 39.33 | 39.29 | 39.42 | 39.37 |
| ALV 2187-1 | 400 | 38.84 | 38.66 | 38.75 | 38.70 | 38.66 | 38.72 |
| Core C1/C2 | 450 | 38.57 | 38.48 | 38.39 | 38.39 | 38.34 | 38.43 |
|  | 300 | 41.41 | 41.53 | 41.21 | 41.16 | 41.23 | 41.31 |
| ALV 2187-1-2 | 400 | 40.98 | 40.78 | 40.85 | 40.85 | 40.98 | 40.89 |
|  | 450 | 40.73 | 40.60 | 40.48 | 40.48 | 40.35 | 40.53 |
|  | 300 | 42.59 | 42.41 | 42.21 | 42.34 | 42.34 | 42.38 |
| ALV 2187-1-2 | 400 | 41.66 | 41.72 | 41.35 | 41.17 | 41.42 | 41.46 |
| Core C3-2 | 450 | 40.61 | 40.55 | 40.43 | 40.68 | 40.68 | 40.59 |
|  | 300 | 43.03 | 42.89 | 42.83 | 42.89 | 42.69 | 42.87 |
|  | 400 | 42.30 | 42.30 | 42.16 | 42.20 | 42.20 | 42.23 |
|  | 450 | 42.10 | 41.96 | 41.81 | 41.96 | 42.20 | 42.00 |
|  | 300 | 47.52 | 47.39 | 47.64 | 47.45 | 47.39 | 47.48 |
|  | 400 | 47.02 | 46.78 | 46.84 | 46.90 | 46.78 | 46.86 |
|  | 450 | 46.65 | 46.84 | 46.41 | 46.60 | 46.97 | 46.69 |
|  | 300 | 48.06 | 47.56 | 47.56 | 47.76 | 47.17 | 47.62 |
|  | 400 | 46.59 | 45.81 | 45.92 | 46.00 | 46.00 | 46.06 |
|  | 450 | 45.42 | 45.04 | 45.15 | 45.34 | 45.34 | 45.26 |
|  | 300 | 41.54 | 40.96 | 41.34 | 41.34 | 41.54 | 41.34 |
| 2190-14-1 | 400 | 41.34 | 40.77 | 40.77 | 41.26 | 40.96 | 41.02 |
|  | 450 | 41.15 | 40.77 | 40.96 | 40.96 | 41.15 | 41.00 |
|  | 300 | 27.98 | 27.94 | 28.15 | 28.11 | 27.98 | 28.03 |
|  | 400 | 27.52 | 27.64 | 27.14 | 27.48 | 27.14 | 27.38 |
|  | 450 | 27.31 | 26.98 | 27.14 | 26.68 | 26.94 | 27.01 |
|  | 300 | 48.02 | 47.42 | 47.12 | 48.19 | 48.32 | 47.81 |
| -137-7-R1 | 400 | 44.75 | 44.16 | 43.86 | 44.16 | 44.45 | 44.27 |
|  | 450 | 43.86 | 43.99 | 43.40 | 43.57 | 43.40 | 43.64 |
|  | 300 | 34.99 | 34.75 | 34.99 | 34.99 | 35.23 | 34.99 |
| J2-222-1-R1 | 400 | 34.22 | 34.04 | 34.22 | 33.87 | 34.00 | 34.07 |
| Core A1 | 450 | 34.04 | 33.69 | 33.33 | 33.69 | 33.82 | 33.72 |
|  | 300 | 36.35 | 35.87 | 35.87 | 36.01 | 35.77 | 35.97 |
|  | 400 | 36.25 | 35.52 | 35.39 | 35.63 | 35.34 | 35.63 |
| Core A2 | 450 | 36.15 | 35.21 | 35.39 | 35.39 | 35.87 | 35.60 |
|  | 300 | 30.01 | 30.26 | 29.51 | 29.76 | 29.26 | 29.76 |
|  | 400 | 29.76 | 29.26 | 29.51 | 29.26 | 29.07 | 29.37 |
| Core Ex | 450 | 29.76 | 29.26 | 29.01 | 29.07 | 29.26 | 29.27 |


| J2-127-1-R2 | 300 | 37.85 | 37.53 | 37.31 | 37.31 | 37.19 | 37.44 |
| :---: | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
|  | 400 | 37.22 | 36.87 | 36.87 | 37.44 | 36.87 | 37.05 |
|  | 450 | 37.09 | 36.43 | 36.65 | 36.87 | 36.78 | 36.76 |
| J2-127-1-R2 | 300 | 36.25 | 36.16 | 36.09 | 36.09 | 36.25 | 36.16 |
|  | 400 | 35.67 | 35.60 | 35.50 | 35.57 | 35.92 | 35.65 |
|  | 450 | 35.43 | 35.60 | 35.50 | 35.50 | 35.50 | 35.51 |

## A3.4 Black Smoker Chimney Data

| Sample | Confining Pressure (psi) | 1 | 2 | rosity (\%) 3 | 4 | 5 | Mean $\phi$ (\%) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{aligned} & \text { ALV 1445-3 } \\ & \text { Core C1 } \end{aligned}$ | 300 | 24.839 | 24.882 | 25.223 | 24.882 | 24.882 | 24.94 |
|  | 400 | 24.070 | 24.691 | 23.537 | 23.876 | 23.876 | 24.01 |
|  | 450 | 23.493 | 22.918 | 23.494 | 22.581 | 23.494 | 23.19 |
| $\begin{gathered} \text { ALV 2179-4-1 } \\ \text { Core A1 } \end{gathered}$ | 300 | 42.799 | 43.004 | 43.336 | 41.995 | 42.594 | 42.74 |
|  | 400 | 41.933 | 41.728 | 41.256 | 41.194 | 41.131 | 41.45 |
|  | 450 | 40.722 | 41.194 | 40.722 | 40.926 | 41.461 | 41.00 |
| $\begin{aligned} & \text { ALV 2179-4-1 } \\ & \text { Core B1 } \end{aligned}$ | 300 | 36.794 | 36.746 | 36.488 | 36.376 | 36.536 | 36.59 |
|  | 400 | 35.288 | 34.872 | 35.080 | 34.458 | 35.080 | 34.95 |
|  | 450 | 34.408 | 34.042 | 34.042 | 34.408 | 34.250 | 34.23 |
| $\begin{aligned} & \text { ALV 2179-4-1 } \\ & \text { Core C1 } \end{aligned}$ | 300 | 42.171 | 42.338 | 41.902 | 42.338 | 42.338 | 42.22 |
|  | 400 | 41.416 | 41.468 | 41.685 | 41.251 | 41.034 | 41.37 |
|  | 450 | 40.599 | 40.817 | 40.217 | 40.432 | 40.599 | 40.53 |
| $\begin{gathered} \text { J2-137-1-R1 } \\ \text { Core D1 } \end{gathered}$ | 300 | 17.085 | 16.700 | 16.827 | 17.085 | 16.926 | 16.92 |
|  | 400 | 16.926 | 16.669 | 16.542 | 16.926 | 16.315 | 16.67 |
|  | 450 | 16.542 | 16.700 | 16.542 | 16.444 | 16.413 | 16.53 |
| $\begin{gathered} \text { J2-137-1-R1 } \\ \text { Core D2 } \end{gathered}$ | 300 | 14.025 | 13.871 | 13.650 | 13.996 | 14.150 | 13.94 |
|  | 400 | 14.025 | 14.120 | 13.899 | 13.431 | 13.775 | 13.85 |
|  | 450 | 13.899 | 13.526 | 13.526 | 13.526 | 13.526 | 13.60 |
| $\begin{gathered} \text { J2-137-1-R1 } \\ \text { Core D3 } \end{gathered}$ | 300 | 18.255 | 18.000 | 18.127 | 18.098 | 17.746 | 18.04 |
|  | 400 | 17.873 | 17.970 | 17.619 | 17.843 | 17.716 | 17.80 |
|  | 450 | 17.843 | 17.716 | 17.619 | 17.619 | 17.619 | 17.68 |
| $\begin{gathered} \text { J2-137-1-R1 } \\ \text { Core D4 } \end{gathered}$ | 300 | 17.749 | 17.621 | 17.749 | 17.621 | 18.005 | 17.75 |
|  | 400 | 17.877 | 17.237 | 17.237 | 17.364 | 17.364 | 17.41 |
|  | 450 | 17.621 | 17.109 | 17.109 | 17.335 | 17.237 | 17.28 |
| $\begin{gathered} \text { J2-213-3-R1 } \\ \text { Core A1 } \end{gathered}$ | 300 | 29.020 | 29.337 | 28.755 | 28.755 | 28.797 | 28.93 |
|  | 400 | 29.295 | 28.396 | 28.438 | 28.618 | 28.438 | 28.64 |
|  | 450 | 28.576 | 28.618 | 28.438 | 28.438 | 28.438 | 28.50 |
| $\begin{gathered} \text { J2-213-3-R1 } \\ \text { Core B1 } \end{gathered}$ | 300 | 30.507 | 30.549 | 30.507 | 30.689 | 30.549 | 30.56 |
|  | 400 | 30.369 | 30.005 | 30.005 | 30.005 | 30.369 | 30.15 |
|  | 450 | 29.825 | 30.145 | 30.145 | 30.187 | 30.005 | 30.06 |
| $\begin{gathered} \text { J2-213-3-R1 } \\ \text { Core D1 } \end{gathered}$ | 300 | 39.484 | 38.969 | 38.969 | 38.713 | 38.969 | 39.02 |
|  | 400 | 39.117 | 38.713 | 38.199 | 38.199 | 38.861 | 38.62 |
|  | 450 | 38.713 | 38.457 | 38.603 | 37.942 | 38.861 | 38.51 |

## A3.5 Relict Spire Data

|  | Confining Pressure (psi) | Porosity (\%) |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| ple |  | 1 | 2 | 3 | 4 | 5 | (\%) |
| J2-129-1-R3 | 300 | 41.124 | 41.128 | 40.946 | 40.946 | 40.946 | 41.02 |
| J-129-1-R3 Core B 2 | 400 | 40.227 | 40.156 | 40.227 | 40.227 | 40.156 | 40.20 |
| Core B2 | 450 | 39.941 | 39.941 | 39.903 | 39.870 | 40.013 | 39.93 |
| $\begin{gathered} \text { J2-136-6-R1 } \\ \text { Core A1-1 } \end{gathered}$ | 300 | 34.964 | 30.802 | 25.403 | 31.456 | 34.744 | 31.27 |
|  | 400 | 33.644 | 29.065 | 27.336 | 29.715 | 30.802 | 30.04 |
|  | 450 | 33.329 | 29.621 | 31.579 | 28.755 | 29.498 | 30.51 |
| J2-136-6-R1Core A1-2 | 300 | 30.960 | 30.433 | 30.257 | 29.480 | 30.006 | 30.22 |
|  | 400 | 30.359 | 30.359 | 29.480 | 28.858 | 30.183 | 29.84 |
|  | 450 | 30.006 | 28.432 | 29.132 | 29.308 | 30.006 | 29.37 |
| $\begin{gathered} \text { J2-136-6-R1 } \\ \text { Core C2 } \end{gathered}$ | 300 | 27.547 | 24.719 | 25.103 | 24.847 | 24.975 | 25.42 |
|  | 400 | 26.514 | 23.827 | 24.464 | 24.591 | 24.791 | 24.82 |
|  | 450 | 26.771 | 24.663 | 24.591 | 25.103 | 24.719 | 25.16 |
| $\begin{gathered} \hline \text { ALV 2944-3-S1 } \\ \text { Pc } 1 \\ \text { Core A1 } \\ \hline \end{gathered}$ | 300 | 43.153 | 43.513 | 33.599 | 34.720 | 34.390 | 37.62 |
|  | 400 | 41.929 | 35.448 | 34.720 | 34.982 | 34.920 | 36.30 |
|  | 450 | 41.570 | 36.774 | 30.705 | 32.544 | 31.555 | 34.40 |
| $\begin{gathered} \text { ALV 2944-3-S1 } \\ \text { Pc } 1 \\ \text { Core A3 } \\ \hline \end{gathered}$ | 300 | 29.564 | 34.089 | 34.089 | 28.905 | 28.340 | 30.89 |
|  | 400 | 33.547 | 33.547 | 29.439 | 28.684 | 30.886 | 31.16 |
|  | 450 | 32.658 | 30.224 | 28.684 | 28.905 | 29.439 | 29.95 |
| $\begin{gathered} \text { ALV 2944-3-S1 } \\ \text { Pc } 1 \\ \text { Core B1 } \end{gathered}$ | 300 | 40.307 | 39.444 | 37.771 | 31.292 | 31.978 | 35.96 |
|  | 400 | 36.769 | 33.338 | 31.729 | 37.771 | 33.775 | 34.60 |
|  | 450 | 36.769 | 26.120 | 27.592 | 27.101 | 30.493 | 29.38 |
| $\begin{gathered} \text { ALV 2944-3-S1 } \\ \text { Pc } 2 \\ \text { Core A1 } \end{gathered}$ | 300 | 32.519 | 32.250 | 38.747 | 31.981 | 32.455 | 33.50 |
|  | 400 | 39.836 | 28.430 | 32.519 | 27.895 | 28.430 | 31.12 |
|  | 450 | 38.138 | 29.232 | 30.778 | 32.519 | 33.737 | 32.74 |
| $\begin{gathered} \text { ALV } 2944-3-\mathrm{S} 1 \\ \text { Pc } 2 \\ \text { Core A2 } \end{gathered}$ | 300 | 42.181 | 42.083 | 43.220 | 39.818 | 45.966 | 42.61 |
|  | 400 | 44.589 | 42.310 | 42.439 | 40.624 | 40.496 | 42.07 |
|  | 450 | 43.805 | 38.594 | 39.495 | 38.370 | 40.624 | 40.13 |
| $\begin{gathered} \text { ALV 2944-3-S1 } \\ \text { Pc } 2 \\ \text { Core A3 } \end{gathered}$ | 300 | 39.539 | 41.847 | 40.197 | 45.493 | 36.588 | 40.63 |
|  | 400 | 47.494 | 40.384 | 41.517 | 40.384 | 40.384 | 41.95 |
|  | 450 | 47.351 | 41.517 | 41.186 | 44.498 | 38.882 | 42.59 |
| $\begin{gathered} \hline \text { ALV } 2944-3-\mathrm{S} 1 \\ \text { Pc } 2 \\ \text { Core A4 } \\ \hline \end{gathered}$ | 300 | 35.785 | 30.412 | 31.574 | 29.891 | 30.876 | 31.64 |
|  | 400 | 35.025 | 29.661 | 27.863 | 28.732 | 28.270 | 29.80 |
|  | 450 | 36.278 | 32.918 | 30.356 | 30.589 | 31.053 | 32.17 |
| $\begin{gathered} \text { ALV 2941-6-S1 } \\ \text { Core A1 } \end{gathered}$ | 300 | 34.089 | 29.885 | 32.331 | 30.972 | 35.340 | 32.46 |
|  | 400 | 35.890 | 32.176 | 29.208 | 31.786 | 34.244 | 32.58 |
|  | 450 | 34.518 | 31.397 | 27.992 | 28.599 | 28.754 | 30.16 |
| $\begin{gathered} \mathrm{J} 2-125-3-\mathrm{B} 1 \\ \text { Core B2-1 } \end{gathered}$ | 300 | 36.834 | 33.842 | 33.416 | 36.498 | 33.295 | 34.74 |
|  | 400 | 35.641 | 36.069 | 33.842 | 35.641 | 33.628 | 34.95 |
|  | 450 | 36.405 | 33.508 | 32.659 | 35.426 | 33.203 | 34.21 |
| $\begin{gathered} \text { J2-125-3-B1 } \\ \text { Core B2-2 } \end{gathered}$ | 300 | 36.501 | 36.245 | 31.005 | 36.612 | 37.126 | 35.42 |
|  | 400 | 35.843 | 29.993 | 29.633 | 34.964 | 33.688 | 32.72 |
|  | 450 | 35.074 | 33.290 | 35.221 | 35.221 | 34.199 | 34.59 |
| $\begin{gathered} \text { J2-125-3-B1 } \\ \text { Core B3-1 } \end{gathered}$ | 300 | 36.058 | 31.102 | 31.405 | 32.864 | 31.982 | 32.64 |
|  | 400 | 35.194 | 31.514 | 32.092 | 31.020 | 33.058 | 32.54 |
|  | 450 | 34.610 | 34.500 | 32.755 | 34.221 | 35.279 | 34.26 |
| $\begin{gathered} \text { J2-125-3-B1 } \\ \text { Core B3-2 } \end{gathered}$ | 300 | 35.960 | 34.149 | 32.618 | 35.002 | 35.583 | 34.64 |
|  | 400 | 34.285 | 33.808 | 29.308 | 30.015 | 33.569 | 32.13 |
|  | 450 | 34.149 | 32.144 | 29.308 | 30.250 | 34.659 | 32.03 |
| $\begin{aligned} & \text { J2-125-3-B1 } \\ & \text { Core Ex B2 } \end{aligned}$ | 300 | 39.833 | 36.423 | 38.571 | 37.272 | 36.990 | 37.80 |
|  | 400 | 39.139 | 38.287 | 37.151 | 36.141 | 34.889 | 37.09 |
|  | 450 | 38.287 | 36.301 | 37.841 | 29.910 | 34.044 | 35.14 |
| $\begin{aligned} & \text { ALV 2178-4-1 } \\ & \text { Core } 4 \end{aligned}$ | 300 | 49.742 | 49.339 | 49.206 | 49.206 | 49.206 | 49.34 |
|  | 400 | 47.613 | 47.685 | 47.480 | 47.420 | 47.420 | 47.52 |
|  | 450 | 46.894 | 46.894 | 46.762 | 46.894 | 46.894 | 46.87 |
| $\begin{aligned} & \text { ALV 2178-4-1 } \\ & \quad \text { Core } 5 \end{aligned}$ | 300 | 45.628 | 45.706 | 45.706 | 45.569 | 45.628 | 45.65 |
|  | 400 | 44.886 | 44.886 | 44.749 | 44.964 | 44.749 | 44.85 |
|  | 450 | 44.344 | 44.418 | 44.208 | 44.149 | 44.149 | 44.25 |
| $\begin{gathered} \text { ALV 2178-4-1 } \\ \text { Core A4 } \end{gathered}$ | 300 | 39.885 | 39.620 | 39.620 | 39.620 | 39.346 | 39.62 |
|  | 400 | 39.081 | 38.815 | 38.665 | 38.665 | 39.081 | 38.86 |
|  | 450 | 38.815 | 39.081 | 38.392 | 38.542 | 38.665 | 38.70 |


| $\begin{gathered} \text { ALV 2178-4-1 } \\ \text { Core A5 } \end{gathered}$ | 300 | 45.057 | 44.940 | 45.057 | 44.784 | 44.504 | 44.87 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 400 | 44.386 | 44.105 | 44.105 | 44.105 | 44.105 | 44.16 |
|  | 450 | 43.831 | 44.105 | 43.831 | 43.551 | 43.551 | 43.77 |
| ALV 2461-R13 Core 1-1 | 300 | 26.730 | 26.664 | 26.767 | 26.767 | 26.700 | 26.73 |
|  | 400 | 26.531 | 26.700 | 26.567 | 26.567 | 26.634 | 26.60 |
|  | 450 | 26.567 | 26.737 | 26.567 | 26.435 | -- | 26.58 |
| ALV 2461-R13 Core 1-2 | 300 | 22.163 | 22.233 | 22.233 | 22.201 | 22.092 | 22.18 |
|  | 400 | 22.131 | 22.061 | 22.061 | 22.030 | 21.991 | 22.06 |
|  | 450 | 22.131 | 21.991 | 21.921 | 22.061 | 22.100 | 22.04 |
| $\begin{gathered} \text { ALV 2461-R13 } \\ \text { Core } 2 \end{gathered}$ | 300 | 28.991 | 28.962 | 28.835 | 28.801 | 28.898 | 28.90 |
|  | 400 | 28.835 | 28.611 | 28.898 | 28.582 | 28.645 | 28.71 |
|  | 450 | -- | -- | -- | -- | -- | -- |
| ALV 2461-R13 Core 3-1 | 300 | 17.353 | 17.528 | 17.403 | 17.003 | 17.353 | 17.33 |
|  | 400 | 17.178 | 16.828 | 16.916 | 16.741 | 17.003 | 16.93 |
|  | 450 | 16.916 | 16.741 | 16.878 | 16.703 | 16.578 | 16.76 |
| ALV 2461-R13 <br> Core 3-2(1-1) | 300 | 6.476 | 4.809 | 4.484 | 3.609 | 6.367 | 5.03 |
|  | 400 | 10.405 | 4.894 | 5.999 | 11.196 | 10.077 | 8.08 |
|  | 450 | 10.866 | 8.589 | 8.961 | 7.520 | 6.738 | 8.42 |
| ALV 2461-R13 <br> Core 3-2(1-2) | 300 | 12.820 | 9.885 | 8.243 | 11.166 | 9.842 | 10.23 |
|  | 400 | 12.820 | 11.124 | 9.294 | 10.390 | 8.607 | 10.35 |
|  | 450 | 14.296 | 11.716 | 10.068 | 9.477 | 8.971 | 10.75 |
| ALV 2461-R13 Core 3-2(2) | 300 | 12.527 | 12.404 | 12.280 | 12.086 | 11.755 | 12.21 |
|  | 400 | 12.651 | 11.940 | 9.916 | 10.925 | 9.732 | 10.98 |
|  | 450 | 12.033 | 11.571 | 12.227 | 10.374 | 11.848 | 11.59 |
| $\begin{aligned} & \text { ALV 2461-R13 } \\ & \text { Core 4-1 } \end{aligned}$ | 300 | 19.582 | 18.698 | 17.729 | 19.526 | 18.185 | 18.73 |
|  | 400 | 19.471 | 18.698 | 17.673 | 17.418 | 19.084 | 18.45 |
|  | 450 | 19.212 | 17.929 | 18.642 | 18.827 | 18.185 | 18.55 |
| $\begin{aligned} & \text { ALV 2461-R13 } \\ & \text { Core 4-2 } \end{aligned}$ | 300 | 22.637 | 22.637 | 22.445 | 21.971 | 20.458 | 22.01 |
|  | 400 | 22.237 | 20.854 | 19.668 | 22.237 | 22.179 | 21.41 |
|  | 450 | 22.503 | 21.118 | 20.326 | 21.971 | 22.445 | 21.66 |
| ALV 2461-R13 Core 6 | 300 | 17.543 | 17.335 | 17.624 | 17.497 | 17.452 | 17.49 |
|  | 400 | 17.497 | 17.254 | 17.335 | 17.254 | 17.254 | 17.32 |
|  | 450 | 17.579 | 17.173 | 17.254 | 17.497 | 17.012 | 17.30 |
| ALV 2461-R13Core 7 | 300 | 23.383 | 24.596 | 24.692 | 24.692 | 24.692 | 24.41 |
|  | 400 | 24.692 | 24.692 | 24.747 | 24.692 | 24.596 | 24.68 |
|  | 450 | 24.498 | 24.692 | 24.596 | 24.596 | 24.498 | 24.58 |
| ALV 2461-R13 Core 8-1 | 300 | 20.783 | 20.606 | 20.645 | 20.783 | 20.683 | 20.70 |
|  | 400 | 20.556 | 20.606 | 20.606 | 20.645 | 20.618 | 20.61 |
|  | 450 | 20.783 | 20.694 | 20.556 | 20.694 | 20.468 | 20.64 |
| ALV 2461-R13 Core 8-2 | 300 | 20.701 | 20.663 | 20.663 | 20.663 | 20.750 | 20.69 |
|  | 400 | 20.528 | 20.490 | 20.528 | 20.490 | 20.501 | 20.51 |
|  | 450 | 20.490 | 20.317 | 20.403 | 20.403 | 20.403 | 20.40 |

## Appendix 4: Microstructure Tables

Mineral abbreviations used in microstructure tables:

| Mineral | Abbreviation |
| :--- | :---: |
| amorphous silica | am Si |
| anhydrite | anh |
| barite | brt |
| calcite | ca |
| chalcocite | ch |
| chalcopyrite | cp |
| clay | cl |
| covellite | co |
| plagioclase | pl |
| palagonatized glass | gl |
| pyrite | py |
| pyrrhotite | po |
| sphalerite | sp |
| stevensite | st |
| wurtzite | wz |

## A4.1 Massive Anhydrite Data

| Sample / Section | $\underset{\left(\times 10^{-15} \mathrm{~m}^{2}\right)}{k}$ | $\begin{gathered} \phi \\ (\%) \end{gathered}$ | Minerals Present | Grain Packing \& Size | Pore Size | Pore Connectivity | Channel Width min-max (avg) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{gathered} J 2-210-8-R 2 \\ \mathrm{~A}-1 \end{gathered}$ | 0.3 | 2.3 | anh (py, cp) | tight; 150-1500 $\mu \mathrm{m}$ (avg: $1000 \mu \mathrm{~m}$ ) sized anh crystals; sulfides 20-150 $\mu \mathrm{m}$ | most pores $60-150 \mu \mathrm{~m}$ (avg: $90 \mu \mathrm{~m}$ ); sulfide crystals precipitated in pore space | low | $10-80 \mu \mathrm{~m}(20 \mu \mathrm{~m})$ |
| $\begin{gathered} J 2-216-14-R 1 \\ \text { A } \\ \text { B (top) } \\ \text { B (bottom) } \end{gathered}$ | $\begin{aligned} & 0.3 \\ & 0.3 \\ & 0.3 \end{aligned}$ | 2.1 <br> 5.6 <br> 5.6 | anh, cp (py, sp) <br> anh, cp (py, sp) <br> anh, cp (py, sp) | tight; areas of large anh 20-2500 $\mu \mathrm{m}$ (avg: $450 \mu \mathrm{~m}$ ); patches of small, euhedral anh crystals $\sim 150 \mu \mathrm{~m}$ <br> tight; areas of large anh $20-1500 \mu \mathrm{~m}$ (avg: $400 \mu \mathrm{~m}$ ); patches of small euhedral anh crystals $\sim 150 \mu \mathrm{~m}$ <br> tight; areas of large anh $20-1600 \mu \mathrm{~m}$ (avg: $400 \mu \mathrm{~m}$ ); patches of small euhedral anh crystals $\sim 100 \mu \mathrm{~m}$ | pores ranging 10-200 $\mu \mathrm{m}$ (avg: $60 \mu \mathrm{~m}$ ) pores ranging 10-200 $\mu \mathrm{m}$ (avg: $60 \mu \mathrm{~m}$ ) pores ranging $10-250 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ) | low <br> low <br> low | $<40 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ $<40 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ $<40 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ |
| $\begin{gathered} \text { J301-3 } \\ \text { A } \\ \text { B } \end{gathered}$ | $\begin{gathered} 38.5 \\ 0.7 \end{gathered}$ | $\begin{gathered} 14.8 \\ 7.9 \end{gathered}$ | anh (py ) <br> anh (py) | moderate; anh (100-1300 $\mu \mathrm{m}$ ) with minor py crystals $20-80 \mu \mathrm{~m}$; anh crystals highly fragmented tight; anh $100-1400 \mu \mathrm{~m}$ (avg: $400 \mu \mathrm{~m}$ ); py crystals $20-150 \mu \mathrm{~m}$; anh crystals highly fragmented | pores ranging $50-500 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) <br> pores ranging $50-500 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) | moderate <br> low | $\begin{aligned} & 10-80 \mu \mathrm{~m}(25 \mu \mathrm{~m}) \\ & 10-50 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \end{aligned}$ |
| $\begin{gathered} \text { ALV } 21837-0 \\ \text { B } \end{gathered}$ | 0.3 | 6.6 | anh (py) | tight; anh $50-1400 \mu \mathrm{~m}$ (avg: $900 \mu \mathrm{~m}$ ); also patches of small euhedral anh crystals $\sim 50-100 \mu \mathrm{~m}$ | pores range from 30-200 $\mu \mathrm{m}$ (avg: $50 \mu \mathrm{~m}$ ); most pores isolated | low | 10-60 $\mu \mathrm{m}(20 \mu \mathrm{~m})$ |
| $\begin{gathered} \text { ALV 2581-8 } \\ \mathrm{A} \end{gathered}$ | 352.6 | 12.7 | anh (py, cp) | moderate; 200-1300 $\mu \mathrm{m}$ sized anh crystals; sulfide crystals $<50 \mu \mathrm{~m}$ | most pores $60-150 \mu \mathrm{~m}$ (avg: $90 \mu \mathrm{~m}$ ); small sulfide crystals precipitated in pore space | high | 20-120 $\mu \mathrm{m}(60 \mu \mathrm{~m})$ |
| MIR 1, 1/74, Sta 2403 <br> A-1 <br> B | $\begin{array}{r} 123.2 \\ 77.1 \end{array}$ | $\begin{aligned} & 10.9 \\ & 9.0 \end{aligned}$ | $\begin{aligned} & \text { anh, py, cp } \\ & \text { anh, py, cp } \end{aligned}$ | loose; anh crystals $100-1700 \mu \mathrm{~m}$ (avg: $500 \mu \mathrm{~m}$ ); py and cp grains < $100 \mu \mathrm{~m}$ <br> moderate; anh crystals $100-1400 \mu \mathrm{~m}$ (avg: $400 \mu \mathrm{~m}$ ); py and cp < $100 \mu \mathrm{~m}$ | most pores $50-300 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ); some sulfide clusters in pore space <br> most pores $20-300 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ); some sulfide clusters in pore space | $\begin{gathered} \text { high } \\ \text { moderate } \end{gathered}$ | $\begin{aligned} & 20-100 \mu \mathrm{~m}(70 \mu \mathrm{~m}) \\ & 10-100 \mu \mathrm{~m}(40 \mu \mathrm{~m}) \end{aligned}$ |
| $\begin{aligned} & \text { MIR 1, 2/78, Sta } 2417 \\ & \text { A } \\ & \text { B } \end{aligned}$ | $\begin{aligned} & 477.9 \\ & 595.9 \end{aligned}$ | $\begin{aligned} & 15.0 \\ & 13.5 \end{aligned}$ | anh (cp) <br> anh (cp) | loose; anh crystals $100-1200 \mu \mathrm{~m}$ (avg: $500 \mu \mathrm{~m}$ ); cp grains $<80 \mu \mathrm{~m}$ <br> loose; anh crystals 100-1200 $\mu \mathrm{m}$ (avg: $400 \mu \mathrm{~m}$ ); cp < $100 \mu \mathrm{~m}$ | pores ranging $50-300 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) pores ranging $50-300 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) | high <br> high | $\begin{aligned} & 10-150 \mu \mathrm{~m}(60 \mu \mathrm{~m}) \\ & 10-150 \mu \mathrm{~m}(60 \mu \mathrm{~m}) \end{aligned}$ |

## A4.2 Flange, Slab and Crust Data

| Sample / Section | $\begin{gathered} k \\ \left(\times 10^{-15} \mathrm{~m}^{2}\right) \end{gathered}$ | $\begin{gathered} \phi \\ (\%) \end{gathered}$ | Minerals Present | Grain Packing \& Size | Pore Size | Pore Connectivity | Channel Width min-max (avg) | Section Cut |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{gathered} \text { ALV 3517-R1 } \\ \text { D1 } \\ \text { D3 } \end{gathered}$ | $\begin{aligned} & 55.1 \\ & 29.5 \end{aligned}$ | 22.1 20.2 | $\mathrm{ca}, \mathrm{st}(\mathrm{cp}, \mathrm{sp})$ $\mathrm{ca}, \mathrm{st}(\mathrm{cp}, \mathrm{sp})$ | tight; patches of large ca crystals (250 $\mu \mathrm{m}$ ), but mostly $\sim 50 \mu \mathrm{~m}$ crystals <br> tight; patches of large ca crystals (250 $\mu \mathrm{m}$ ), but mostly $\sim 50 \mu \mathrm{~m}$ crystals | some $\sim 350 \mu \mathrm{~m}$ pores around smaller crystals; mostly $\sim 60 \mu \mathrm{~m}$ <br> some $\sim 350 \mu \mathrm{~m}$ pores around smaller crystals; mostly $\sim 60 \mu \mathrm{~m}$ | low <br> low | $\begin{aligned} & <20 \mu \mathrm{~m}(10 \mu \mathrm{~m}) \\ & <20 \mu \mathrm{~m}(10 \mu \mathrm{~m}) \end{aligned}$ | (2) (3) |
| $\begin{gathered} \text { ALV 3521-R2 } \\ \text { Ex } 1 \\ \text { Ex } 2 \end{gathered}$ | $\begin{gathered} 1243.4 \\ 954.7 \end{gathered}$ | 45.3 45.9 | ca, st (brt, cp, sp) <br> ca, st (brt, cp, sp) | tight; large areas of small $(\sim 50 \mu \mathrm{~m})$ ca crystals, few patches of larger $\sim 100 \mu \mathrm{~m}$ tight; large areas of small $(\sim 50 \mu \mathrm{~m}) \mathrm{ca}$ crystals, few patches of larger $\sim 100 \mu \mathrm{~m}$ | most pores $20-150 \mu \mathrm{~m}$ regularly distributed, some larger $\sim 300 \mu \mathrm{~m}$ pores <br> most pores $20-150 \mu \mathrm{~m}$ regularly distributed, some larger $\sim 300 \mu \mathrm{~m}$ pores | moderate <br> moderate | $<40 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ $<40 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ | (2) (3) |
| $\begin{gathered} \hline A L V 2927-3 \\ \text { B1 } \\ 3 \\ 2 \end{gathered}$ | $\begin{gathered} 1889.1 \\ 967.3 \\ 148.7 \end{gathered}$ | $\begin{aligned} & 38.2 \\ & 40.7 \\ & 32.7 \end{aligned}$ | $\begin{aligned} & \text { py, wz (po, cp, cl) } \\ & \text { py, wz (po, cp, cl) } \\ & \text { py, wz (po, cp, cl) } \end{aligned}$ | loose; small ( $20-125 \mu \mathrm{~m}$ ) sulfide crystals with small amounts of clay <br> moderate; small (20-125 $\mu \mathrm{m}$ ) sulfide crystals with small amounts of clay <br> tight-moderate; small $(5-130 \mu \mathrm{~m})$ sulfide crystals with small amounts of clay | irregularly shaped pores ranging $40-350 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ); smaller isolated pores irregularly shaped pores ranging 40-200 $\mu \mathrm{m}$ (avg: $80 \mu \mathrm{~m}$ ); minor isolated pores <br> irregularly shaped pores ranging $30-250 \mu \mathrm{~m}$ (avg: $50 \mu \mathrm{~m}$ ); minor isolated pores | high <br> high <br> moderate | $\begin{aligned} & 10-75 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \\ & 10-180 \mu \mathrm{~m}(40 \mu \mathrm{~m}) \\ & 10-70 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \end{aligned}$ | (1) <br> (3) <br> (2) |
| $\begin{gathered} A L V 2415-1 B \\ \text { B1 } \\ \text { A1 } \end{gathered}$ | $\begin{array}{r} 1552.3 \\ 55.4 \end{array}$ | 31.2 20.2 | $\begin{aligned} & \text { py, po (am Si, cp, wz) } \\ & \text { py, po, am Si (cp, wz) } \end{aligned}$ | moderate; $10-300 \mu \mathrm{~m}$ (avg: $75 \mu \mathrm{~m}$ ) sized, poorly sorted sulfide crystals <br> tight; $10-280 \mu \mathrm{~m}$ (avg: $75 \mu \mathrm{~m}$ ) sized sulfide crystals; lots of am Si | pores ranging from $30-250 \mu \mathrm{~m}$ (avg: $50 \mu \mathrm{~m}$ ); some am Si restricting pore channels <br> pores ranging from 30-200 $\mu \mathrm{m}$ (avg: $75 \mu \mathrm{~m}$ ); abundant am Si restricting pore channels | moderate <br> (high in 1 <br> layer) <br> low | $\begin{aligned} & 10-50 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \\ & (20-130 \mu \mathrm{~m}(30 \mu \mathrm{~m})) \\ & 10-50 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \end{aligned}$ | (1) (2) |
| $J 2-286$ C3-1 <br> C3-2 <br> C2-1 <br> A1 | $\begin{array}{r} 146.0 \\ 48.7 \\ 0.8 \\ 0.5 \end{array}$ | $\begin{aligned} & 27.6 \\ & 21.2 \\ & 27.2 \\ & 20.7 \end{aligned}$ | $\begin{aligned} & \text { py, wz, po (am Si) } \\ & \text { py, am Si (wz, cp) } \\ & \text { py, wz (po) } \\ & \text { py, po, am Si (wz) } \end{aligned}$ | tight; $10-50 \mu \mathrm{~m}$ sized sulfide crystals; large bladed po crystals ( $200-3000 \mu \mathrm{~m}$ ) tight; 20-200 $\mu \mathrm{m}$ sized sulfide crystals with lots of am Si around crystals <br> moderate; $10-175 \mu \mathrm{~m}$ sulfide crystals <br> moderate; $20-200 \mu \mathrm{~m}$ (avg: $50 \mu \mathrm{~m}$ ) sized sulfide crystals; lots of am Si | pores ranging from $20-300 \mu \mathrm{~m}$ (avg: $20 \mu \mathrm{~m}$ ); remnant worm tubes and possible flange edge pores ranging from $10-400 \mu \mathrm{~m}$ (avg: $20 \mu \mathrm{~m}$ ); lots of remnant worm tubes <br> pores ranging from $10-250 \mu \mathrm{~m}$ (avg: $75 \mu \mathrm{~m}$ ); remnant worm tubes <br> pores ranging from $10-300 \mu \mathrm{~m}$ (avg: $75 \mu \mathrm{~m}$ ); remnant worm tubes |  | $\begin{aligned} & 10-50 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \\ & 10-50 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \\ & 10-100 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \\ & 10-100 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \end{aligned}$ | (2) <br> (2) <br> (3) <br> (2) |


| $\begin{gathered} \hline A L V 2608-3-3 \\ \text { B2 } \\ \text { C1 } \end{gathered}$ | $2737.1$ <br> 1447.1 | $\begin{aligned} & 42.9 \\ & 29.2 \end{aligned}$ | $\mathrm{gl}(\mathrm{pl}, \mathrm{am} \mathrm{Si}, \mathrm{py})$ <br> gl, am Si (pl, py) | loose but variable packing; poorly sorted gl shards $50-1400 \mu \mathrm{~m}$ (avg: $300 \mu \mathrm{~m}$ ) <br> moderate, but variable; poorly sorted gl shards $30-1000 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) | very porous; pores ranging from $50-800 \mu \mathrm{~m}$ (avg: $175 \mu \mathrm{~m}$ ) <br> very porous, but connectivity lost to am Si; pores range from $50-600 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) | high <br> moderate | $10-200 \mu \mathrm{~m}(40 \mu \mathrm{~m})$ <br> $10-50 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ | (1) <br> (1) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{aligned} & \text { ALV 2608-4-1, Pc 1 } \\ & \text { A1 } \\ & \text { C3-2 } \\ & \text { C3-1 } \end{aligned}$ | $\begin{array}{r} 740.6 \\ 111.7 \\ 24.1 \end{array}$ | $\begin{aligned} & 34.9 \\ & 37.2 \\ & 35.0 \end{aligned}$ | $\mathrm{gl}, \mathrm{am} \mathrm{Si}, \mathrm{cl}(\mathrm{pl})$ <br> $\mathrm{gl}, \mathrm{am} \mathrm{Si}, \mathrm{cl}(\mathrm{pl})$ <br> $\mathrm{gl}, \mathrm{am} \mathrm{Si}, \mathrm{cl}(\mathrm{pl})$ | tight; poorly sorted, angular gl shards 50-200 $\mu \mathrm{m}$ (avg: $80 \mu \mathrm{~m}$ ) and am Si $\pm \mathrm{cl}$ <br> tight; poorly sorted, angular gl shards $50-200 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ) and am Si $\pm \mathrm{cl}$ <br> moderate; poorly sorted, angular gl shards $50-300 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) | pores range from $10-400 \mu \mathrm{~m}$ (avg: $30 \mu \mathrm{~m}$ ); some channels blocked by am Si $\pm \mathrm{cl}$ <br> pores range from 20-500 $\mu \mathrm{m}$ (avg: $60 \mu \mathrm{~m}$ ); few channels blocked by am Si $\pm$ cl <br> pores range from $20-500 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ); few channels blocked by am $\mathrm{Si} \pm \mathrm{cl}$ | $\begin{gathered} \text { low } \\ \text { moderate } \\ \substack{\text { moderate } \\ \text { (very low in } \\ 1 \text { layer) }} \end{gathered}$ | $\begin{aligned} & 10-40 \mu \mathrm{~m}(15 \mu \mathrm{~m}) \\ & 10-80 \mu \mathrm{~m}(20 \mu \mathrm{~m}) \\ & \begin{array}{c} 10-200 \mu \mathrm{~m}(50 \mu \mathrm{~m}) \\ (<30 \mu \mathrm{~m}(5 \mathrm{~m})) \end{array} \end{aligned}$ | (1) <br> (2) <br> (2) |
| $\begin{aligned} & \text { ALV 2608-4-1, Pc } 2 \\ & \text { A3 } \\ & 2 \\ & \text { B3 } \end{aligned}$ | $\begin{array}{r} 2016.9 \\ 574.6 \\ 538.3 \end{array}$ | $\begin{aligned} & 39.8 \\ & 48.0 \\ & 43.1 \end{aligned}$ | $\mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl})$ <br> $\mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl})$ <br> $\mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl})$ | moderate; angular gl shards $30-400 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) with am Si coating <br> tight; angular gl shards $50-200 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m})$ with am Si coating <br> tight; gl shards $50-200 \mu \mathrm{~m}$ (avg: 100 $\mu \mathrm{m})$ and $2000-3000 \mu \mathrm{~m}$ in some areas | very porous; pores ranging from $50-1000 \mu \mathrm{~m}$ (avg: $175 \mu \mathrm{~m}$ ) <br> very porous; pores ranging from $50-1000 \mu \mathrm{~m}$ (avg: $175 \mu \mathrm{~m}$ ); minor am Si $\pm \mathrm{cl}$ <br> very porous; pores ranging from $50-1000 \mu \mathrm{~m}$ (avg: $175 \mu \mathrm{~m}$ ); minor am Si $\pm \mathrm{cl}$ | high <br> high <br> high | 10-200 $\mu \mathrm{m}(40 \mu \mathrm{~m})$ <br> $10-200 \mu \mathrm{~m}(40 \mu \mathrm{~m})$ <br> $10-200 \mu \mathrm{~m}(40 \mu \mathrm{~m})$ | (1) <br> (2) <br> (2) |
| JAS 177-2-1 <br> B2 <br> B1 <br> A2 | $\begin{array}{r} 2954.2 \\ 1774.8 \\ 170.4 \end{array}$ | $\begin{aligned} & 41.2 \\ & 38.8 \\ & 46.1 \end{aligned}$ | $\mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl})$ <br> $\mathrm{gl}(\mathrm{am} \mathrm{Si}, \mathrm{cl}, \mathrm{pl})$ <br> $\mathrm{gl}, \mathrm{pl}(\mathrm{am} \mathrm{Si}, \mathrm{cl})$ | loose; angular gl shards $30-900 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) with am Si coating <br> loose; angular gl shards $30-900 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) with am Si coating <br> loose; angular gl shards $30-900 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) with am Si coating | very porous; pores ranging from $50-800 \mu \mathrm{~m}$ (avg: $175 \mu \mathrm{~m}$ ); minor am Si $\pm \mathrm{cl}$ <br> very porous; pores ranging from $50-800 \mu \mathrm{~m}$ (avg: $175 \mu \mathrm{~m}$ ); minor am Si $\pm \mathrm{cl}$ <br> very porous; pores ranging from $20-650 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ); areas of high am $\mathrm{Si} \pm \mathrm{cl}$ | high <br> high <br> high | $10-1000 \mu \mathrm{~m}(50 \mu \mathrm{~m})$ <br> $10-1000 \mu \mathrm{~m}(50 \mu \mathrm{~m})$ $20-350 \mu \mathrm{~m}(40 \mu \mathrm{~m})$ | (1) <br> (1) <br> (2) |
| $\begin{gathered} \hline A L V 2179-1-1 \\ \text { B2 } \\ \text { A1 } \end{gathered}$ | $\begin{array}{r} 1476.5 \\ 575.9 \end{array}$ | $\begin{aligned} & 43.2 \\ & 36.9 \end{aligned}$ | $\begin{aligned} & \mathrm{cp}, \mathrm{py}, \mathrm{sp} \\ & \mathrm{cp}, \mathrm{py}, \mathrm{sp} \end{aligned}$ | moderate; $10-100 \mu \mathrm{~m}$ (avg: $50 \mu \mathrm{~m}$ ) sized sulfide crystals <br> moderate; $<50 \mu \mathrm{~m}$ (avg: $30 \mu \mathrm{~m}$ ) sized sulfide crystals; well sorted | pores ranging from 30-200 $\mu \mathrm{m}$ (avg: $60 \mu \mathrm{~m}$ ); remnant worm tubes <br> pores ranging from 30-200 $\mu \mathrm{m}$ (avg: $60 \mu \mathrm{~m}$ ); remnant worm tubes | high <br> high | $10-150 \mu \mathrm{~m}(60 \mu \mathrm{~m})$ $<30 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ | (1) (3) |
| ${ }^{(1)}$ parallel-to-layering core thin section cut radially transects different layers |  |  |  | (3) <br> perpendicular-to-layering core thin section cut radially transects different layers consists of a single layer |  |  |  |  |

## A4.3 Zn-Rich Actively Diffusing Spire Data

| Sample / Section | $\begin{gathered} k \\ \left(\times 10^{-15} \mathrm{~m}^{2}\right) \end{gathered}$ | $\begin{gathered} \phi \\ (\%) \end{gathered}$ | Minerals Present | Grain Packing \& Size | Pore Size | Pore Connectivity | Channel Width min-max (avg) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{aligned} & \text { ALV } 2187-1-1 \text { (top) } \\ & \text { A2 } \\ & \text { A4 } \\ & \text { B1 } \end{aligned}$ | $\begin{gathered} 5462.3 \\ 3433.7 \\ 596.3 \end{gathered}$ | $\begin{aligned} & 45.5 \\ & 36.5 \\ & 45.5 \end{aligned}$ | sp (cp, co) <br> sp (cp) <br> sp (cp) | loose; sulfide crystals $10-300 \mu \mathrm{~m}$ (avg: $120 \mu \mathrm{~m}$ ) <br> loose; sulfide crystals $10-250 \mu \mathrm{~m}$ (avg: $120 \mu \mathrm{~m}$ ) <br> loose; sulfide crystals $10-200 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m})$ | pores $10-600 \mu \mathrm{~m}$ (avg: $140 \mu \mathrm{~m}$ ) pores $10-600 \mu \mathrm{~m}$ (avg: $140 \mu \mathrm{~m}$ ) pores $10-600 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) | high <br> high <br> high | $10-300 \mu \mathrm{~m}(100 \mu \mathrm{~m})$ <br> $10-250 \mu \mathrm{~m}(100 \mu \mathrm{~m})$ <br> $10-150 \mu \mathrm{~m}(80 \mu \mathrm{~m})$ |
| $\begin{gathered} A L V 2187-1-2 \\ \mathrm{C} 1 / \mathrm{C} 2 \\ \\ \mathrm{C} 5 \end{gathered}$ | $\begin{aligned} & 1619.4 \\ & 1426.3 \end{aligned}$ | $\begin{aligned} & 39.4 \\ & 42.9 \end{aligned}$ | $\begin{aligned} & \mathrm{sp}(\mathrm{cp}) \\ & \mathrm{sp}(\mathrm{cp}) \end{aligned}$ | loose; sulfide crystals $10-200 \mu \mathrm{~m}$ (avg: $130 \mu \mathrm{~m})$ <br> loose; sulfide crystals $10-200 \mu \mathrm{~m}$ (avg: $130 \mu \mathrm{~m})$ | pores $10-650 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) <br> pores $10-650 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) | high <br> high | $\begin{aligned} & 10-150 \mu \mathrm{~m}(70 \mu \mathrm{~m}) \\ & 10-150 \mu \mathrm{~m}(70 \mu \mathrm{~m}) \end{aligned}$ |
| $\begin{gathered} A L V 2190-14-1 \\ \text { A1 } \end{gathered}$ | 2726.3 | 47.5 | cp, sp | loose; sulfide crystals $10-120 \mu \mathrm{~m}$ (avg: $75 \mu \mathrm{~m})$ | pores $10-450 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) | high | 10-140 $\mu \mathrm{m}(70 \mu \mathrm{~m})$ |
| $\begin{gathered} J 2-128-8-R 1 \\ E x \end{gathered}$ | 22.2 | 28.0 | cp, wz, am Si | mod-tight; sulfide crystals 5-300 $\mu \mathrm{m}$ (avg: $100 \mu \mathrm{~m}$ ); 1 layer with lots of am Si | pores 10-450 $\mu \mathrm{m}$ (avg: $120 \mu \mathrm{~m}$ ) | moderate (low in 1 layer) | $\begin{aligned} & 10-150 \mu \mathrm{~m} \\ & (10-40 \mu \mathrm{~m}) \end{aligned}$ |
| $\begin{gathered} J 2-222-1-R 1 \\ \mathrm{~A} 2 \\ \mathrm{Ex} \end{gathered}$ | $\begin{array}{r} 327.3 \\ 202.9 \end{array}$ | $\begin{aligned} & 35.0 \\ & 29.8 \end{aligned}$ | $\mathrm{sp}, \mathrm{anh}, \mathrm{cp}, \mathrm{ch}(\mathrm{co})$ <br> $\mathrm{sp}, \mathrm{anh}, \mathrm{cp}, \mathrm{ch}(\mathrm{co})$ | moderate; sulfides $10-300 \mu \mathrm{~m}$ (avg: 100 $\mu \mathrm{m}$ ); anh $\sim 150 \mu \mathrm{~m}$ <br> moderate; sulfides $10-400 \mu \mathrm{~m}$ (avg: 150 $\mu \mathrm{m}$ ); anh $\sim 200 \mu \mathrm{~m}$ | pores $10-550 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) <br> pores $10-600 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) | moderate <br> moderate | $\begin{aligned} & 10-120 \mu \mathrm{~m}(60 \mu \mathrm{~m}) \\ & 10-120 \mu \mathrm{~m}(60 \mu \mathrm{~m}) \end{aligned}$ |
| $\begin{gathered} J 2-127-1-R 2 \\ \text { B3 } \end{gathered}$ | 438.2 | 36.2 | cp, wz | moderate; sulfide crystals $10-300 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) | pores 10-350 $\mu \mathrm{m}$ (avg: $120 \mu \mathrm{~m}$ ) | moderate | 10-150 $\mu \mathrm{m}(45 \mu \mathrm{~m})$ |

## A4.4 Black Smoker Chimney Data

| Sample / Section | $\stackrel{k}{\left(\times 10^{-15} \mathrm{~m}^{2}\right)}$ | $\begin{gathered} \phi \\ (\%) \end{gathered}$ | Minerals Present | Grain Packing \& Size | Pore Size | Pore Connectivity | Channel Width min-max (avg) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{gathered} J 2-137-1-R 1 \\ \text { D1 } \end{gathered}$ | 0.6 | 16.9 | cp, anh, wz, py | tight; sulfide crystals $10-200 \mu \mathrm{~m}$ (avg: $90 \mu \mathrm{~m}$ ); anh $50-300 \mu \mathrm{~m}$ (avg: $120 \mu \mathrm{~m}$ ) | pores 10-200 $\mu \mathrm{m}$ (avg: $75 \mu \mathrm{~m}$ ) | low | $<20 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ |
| $\begin{gathered} \hline J 2-213-3-R 1 \\ \text { A1 } \\ \text { D1 } \end{gathered}$ | 18.4 <br> 99.5 | $\begin{aligned} & 28.9 \\ & 39.0 \end{aligned}$ | cp, anh, wz (py) <br> cp, anh, wz (co, py) | tight; sulfide crystals $10-150 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ); anh $50-350 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) <br> moderate; sulfides $10-250 \mu \mathrm{~m}$ (avg: 120 $\mu \mathrm{m}$ ); anh $50-350 \mu \mathrm{~m}$ (avg: $200 \mu \mathrm{~m}$ ) | pores $10-150 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) pores $10-300 \mu \mathrm{~m}$ (avg: $175 \mu \mathrm{~m}$ ) | low <br> moderate | $10-30 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ <br> $10-150 \mu \mathrm{~m}(40 \mu \mathrm{~m})$ |
| $\begin{gathered} \text { ALV 1445-3 } \\ \mathrm{C} 1 \end{gathered}$ | 16.1 | 24.9 | cp, anh, py | tight; sulfide crystals $10-150 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ); anh $50-300 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) | pores 10-200 $\mu \mathrm{m}$ (avg: $75 \mu \mathrm{~m}$ ) | low | $<20 \mu \mathrm{~m}$ (10 $\mu \mathrm{m}$ ) |

## A4.5 Relict Spire Data

| Sample / Section | $\begin{gathered} k \\ \left(\times 10^{-15} \mathrm{~m}^{2}\right) \end{gathered}$ | $\begin{gathered} \phi \\ (\%) \\ \hline \end{gathered}$ | Minerals Present | Grain Packing \& Size | Pore Size | Pore Connectivity | Channel Width min-max (avg) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{gathered} J 2-129-1-R 3 \\ \mathrm{~B} 2 \end{gathered}$ | 740.4 | 41.0 | $\mathrm{wz}, \mathrm{cp}, \mathrm{am} \mathrm{Si}, \mathrm{anh}$ (py) | loose-mod; most sulfide crystals between $5-200 \mu \mathrm{~m}$ (avg: $60 \mu \mathrm{~m}$ ) | most pores $10-350 \mu \mathrm{~m}$ (avg: $130 \mu \mathrm{~m}$ ); lots of am Si in pore space around smaller grains | moderate | 10-120 $\mu \mathrm{m}(50 \mu \mathrm{~m})$ |
| $\begin{gathered} \hline \begin{array}{c} J 2-136-6-R 1 \\ \mathrm{~A} 1-2 \end{array} \\ \mathrm{C} 2 \end{gathered}$ | $\begin{array}{r} 190.8 \\ 19.4 \end{array}$ | $\begin{aligned} & 30.2 \\ & 25.4 \end{aligned}$ | $\mathrm{wz}, \mathrm{cp}, \mathrm{py}$ <br> $\mathrm{wz}, \mathrm{cp}$ (py) | moderate; sulfide crystals $10-125 \mu \mathrm{~m}$ (avg: $70 \mu \mathrm{~m}$ ) <br> mod-tight; sulfide crystals $10-120 \mu \mathrm{~m}$ (avg: $70 \mu \mathrm{~m}$ ) | pores $10-350 \mu \mathrm{~m}$ (avg: $130 \mu \mathrm{~m}$ ) <br> pores $10-240 \mu \mathrm{~m}$ (avg: $60 \mu \mathrm{~m}$ ); lots of am Si in pore space around smaller grains | moderate <br> low | $\begin{aligned} & 10-100 \mu \mathrm{~m}(50 \mu \mathrm{~m}) \\ & 10-80 \mu \mathrm{~m}(30 \mu \mathrm{~m}) \end{aligned}$ |
| $\begin{aligned} & \text { ALV 2944-3-Sl, Pc } 1 \\ & \text { A1 } \\ & \text { A3 } \end{aligned}$ | $\begin{array}{r} 626.4 \\ 40.9 \end{array}$ | 37.6 30.9 | $\mathrm{cp}, \mathrm{wz}, \mathrm{py}(\mathrm{am} \mathrm{Si})$ <br> cp, wz, py, am Si, anh | moderate; sulfide crystals $5-130 \mu \mathrm{~m}$ (avg: $40 \mu \mathrm{~m}$ ) <br> tight; sulfide crystals $10-100 \mu \mathrm{~m}$ (avg: $50 \mu \mathrm{~m}$ ) | pores $10-400 \mu \mathrm{~m}$ (avg: $150 \mu \mathrm{~m}$ ) pores $10-350 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) | high <br> low | $10-120 \mu \mathrm{~m}(50 \mu \mathrm{~m})$ <br> $10-100 \mu \mathrm{~m}(30 \mu \mathrm{~m})$ |
| $\begin{aligned} & \text { ALV 2944-3-S1, Pc } 2 \\ & \text { A1 } \\ & \text { A2 } \end{aligned}$ | $\begin{aligned} & 327.2 \\ & 645.8 \end{aligned}$ | $\begin{aligned} & 33.5 \\ & 42.6 \end{aligned}$ | $\begin{aligned} & \text { cp, wz, py } \\ & \text { cp, wz, py } \end{aligned}$ | loose; sulfide crystals $10-180 \mu \mathrm{~m}$ (avg: <br> $90 \mu \mathrm{~m})$ <br> loose; sulfide crystals $10-130 \mu \mathrm{~m}$ (avg: <br> $80 \mu \mathrm{~m})$ | pores $10-600 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) pores $10-250 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) | high <br> high | $\begin{aligned} & 10-200 \mu \mathrm{~m}(60 \mu \mathrm{~m}) \\ & 10-80 \mu \mathrm{~m}(60 \mu \mathrm{~m}) \end{aligned}$ |
| $\begin{aligned} & \text { ALV 2941-6-S1 } \\ & \quad \text { A1 } \end{aligned}$ | 140.6 | 32.5 | cp, wz (py) | loose; sulfide crystals $10-275 \mu \mathrm{~m}$ (avg: <br> $125 \mu \mathrm{~m})$ | pores $10-550 \mu \mathrm{~m}$ (avg: $140 \mu \mathrm{~m}$ ) | high | 10-150 $\mu \mathrm{m}$ (40 $\mu \mathrm{m}$ ) |
| $\begin{gathered} J 2-125-3-B 1 \\ \text { B2-1 } \\ \text { B3-1 } \\ \text { Ex } \end{gathered}$ | $\begin{array}{r} 356.3 \\ 97.0 \\ 502.5 \end{array}$ | 34.7 32.6 37.8 | cp (wz, py) <br> $\mathrm{cp}, \mathrm{wz}, \mathrm{py}$ <br> cp, wz, py | loose; sulfide crystals $10-175 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m})$ <br> loose-mod; sulfide crystals 5-150 $\mu \mathrm{m}$ (avg: $70 \mu \mathrm{~m}$ ) <br> loose-mod; sulfide crystals $10-150 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ) | pores $10-350 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) <br> pores $10-450 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) <br> pores $10-500 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) | high <br> moderate <br> mod -high | $10-175 \mu \mathrm{~m}(60 \mu \mathrm{~m})$ <br> $10-125 \mu \mathrm{~m}(40 \mu \mathrm{~m})$ <br> $10-250 \mu \mathrm{~m}(40 \mu \mathrm{~m})$ |


| $\begin{gathered} \text { ALV } 2178-4-1 \\ 4 \\ \text { A5 } \end{gathered}$ | $\begin{gathered} 2224.3 \\ 766.9 \end{gathered}$ | $\begin{aligned} & 49.3 \\ & 44.9 \end{aligned}$ | $\begin{aligned} & \text { py, cp (wz) } \\ & \text { py, cp (wz) } \end{aligned}$ | loose; sulfide crystals $10-250 \mu \mathrm{~m}$ (avg: $120 \mu \mathrm{~m})$ <br> loose; sulfide crystals $10-350 \mu \mathrm{~m}$ (avg: $120 \mu \mathrm{~m})$ | pores $10-600 \mu \mathrm{~m}$ (avg: $120 \mu \mathrm{~m}$ ) pores $10-400 \mu \mathrm{~m}$ (avg: $70 \mu \mathrm{~m}$ ) | high <br> high | $\begin{aligned} & 10-250 \mu \mathrm{~m}(70 \mu \mathrm{~m}) \\ & 10-130 \mu \mathrm{~m}(70 \mu \mathrm{~m}) \end{aligned}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{gathered} \text { ALV } 2461-R 13 \\ \text { C1-1 } \end{gathered}$ | 822.4 | 26.7 | cp, wz, py (am Si) | moderate; sulfide crystals $10-150 \mu \mathrm{~m}$ (avg: $70 \mu \mathrm{~m}$ ) | pores 10-350 $\mu \mathrm{m}$ (avg: $120 \mu \mathrm{~m}$ ) | moderate | 10-160 $\mu \mathrm{m}(40 \mu \mathrm{~m})$ |
| C1-2 | 1142.5 | 22.2 | cp, wz, py, po, am Si | moderate; sulfide crystals $10-300 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) | pores $10-450 \mu \mathrm{~m}$ (avg: $120 \mu \mathrm{~m}$ ) | moderate | $10-140 \mu \mathrm{~m}(50 \mu \mathrm{~m})$ |
| C3-2(1-2) | 0.3 | 10.4 | cp, wz, py, po, am Si | tight; sulfide crystals $10-320 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m})$ | pores $10-400 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ) | low | $10-50 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ |
| C3-2(2) | 6.4 | 12.2 | cp, wz, po, am Si (py) | tight; sulfide crystals 10-275 $\mu \mathrm{m}$ (avg: $80 \mu \mathrm{~m}$ ) | pores $10-400 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ) | low | $10-50 \mu \mathrm{~m}(10 \mu \mathrm{~m})$ |
| C4-1 | 6.6 | 18.7 | cp, wz, am Si (py, po) | tight; sulfide crystals $10-150 \mu \mathrm{~m}$ (avg: $75 \mu \mathrm{~m}$ ) | pores $10-500 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ) | low | $10-60 \mu \mathrm{~m}(15 \mu \mathrm{~m})$ |
| C4-2 | 4.4 | 22.0 | cp, wz, py, am Si (po) | tight; sulfide crystals $10-120 \mu \mathrm{~m}$ (avg: $70 \mu \mathrm{~m})$ | pores $10-500 \mu \mathrm{~m}$ (avg: $80 \mu \mathrm{~m}$ ) | low | $10-60 \mu \mathrm{~m}(15 \mu \mathrm{~m})$ |
| C8-2 | 112.6 | 20.7 | cp, wz, po,(py, am Si) | moderate; sulfide crystals $10-175 \mu \mathrm{~m}$ (avg: $100 \mu \mathrm{~m}$ ) | pores $10-550 \mu \mathrm{~m}$ (avg: $120 \mu \mathrm{~m}$ ) | moderate | 10-250 $\mu \mathrm{m}(20 \mu \mathrm{~m})$ |

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